

# Consistent section-averaged shallow water equations with bottom friction

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# Motivation: 2D/1D coupling for estuary simulation



Gironde estuary: satellite picture



Gironde estuary: 2D mesh

Regarding the shape of the river bed, as of now,

- the derivation of 1D models is **well-understood**<sup>1,2</sup> in the ideal case of a **U-shaped channel**;
- for **more complex shapes**, the water surface of uniform stationary flows is recovered<sup>3,4</sup> using a **empiric terms** or **data assimilation**;
- fully 2D models are used but they are computationally costly.

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<sup>1</sup>see Bresch and Noble, 2007, in the context of laminar flows

<sup>2</sup>see Richard, Rambaud and Vila, 2017, in the context of turbulent flows

<sup>3</sup>see Decoene, Bonaventura, Miglio and Saleri, 2009

<sup>4</sup>see Marin and Monnier, 2009

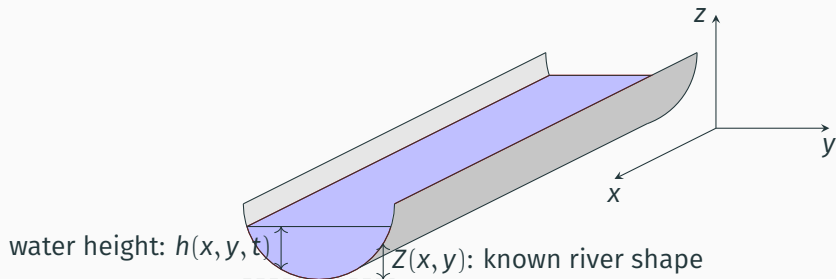
# Specifications of the 1D model

The goal of this work is to **develop a new model**, based on the shallow water equations, that is:

- **generic** enough to not require empiric friction coefficients;
- **consistent** with the 2D shallow water in the **asymptotic regime** corresponding to an **estuary** or a **river**;
- **hyperbolic**;
- **easily implementable** (collaboration with the SHOM for flood simulations, ocean model forcing, ...).

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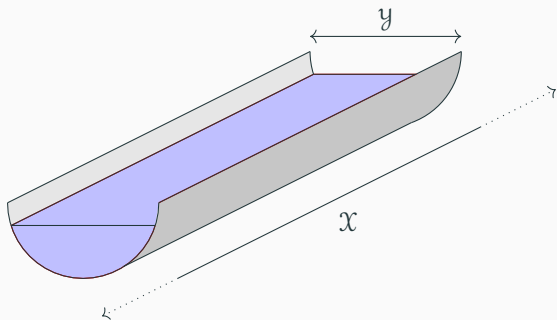
# The non-conservative 2D shallow water system



$$\begin{cases} h_t + \nabla \cdot (h\mathbf{u}) = 0 \\ \mathbf{u}_t + \mathbf{u} \cdot \nabla \mathbf{u} + g\nabla h = g \left( -\nabla Z - \frac{\mathbf{u}\|\mathbf{u}\|}{C_h^2 h^p} \right) \end{cases}$$

- $\mathbf{u} = (u, v)$  is the water velocity
- $g$  is the gravity constant
- $C_h(x, y)$  is the (known) Chézy friction coefficient
- $p = 4/3$  is the friction exponent

# Introduction of reference scales: the coordinates



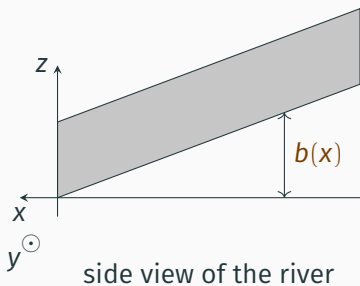
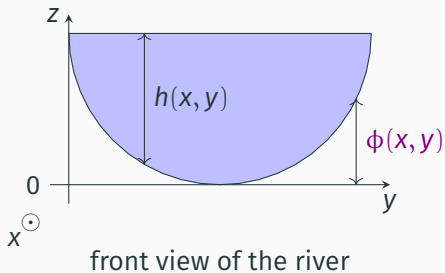
	dimensional quantity	reference scale	non-dimensional quantity
longitudinal coordinates	$x \in (0\text{m}, 60000\text{m})$	$\mathcal{X} = 2000\text{m}$	$\bar{x} = \frac{x}{\mathcal{X}} \in (0, 30)$
transverse coordinates	$y \in (-25\text{m}, 25\text{m})$	$\mathcal{Y} = 50\text{m}$	$\bar{y} = \frac{y}{\mathcal{Y}} \in (-0.5, 0.5)$

# Introduction of reference scales: the topography

Regarding the geometry, we assume that  $Z(x,y) = b(x) + \phi(x,y)$ , where:

- $b(x)$  represents the main **longitudinal topography**, driving the flow from upstream to downstream;
- $\phi(x,y)$  represents **small** longitudinal and transverse **variations**.

