

A fully well-balanced scheme for the shallow-water equations with topography

C. Berthon¹, S. Clain², F. Foucher^{1,3}, V. Michel-Dansac⁴

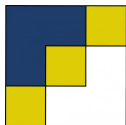
¹Laboratoire de Mathématiques Jean Leray, Université de Nantes

²Centre of Mathematics, Minho University

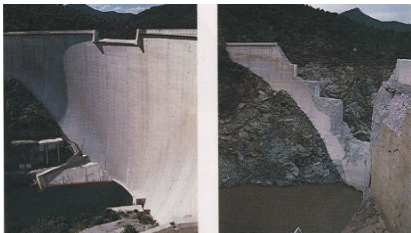
³École Centrale de Nantes

⁴Institut de Mathématiques de Toulouse

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Several kinds of destructive geophysical flows



Dam failure (Malpasset, France, 1959)



Tsunami (Tōhoku, Japan, 2011)



Flood (La Faute sur Mer, France, 2010)

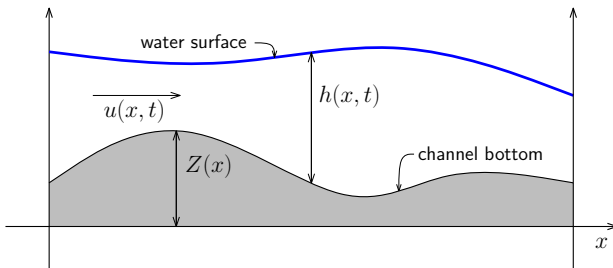


Mudslide (Madeira, Portugal, 2010)

The shallow-water equations

$$\begin{cases} \partial_t h + \partial_x(hu) = 0 \\ \partial_t(hu) + \partial_x\left(hu^2 + \frac{1}{2}gh^2\right) = -gh\partial_x Z \end{cases}$$

We can rewrite the equations as $\partial_t W + \partial_x F(W) = S(W)$, with $W = \begin{pmatrix} h \\ q \end{pmatrix}$.



- $Z(x)$ is the known topography
- g is the gravitational constant
- we label the water discharge $q := hu$

Steady state solutions

Definition: Steady state solutions

W is a steady state solution iff $\partial_t W = 0$, i.e. $\partial_x F(W) = S(W)$.

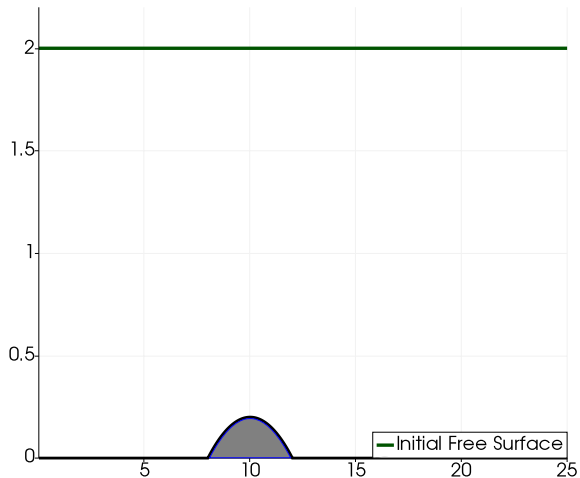
Taking $\partial_t W = 0$ in the shallow-water equations leads to

$$\begin{cases} \partial_x q = 0 \\ \partial_x \left(\frac{q^2}{h} + \frac{1}{2}gh^2 \right) = -gh\partial_x Z. \end{cases}$$

The **smooth** steady state solutions are therefore given by the following statement of Bernoulli's principle:

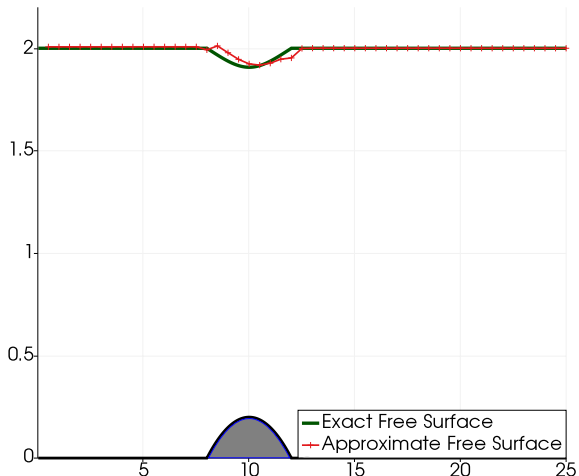
$$\begin{cases} q = \text{cst} = q_0 \\ \frac{q_0^2}{2} \partial_x \left(\frac{1}{h^2} \right) + g\partial_x (h + Z) = 0. \end{cases}$$

Steady state not captured in 1D



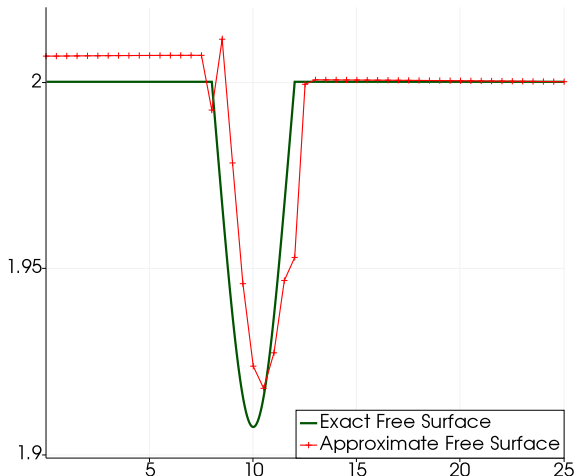
The initial condition is at rest; water is injected through the left boundary.