Thus,  $h + \phi$  represents the altitude of the water surface.





# Non-dimensional form of the 2D shallow water system

We introduce the following non-dimensional numbers to emphasize the different scales of the flow:

- $F^2$ , the reference Froude number (ratio material/acoustic velocity),
- $\delta$ , the shallow water parameter (ratio height/reference length),
- $R_u$ , the quasi-1D parameter (ratio transverse/longitudinal velocity),
- $l_0$  and  $J_0$ , the reference topography and friction slopes.

Finally, the non-dimensional form of the 2D shallow water system is:

$$\begin{cases} \bar{h}_{\bar{t}} + (\bar{h}\bar{u})_{\bar{x}} + (\bar{h}\bar{v})_{\bar{y}} = 0, \\ \bar{u}_{\bar{t}} + \bar{u}\bar{u}_{\bar{x}} + \bar{v}\bar{u}_{\bar{y}} + \frac{1}{F^2} (\bar{h} + \bar{\Phi})_{\bar{x}} = \frac{1}{\delta F^2} \left( -J_0 \frac{\bar{u}\sqrt{\bar{u}^2 + R_u^2\bar{v}^2}}{\bar{c}^2\bar{h}^p} - l_0\bar{b}_{\bar{x}} \right), \\ \bar{v}_{\bar{t}} + \bar{u}\bar{v}_{\bar{x}} + \bar{v}\bar{v}_{\bar{y}} + \frac{1}{R_u^2 F^2} (\bar{h} + \bar{\Phi})_{\bar{y}} = -\frac{J_0}{\delta F^2} \frac{\bar{v}\sqrt{\bar{u}^2 + R_u^2\bar{v}^2}}{\bar{c}^2\bar{h}^p}. \end{cases}$$

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# Asymptotic expansions setup

In the regime under consideration, we have

- $\varepsilon := \frac{\delta F^2}{J_0} \ll 1$  (in practice,  $F^2 \ll 1$ ,  $\delta \ll 1$ ,  $J_0 \ll 1$  and  $J_0 \sim \delta$ ),
- $R_u \ll 1$  (quasi-unidimensional setting), and  $R_u = \mathcal{O}(\varepsilon)$ .

Highlighting the **dominant terms** in the system, we get:

$$\begin{cases} h_t + (hu)_x + (hv)_y = 0, \\ u_t + uu_x + vu_y + \frac{1}{\varepsilon} \frac{\delta}{J_0} (h + \phi)_x = \frac{1}{\varepsilon} \left( -\frac{u\sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^p} - \frac{l_0}{J_0} b_x \right), \\ v_t + uv_x + vv_y + \frac{1}{\varepsilon^3} \frac{\delta}{J_0} (h + \phi)_y = -\frac{1}{\varepsilon} \frac{v\sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^p}. \end{cases}$$

**Goal:** Perform asymptotic expansions in this regime, to better understand the weak dependency of the solution in  $y$ .

# Free surface expansion

We consider the third equation:

$$v_t + uv_x + vv_y + \frac{1}{\varepsilon^3} \frac{\delta}{J_0} (h + \phi)_y = -\frac{1}{\varepsilon} \frac{v\sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^p},$$

which we rewrite as follows to highlight the dominant term:

$$\frac{\delta}{J_0} (h + \phi)_y = \varepsilon^2 \frac{v\sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^p} + \varepsilon^3 (v_t + uv_x + vv_y).$$

Neglecting the  $\mathcal{O}(\varepsilon^2)$  terms, we get

$$\frac{\delta}{J_0} (h + \phi)_y = \mathcal{O}(\varepsilon^2),$$

and there exists  $H = H(x, t)$  such that



$$H(x, t) = h(x, y, t) + \phi(x, y) + \mathcal{O}(\varepsilon^2).$$

$\rightsquigarrow$  the free surface is almost flat in the  $y$ -direction, up to  $\mathcal{O}(\varepsilon^2)$

# Longitudinal velocity expansion

Highlighting the **dominant terms**, the second equation (times  $\varepsilon$ ) reads:

$$u_t + uu_x + vu_y + \frac{1}{\varepsilon} \frac{\delta}{J_0} (h + \phi)_x = \frac{1}{\varepsilon} \left( -\frac{u\sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^p} - \frac{I_0}{J_0} b_x \right).$$

To perform the asymptotic expansion of  $u$  with respect to  $\varepsilon$ , we write

$$u(x, y, t) = u_{2D}^{(0)}(x, y, t) + \mathcal{O}(\varepsilon).$$

Since  $h + \phi = H + \mathcal{O}(\varepsilon^2)$ , straightforward computations yield:

$$u_{2D}^{(0)} = C \frac{\Lambda}{\sqrt{|\Lambda|}} (H - \phi)^{p/2},$$

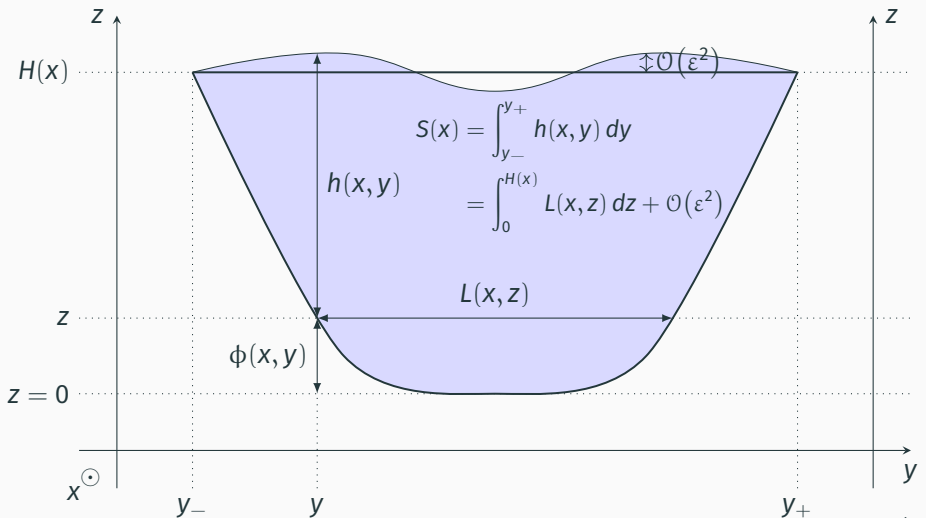
where we have defined the corrected slope  $\Lambda(x, t) = -\frac{I_0}{J_0} b_x - \frac{\delta}{J_0} H_x$ .

**Next step:** Build a **1D model** consistent with these expansions.

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# The river cross-section

To obtain a 1D model, we start by averaging the 2D equations: below, we display the cross-section of the river, with respect to  $x$ .



# Averaging the 2D system over the river width

1. The original mass conservation equation reads:

$$h_t + (hu)_x + (hv)_y = 0.$$

Therefore, since  $v(y_-) = v(y_+) = 0$ , we get:

$$\int_{y_-}^{y_+} h_t dy + \int_{y_-}^{y_+} (hu)_x dy = 0 \implies S_t + Q_x = 0,$$

where the averaged discharge  $Q$  is given by  $Q = \int_{y_-}^{y_+} hu dy$ .

2. Arguing the mass conservation and integrating the second equation between  $y_-$  and  $y_+$  yields:

$$Q_t + \left( \int_{y_-}^{y_+} hu^2 dy \right)_x = \frac{1}{\varepsilon} \int_{y_-}^{y_+} h \left( -\frac{l_0}{J_0} b_x - \frac{\delta}{J_0} (h + \phi)_x \right) dy - \frac{1}{\varepsilon} \int_{y_-}^{y_+} \frac{u \sqrt{u^2 + \varepsilon^2 v^2}}{C^2 h^{p-1}} dy.$$



## Averaging the 2D system

Finally, the averaged system reads as follows, up to  $\mathcal{O}(\varepsilon^2)$ :

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left( \int_{y_-}^{y_+} hu^2 dy \right)_x = \frac{1}{\varepsilon} \left( \Lambda S - \int_{y_-}^{y_+} \frac{u|u|}{C^2 h^{p-1}} dy \right) + \mathcal{O}(\varepsilon). \end{cases}$$

**Next step:** From the averaged system, build a truly 1D model that is **zeroth-order** accurate (up to  $\mathcal{O}(\varepsilon)$ ).