Steady state not captured in 1D



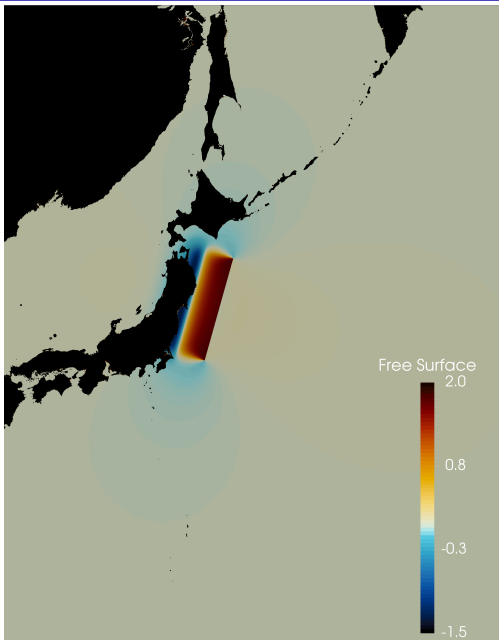
The classical HLL numerical scheme converges towards a **numerical** steady state which does not correspond to the **physical** one.

Steady state not captured in 1D



The classical HLL numerical scheme converges towards a **numerical** steady state which does not correspond to the **physical** one.

A real-life simulation:
the 2011 Tōhoku
tsunami. The water is
close to a steady state
at rest far from the
tsunami.



Objectives

Our goal is to derive a **numerical method** for the shallow-water model that **exactly preserves** its **stationary solutions** on every mesh.

To that end, we seek a numerical scheme that:

- 1** is **well-balanced** for the shallow-water equations with topography, i.e. it exactly preserves and captures the steady states without having to solve the governing nonlinear equation;
- 2** preserves the **non-negativity** of the water height;
- 3** can be easily extended for **other source terms** of the shallow-water equations (e.g. friction or breadth).

Contents

- 1 Brief introduction to Godunov-type schemes
- 2 Derivation of a well-balanced scheme
- 3 Numerical simulations
- 4 Conclusion and perspectives

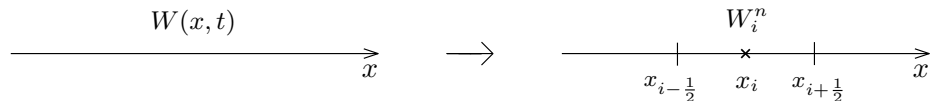
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Setting

Objective: Approximate the solution $W(x, t)$ of the hyperbolic system $\partial_t W + \partial_x F(W) = S(W)$, with suitable initial and boundary conditions.

We partition $[a, b]$ in *cells*, of volume Δx and of evenly spaced centers x_i , and we define:

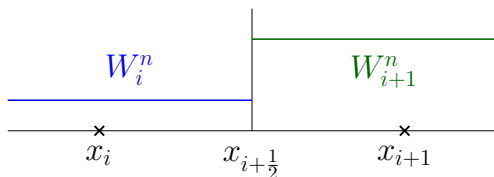
- $x_{i-\frac{1}{2}}$ and $x_{i+\frac{1}{2}}$, the boundaries of the cell i ;
- W_i^n , an approximation of $W(x, t)$, constant in the cell i and at time t^n , which is defined as $W_i^n = \frac{1}{\Delta x} \int_{\Delta x/2}^{\Delta x/2} W(x, t^n) dx$.



Using an approximate Riemann solver

As a consequence, at time t^n , we have a succession of Riemann problems (Cauchy problems with discontinuous initial data) at the interfaces between cells:

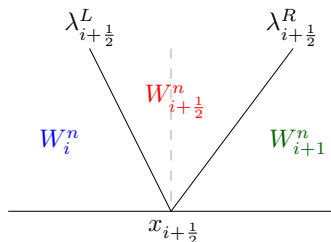
$$\begin{cases} \partial_t W + \partial_x F(W) = S(W) \\ W(x, t^n) = \begin{cases} W_i^n & \text{if } x < x_{i+\frac{1}{2}} \\ W_{i+1}^n & \text{if } x > x_{i+\frac{1}{2}} \end{cases} \end{cases}$$



For $S(W) \neq 0$, the exact solution to these Riemann problems is unknown or costly to compute \rightsquigarrow we require an approximation.

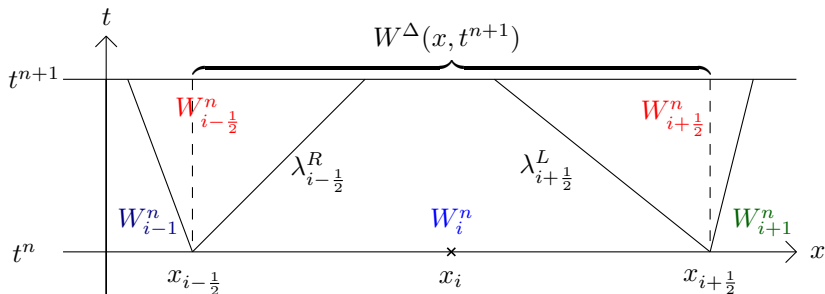
Using an approximate Riemann solver

We choose to use an approximate Riemann solver, as follows.



- $W_{i+\frac{1}{2}}^n$ is an approximation of the interaction between W_i^n and W_{i+1}^n (i.e. of the solution to the Riemann problem), possibly made of several constant states separated by discontinuities.
- $\lambda_{i+\frac{1}{2}}^L$ and $\lambda_{i+\frac{1}{2}}^R$ are approximations of the wave speeds.

Godunov-type scheme (approximate Riemann solver)



We define the time update as follows :

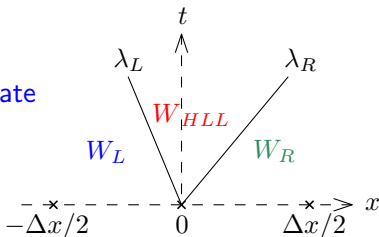
$$W_i^{n+1} := \frac{1}{\Delta x} \int_{x_{i-1/2}}^{x_{i+1/2}} W^\Delta(x, t^{n+1}) dx.$$

Since $W_{i-1/2}^n$ and $W_{i+1/2}^n$ are made of constant states, the above integral is easy to compute.

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The HLL approximate Riemann solver

To approximate solutions of $\partial_t W + \partial_x F(W) = 0$, the **HLL approximate Riemann solver** (Harten, Lax, van Leer (1983)) may be chosen; it is denoted by W^Δ and displayed on the right.



The consistency condition (as per Harten and Lax) holds if:

$$\frac{1}{\Delta x} \int_{-\Delta x/2}^{\Delta x/2} W^\Delta(\Delta t, x; W_L, W_R) dx = \frac{1}{\Delta x} \int_{-\Delta x/2}^{\Delta x/2} W_R(\Delta t, x; W_L, W_R) dx,$$

$$\text{which gives } W_{HLL} = \frac{\lambda_R W_R - \lambda_L W_L}{\lambda_R - \lambda_L} - \frac{F(W_R) - F(W_L)}{\lambda_R - \lambda_L} = \begin{pmatrix} h_{HLL} \\ q_{HLL} \end{pmatrix}.$$

Note that, if $h_L > 0$ and $h_R > 0$, then $h_{HLL} > 0$ for $|\lambda_L|$ and $|\lambda_R|$ large enough.

Modification of the HLL approximate Riemann solver

The shallow-water equations with the topography source term read as follows:

$$\begin{cases} \partial_t h + \partial_x q = 0, \\ \partial_t q + \partial_x \left(\frac{q^2}{h} + \frac{1}{2} g h^2 \right) + g h \partial_x Z = 0. \end{cases}$$

Modification of the HLL approximate Riemann solver

We can add the equation $\partial_t Z = 0$, which corresponds to the fixed geometry of the problem:

$$\begin{cases} \partial_t h + \partial_x q = 0, \\ \partial_t q + \partial_x \left(\frac{q^2}{h} + \frac{1}{2} g h^2 \right) + g h \partial_x Z = 0, \\ \partial_t Z = 0. \end{cases}$$

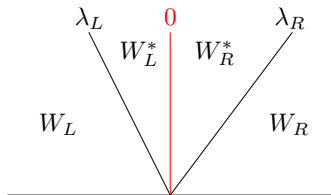
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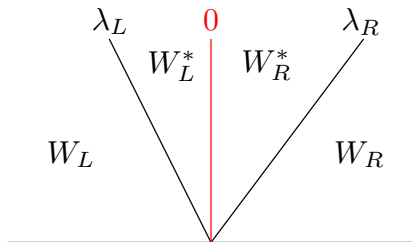
The equation $\partial_t Z = 0$ induces a **stationary wave** associated to the source term; we also note that q is a Riemann invariant for this wave.

To approximate solutions of $\partial_t W + \partial_x F(W) = S(W)$, we thus use the approximate Riemann solver displayed on the right (assuming $\lambda_L < 0 < \lambda_R$).



Modification of the HLL approximate Riemann solver

We have 4 unknowns to determine: $W_L^* = \begin{pmatrix} h_L^* \\ q_L^* \end{pmatrix}$ and $W_R^* = \begin{pmatrix} h_R^* \\ q_R^* \end{pmatrix}$.