That is to say, the new model needs to ensure  $Q = Q_{2D}^{(0)} + \mathcal{O}(\varepsilon)$ , where

$$\begin{aligned} Q_{2D}^{(0)} &= \int_{y_-}^{y_+} hu_{2D}^{(0)} dy \\ &= \sqrt{|\Lambda|} \operatorname{sgn}(\Lambda) \int_{y_-}^{y_+} C (H - \phi)^{1+p/2} dy. \end{aligned}$$

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## Setting up the model

The integrated discharge equation, highlighting the **dominant terms** and multiplying by  $\varepsilon$ , is

$$\Lambda S - \int_{y_-}^{y_+} \frac{u|u|}{C^2 h^{p-1}} dy = \varepsilon \left( Q_t + \left( \int_{y_-}^{y_+} hu^2 dy \right)_x \right) + \mathcal{O}(\varepsilon^2).$$

At the **zeroth order**, i.e. **up to**  $\mathcal{O}(\varepsilon)$ , the right-hand side of this equation is neglected, and we get:

$$\Lambda S - \int_{y_-}^{y_+} \frac{u|u|}{C^2 h^{p-1}} dy = \mathcal{O}(\varepsilon).$$

We cannot directly use this equation in a 1D model, since it contains the unknown  $u$ , which depends on  $y$ .

Instead, we approximate the integral, up to  $\mathcal{O}(\varepsilon)$ , with a **new 1D friction term**.

# The friction model

First, we choose this 1D friction term as a **usual hydraulic engineering model**. Thus, we impose the following formula:

$$\frac{Q|Q|}{C_{1D}^2 S} = \int_{y_-}^{y_+} \frac{u|u|}{C^2 h^{p-1}} dy + \mathcal{O}(\varepsilon).$$

It contains a 1D friction coefficient<sup>5</sup>  $C_{1D}$ , to be determined.

According to the discharge equation, we get, up to  $\mathcal{O}(\varepsilon)$ :

$$\frac{Q|Q|}{C_{1D}^2 S} = \Lambda S + \mathcal{O}(\varepsilon) \quad \Longrightarrow \quad C_{1D}^2 = \frac{Q|Q|}{\Lambda S^2} + \mathcal{O}(\varepsilon).$$

Second, we impose  $Q = Q_{2D}^{(0)} + \mathcal{O}(\varepsilon)$ , to get the following expression of the friction coefficient:

$$C_{1D}^2 = \frac{Q_{2D}^{(0)} |Q_{2D}^{(0)}|}{\Lambda S^2}.$$

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<sup>5</sup>The coefficient  $C_{1D}^2$  usually contains the hydraulic radius, the Chézy coefficient, ...

# The final system

With the **new friction model**, the discharge equation reads

$$\Lambda S - \frac{Q|Q|}{C_{1D}^2 S} = \varepsilon \left( Q_t + \left( \int_{y_-}^{y_+} hu^2 dy \right)_x \right) + \mathcal{O}(\varepsilon).$$

We choose to approximate the **integral in the flux** to describe the advection of the discharge:

$$\varepsilon \int_{y_-}^{y_+} hu^2 dy = \varepsilon \frac{\left( \int_{y_-}^{y_+} hu dy \right)^2}{\int_{y_-}^{y_+} h dy} + \mathcal{O}(\varepsilon) = \varepsilon \frac{Q^2}{S} + \mathcal{O}(\varepsilon).$$

The resulting discharge equation is

$$\varepsilon \left( Q_t + \left( \frac{Q^2}{S} \right)_x \right) = S \left( \Lambda - \underbrace{\frac{Q|Q|}{C_{1D}^2 S^2}}_{\mathcal{J}} \right) + \mathcal{O}(\varepsilon).$$

# The final system

Finally, the zeroth-order accurate 1D system reads:

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left(\frac{Q^2}{S}\right)_x = \frac{1}{\varepsilon} S(\Lambda - \mathcal{J}). \end{cases}$$

Let us double check that this model is sufficient to **recover the zeroth-order expansion of  $Q$** .