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- $q^* = q_{HLL} + \frac{\bar{S} \Delta x}{\lambda_R - \lambda_L}$ (relation 3),

where $\bar{S} \simeq \frac{1}{\Delta x} \frac{1}{\Delta t} \int_{-\Delta x/2}^{\Delta x/2} \int_0^{\Delta t} S(W_{\mathcal{R}}(x, t)) dt dx$

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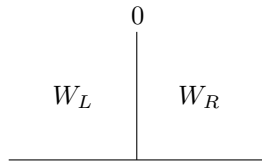
- next step: obtain a fourth relation

Obtaining an additional relation

Assume that W_L and W_R define a steady state, i.e. that they satisfy the following discrete version of the steady relation $\partial_x F(W) = S(W)$ (where $[X] = X_R - X_L$):

$$\frac{1}{\Delta x} \left(q_0^2 \left[\frac{1}{h} \right] + \frac{g}{2} [h^2] \right) = \bar{S}.$$

For the steady state to be preserved, it is sufficient to have $h_L^* = h_L$, $h_R^* = h_R$ and $q^* = q_0$.



Assuming a steady state, we easily show that $q^* = q_0$. Therefore, the additional relation should only link h_L^* and h_R^* .

Obtaining an additional relation

In order to determine an additional relation, we consider the discrete steady relation, satisfied when W_L and W_R define a steady state:

$$q_0^2 \left(\frac{1}{h_R} - \frac{1}{h_L} \right) + \frac{g}{2} ((h_R)^2 - (h_L)^2) = \bar{S} \Delta x.$$

To ensure that $h_L^* = h_L$ and $h_R^* = h_R$, we impose that h_L^* and h_R^* satisfy the above relation, as follows:

$$q_0^2 \left(\frac{1}{h_R^*} - \frac{1}{h_L^*} \right) + \frac{g}{2} ((h_R^*)^2 - (h_L^*)^2) = \bar{S} \Delta x.$$

To avoid solving a nonlinear system, we elect to use the following linearization of this relation:

$$\underbrace{\left(\frac{-(q^*)^2}{h_L h_R} + \frac{g}{2} (h_L + h_R) \right)}_{\alpha} (h_R^* - h_L^*) = \bar{S} \Delta x.$$

Determination of h_L^* and h_R^*

With the consistency relation between h_L^* and h_R^* , the intermediate water heights satisfy the following linear system:

$$\begin{cases} \alpha(h_R^* - h_L^*) = \bar{S}\Delta x, \\ \lambda_R h_R^* - \lambda_L h_L^* = (\lambda_R - \lambda_L)h_{HLL}. \end{cases}$$

Using both relations linking h_L^* and h_R^* , we obtain

$$\begin{cases} h_L^* = h_{HLL} - \frac{\lambda_R \bar{S} \Delta x}{\alpha(\lambda_R - \lambda_L)}, \\ h_R^* = h_{HLL} - \frac{\lambda_L \bar{S} \Delta x}{\alpha(\lambda_R - \lambda_L)}, \end{cases}$$

where $\alpha = \left(\frac{-(q^*)^2}{h_L h_R} + \frac{g}{2}(h_L + h_R) \right)$ with $q^* = q_{HLL} + \frac{\bar{S} \Delta x}{\lambda_R - \lambda_L}$.

Correction to ensure non-negative h_L^* and h_R^*

However, these expressions of h_L^* and h_R^* do not guarantee that the intermediate heights are non-negative: instead, we use the following cutoff (see Audusse, Chalons, Ung (2014)):

$$\begin{cases} h_L^* = \min\left(\left(h_{HLL} - \frac{\lambda_R \bar{S} \Delta x}{\alpha(\lambda_R - \lambda_L)}\right)_+, \left(1 - \frac{\lambda_R}{\lambda_L}\right) h_{HLL}\right), \\ h_R^* = \min\left(\left(h_{HLL} - \frac{\lambda_L \bar{S} \Delta x}{\alpha(\lambda_R - \lambda_L)}\right)_+, \left(1 - \frac{\lambda_L}{\lambda_R}\right) h_{HLL}\right). \end{cases}$$

Note that this cutoff does not interfere with:

- the consistency condition $\lambda_R h_R^* - \lambda_L h_L^* = (\lambda_R - \lambda_L) h_{HLL}$;
- the well-balance property, since it is not activated when W_L and W_R define a steady state.

Summary

The two-state approximate Riemann solver with intermediate states

$$W_L^* = \begin{pmatrix} h_L^* \\ q^* \end{pmatrix} \text{ and } W_R^* = \begin{pmatrix} h_R^* \\ q^* \end{pmatrix} \text{ given by}$$

$$\begin{cases} q^* = q_{HLL} + \frac{\bar{S}\Delta x}{\lambda_R - \lambda_L}, \\ h_L^* = \min\left(\left(h_{HLL} - \frac{\lambda_R \bar{S}\Delta x}{\alpha(\lambda_R - \lambda_L)}\right)_+, \left(1 - \frac{\lambda_R}{\lambda_L}\right)h_{HLL}\right), \\ h_R^* = \min\left(\left(h_{HLL} - \frac{\lambda_L \bar{S}\Delta x}{\alpha(\lambda_R - \lambda_L)}\right)_+, \left(1 - \frac{\lambda_L}{\lambda_R}\right)h_{HLL}\right), \end{cases}$$

is **consistent**, **non-negativity-preserving** and **well-balanced**.

next step: determination of \bar{S} according to the **source term definition** (topography).

The topography source term

We now consider $S(W) = S^t(W) = -gh\partial_x Z$:
the smooth steady states are governed by

$$\left. \begin{aligned} \partial_x \left(\frac{q_0^2}{h} \right) + \frac{g}{2} \partial_x (h^2) &= -gh\partial_x Z, \\ \frac{q_0^2}{2} \partial_x \left(\frac{1}{h^2} \right) + g\partial_x (h + Z) &= 0, \end{aligned} \right\} \xrightarrow{\text{discretization}} \begin{cases} q_0^2 \left[\frac{1}{h} \right] + \frac{g}{2} [h^2] = \bar{S}^t \Delta x, \\ \frac{q_0^2}{2} \left[\frac{1}{h^2} \right] + g[h + Z] = 0. \end{cases}$$

We can exhibit an expression of q_0^2 and thus obtain

$$\bar{S}^t = -g \frac{2h_L h_R}{h_L + h_R} \frac{[Z]}{\Delta x} + \frac{g}{2\Delta x} \frac{[h]^3}{h_L + h_R}.$$

However, when $Z_L = Z_R$, we have $\bar{S}^t \neq \mathcal{O}(\Delta x)$, i.e. a **loss of consistency with S^t** (see for instance Berthon, Chalons (2016)).

The topography source term

Instead, we set, for some constant $C > 0$,

$$\left\{ \begin{array}{l} \bar{S}^t = -g \frac{2h_L h_R}{h_L + h_R} \frac{[Z]}{\Delta x} + \frac{g}{2\Delta x} \frac{[h]_c^3}{h_L + h_R}, \\ [h]_c = \begin{cases} h_R - h_L & \text{if } |h_R - h_L| \leq C\Delta x, \\ \text{sgn}(h_R - h_L) C\Delta x & \text{otherwise.} \end{cases} \end{array} \right.$$

Theorem: Well-balance for the topography source term

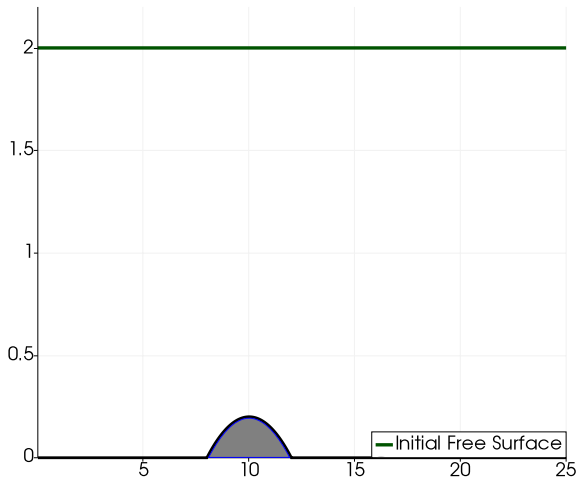
If W_L and W_R define a smooth steady state, i.e. if they satisfy

$$\frac{q_0^2}{2} \left[\frac{1}{h^2} \right] + g[h + Z] = 0,$$

then we have $W_L^* = W_L$ and $W_R^* = W_R$ and the approximate Riemann solver is well-balanced. By construction, the Godunov-type scheme using this approximate Riemann solver is **consistent**, **fully well-balanced** and **positivity-preserving**.

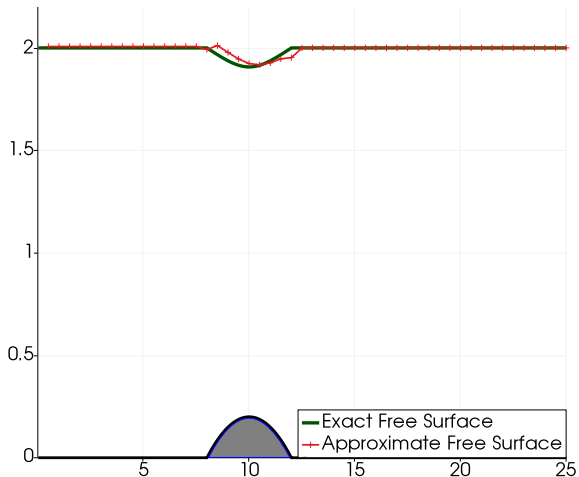
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Verification of the well-balance: topography



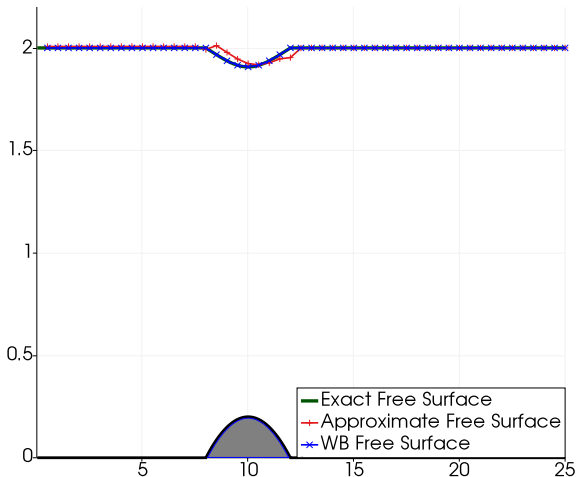
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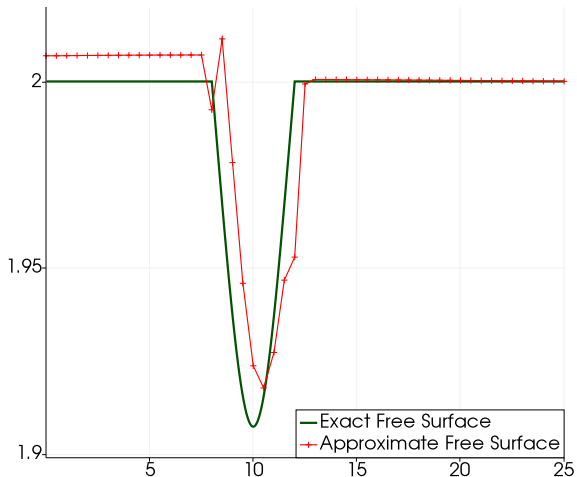
The non-well-balanced HLL scheme converges towards a **numerical** steady state which does not correspond to the **physical** one.