With  $Q = Q_{\text{model}}^{(0)} + \mathcal{O}(\varepsilon)$ , we get, **at the zeroth order**:

$$\begin{aligned} \Lambda = \mathcal{J} + \mathcal{O}(\varepsilon) &\implies \Lambda = \overbrace{\Lambda \frac{Q|Q|}{Q_{2D}^{(0)}|Q_{2D}^{(0)}|}}^{\mathcal{J}} + \mathcal{O}(\varepsilon) = \Lambda \frac{Q_{\text{model}}^{(0)}|Q_{\text{model}}^{(0)}|}{Q_{2D}^{(0)}|Q_{2D}^{(0)}|} + \mathcal{O}(\varepsilon) \\ &\implies Q_{\text{model}}^{(0)} = Q_{2D}^{(0)} + \mathcal{O}(\varepsilon). \end{aligned}$$

## The final system

Finally, the zeroth-order accurate 1D system reads:

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left(\frac{Q^2}{S}\right)_x = \frac{1}{\varepsilon} S \left( -\frac{I_0}{J_0} b_x - \frac{\delta}{J_0} H_x - \mathcal{J} \right). \end{cases}$$

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# The final system

Finally, the zeroth-order accurate 1D system reads:

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left(\frac{Q^2}{S}\right)_x + \frac{SH_x}{F^2} = \frac{1}{\varepsilon} S(\mathcal{J} - \mathcal{J}). \end{cases}$$

This form is quite similar to that of the the usual models. All the complexity lies within the **friction model**  $\mathcal{J}$  and in the expression of the **friction coefficient**  $C_{1D}$ .

↪ We have derived a **zeroth-order model** governed by a hyperbolic system of balance laws.

↪ We also enhance this approach to derive a **first-order model**, based on the energy equation.

**Next step:** Numerical validation of these models.

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# Numerical schemes

To handle the **stiff relaxation source term**, we introduce an **implicit splitting procedure**.

The zeroth-order model is made of a **non-stiff part** and a **stiff part**:

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left(\frac{Q^2}{S}\right)_x + \frac{1}{\varepsilon} \frac{\delta}{J_0} S H_x = \frac{1}{\varepsilon} S(\mathcal{J} - \mathcal{J}). \end{cases}$$

**First**, we consider the **non-stiff part**:

$$\begin{cases} S_t + Q_x = 0, \\ Q_t + \left(\frac{Q^2}{S}\right)_x = 0, \end{cases}$$

which we discretize using an **upwind finite difference** scheme.

# Numerical schemes

**Second**, we consider the **stiff part**:

$$\begin{cases} S_t = 0, \\ Q_t + \frac{1}{\varepsilon} \frac{\delta}{J_0} S H_x = \frac{1}{\varepsilon} S(\mathcal{J} - \mathcal{J}). \end{cases}$$

Since  $S_t = 0$ , we are left with the following ODE on  $Q$ :

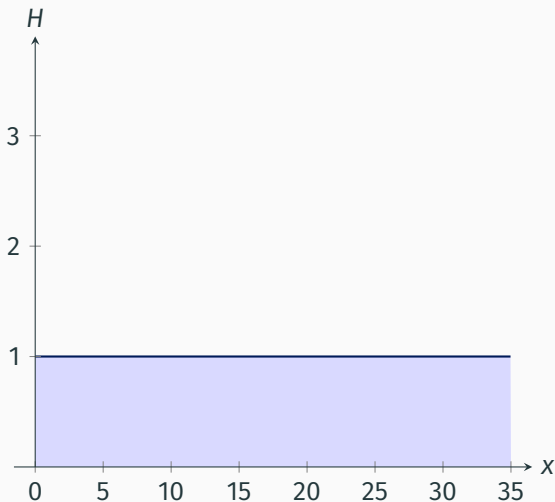
$$Q_t = \frac{1}{\varepsilon} S \Lambda \left( 1 - \frac{Q^2}{(Q_{2D}^{(0)})^2} \right),$$

which we can solve exactly, to get

$$Q(t) = Q_{2D}^{(0)} \frac{\tanh\left(\frac{1}{\varepsilon} \frac{S|\Lambda|}{|Q_{2D}^{(0)}|} t\right) + \frac{Q(0)}{Q_{2D}^{(0)}}}{1 + \tanh\left(\frac{1}{\varepsilon} \frac{S|\Lambda|}{|Q_{2D}^{(0)}|} t\right) \frac{Q(0)}{Q_{2D}^{(0)}}} \xrightarrow{\varepsilon \rightarrow 0} Q_{2D}^{(0)}.$$