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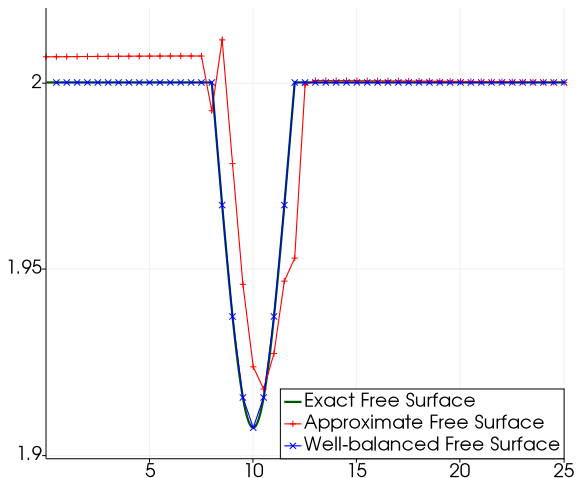
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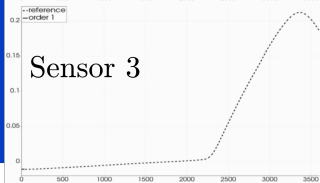
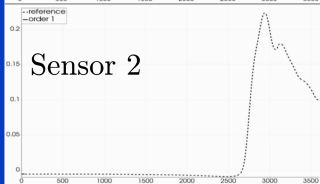
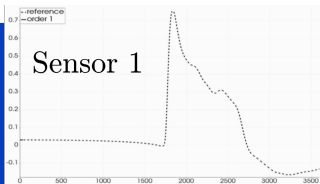
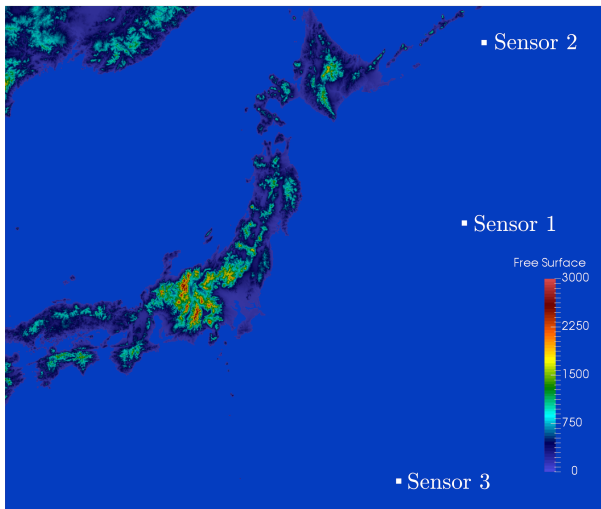
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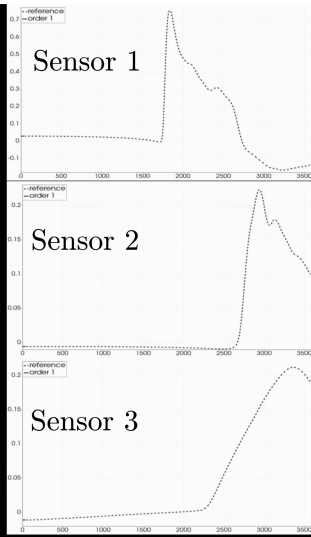
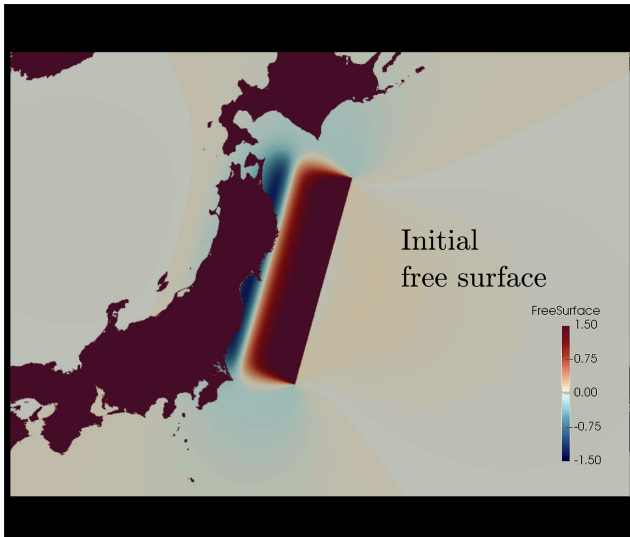


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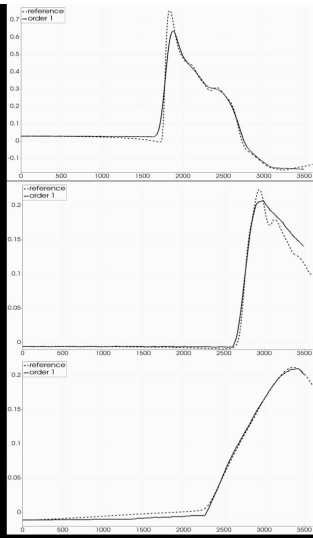
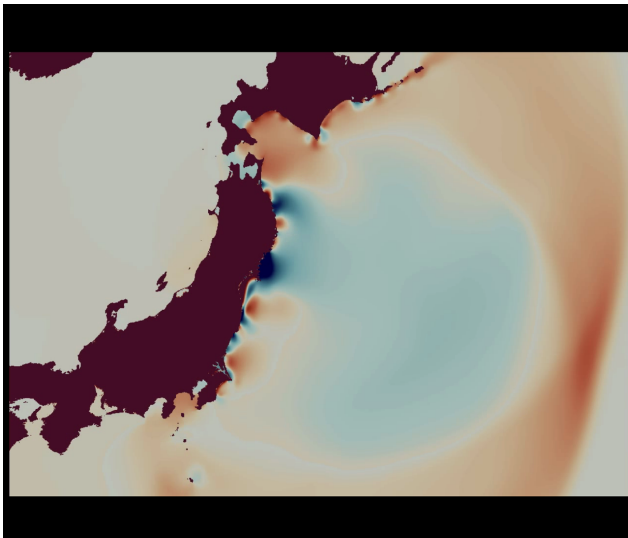
Simulation of the 2011 Tōhoku tsunami



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Conclusion

- We have presented a **well-balanced** and **non-negativity-preserving** numerical scheme, able to be applied to other source terms or combinations of source terms.
- We have also displayed results from a **2D well-balanced** numerical method, coded in Fortran and **parallelized**.

This work has been published:

V. Michel-Dansac et al. “A well-balanced scheme for the shallow-water equations with topography”. In: *Comput. Math. Appl.* 72.3 (2016), pp. 568–593

V. Michel-Dansac et al. “A well-balanced scheme for the shallow-water equations with topography or Manning friction”. In: *J. Comput. Phys.* (*accepted*) (2017)

Perspectives

Work in progress or completed

- application to other source terms:
 - friction source term (completed, article accepted)
 - Coriolis force source term (work in progress)
 - breadth variation source term (work in progress)
- high-order and 2D extensions (work in progress, collaboration with R. Loubère)

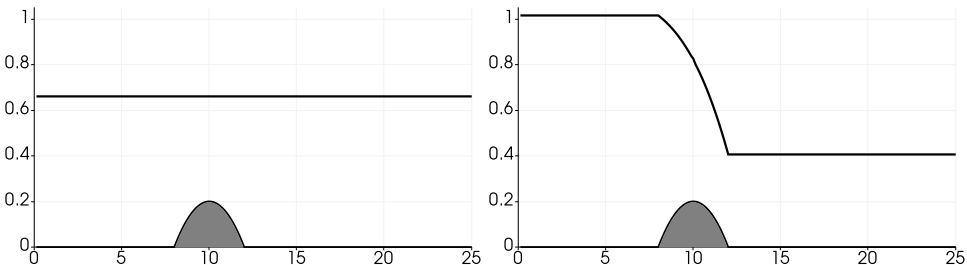
Long-term perspectives

- stability of the scheme: values of C , λ_L and λ_R to ensure the entropy preservation
- ensure the entropy preservation for the high-order scheme (use of a MOOD method)

Thank you for your attention!

Verification of the well-balance: topography

transcritical flow test case (see Goutal, Maurel (1997))



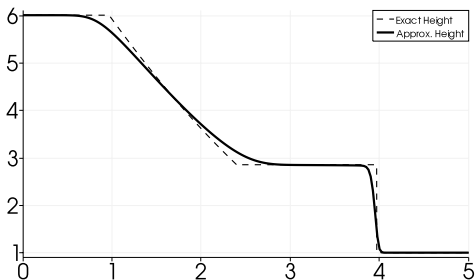
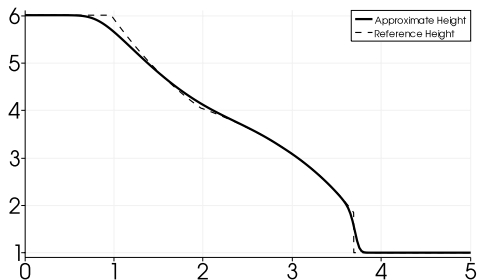
left panel: initial free surface at rest; water is injected from the left boundary

right panel: free surface for the steady state solution, after a transient state

$$\Phi = \frac{u^2}{2} + g(h + Z)$$

	L^1	L^2	L^∞
errors on q	1.47e-14	1.58e-14	2.04e-14
errors on Φ	1.67e-14	2.13e-14	4.26e-14

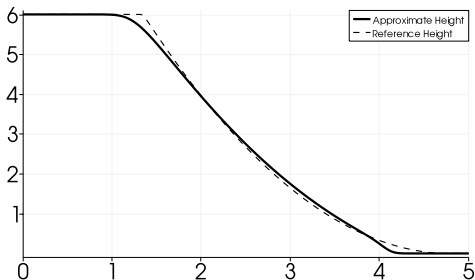
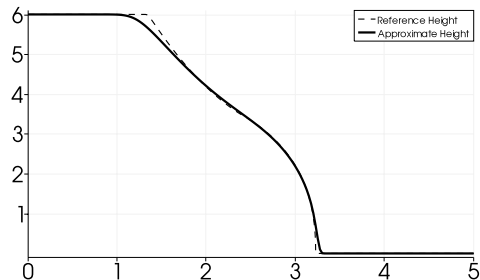
Riemann problems between two wet areas

left: $k = 0$ left: $k = 10$

both Riemann problems have initial data $W_L = \begin{pmatrix} 6 \\ 0 \end{pmatrix}$ and

$W_R = \begin{pmatrix} 1 \\ 0 \end{pmatrix}$, on $[0, 5]$, with 200 points, and final time 0.2s

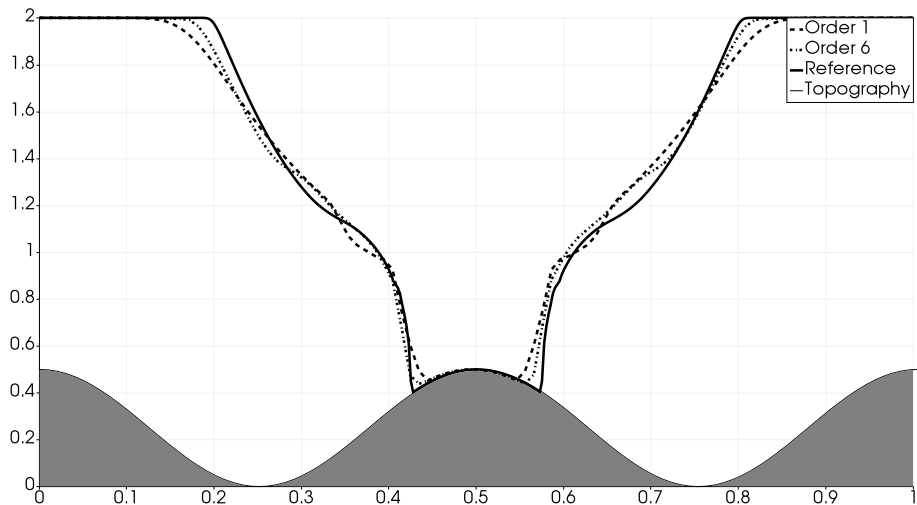
Riemann problems with a wet/dry transition

left: $k = 0$ left: $k = 10$

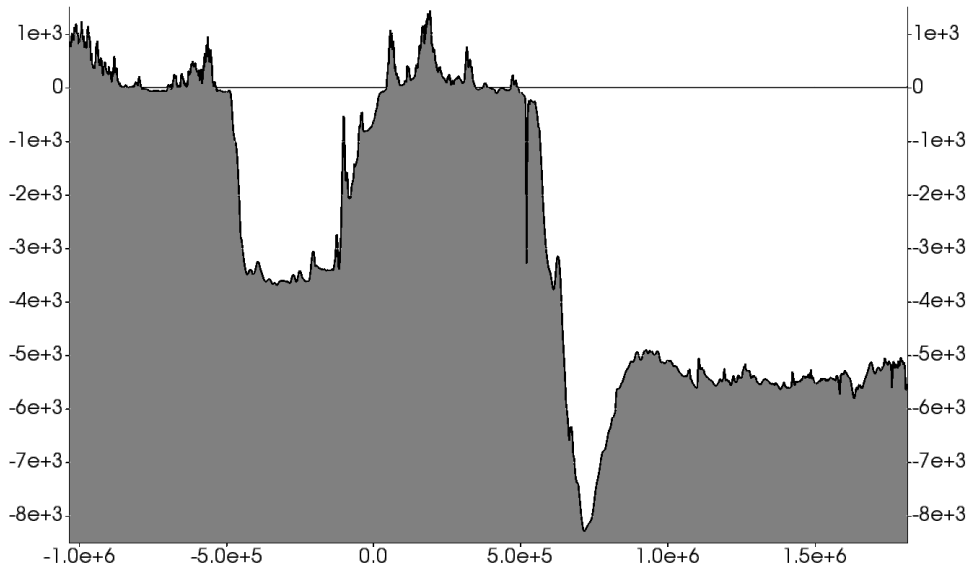
both Riemann problems have initial data $W_L = \begin{pmatrix} 6 \\ 0 \end{pmatrix}$ and

$W_R = \begin{pmatrix} 0 \\ 0 \end{pmatrix}$, on $[0, 5]$, with 200 points, and final time 0.15s

Double dry dam-break on a sinusoidal bottom



Japan tsunami: 1D slice



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