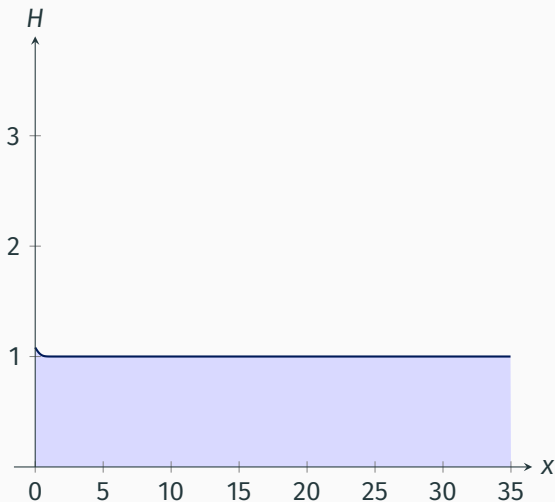
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We consider a 5-year flood for the Garonne river upstream of Toulouse; we take  $F = 0.09$  and  $\varepsilon \simeq 0.175$ .



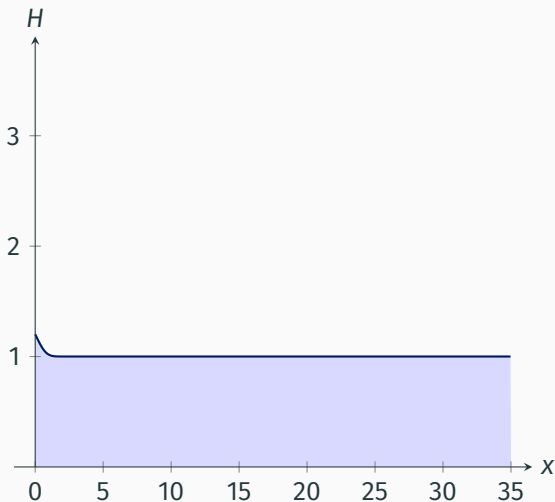
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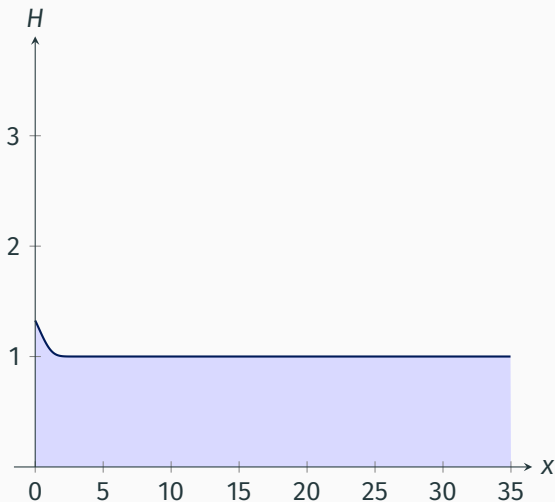
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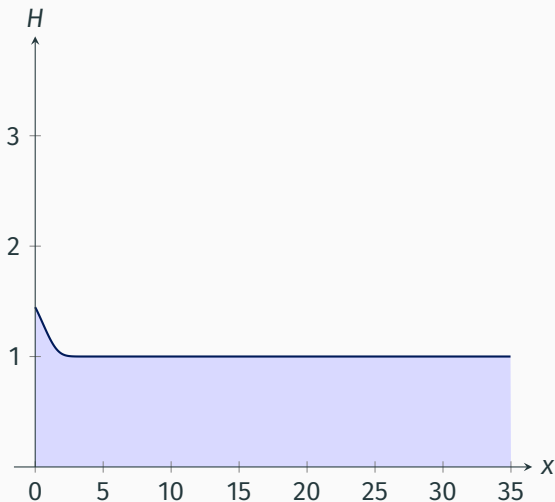
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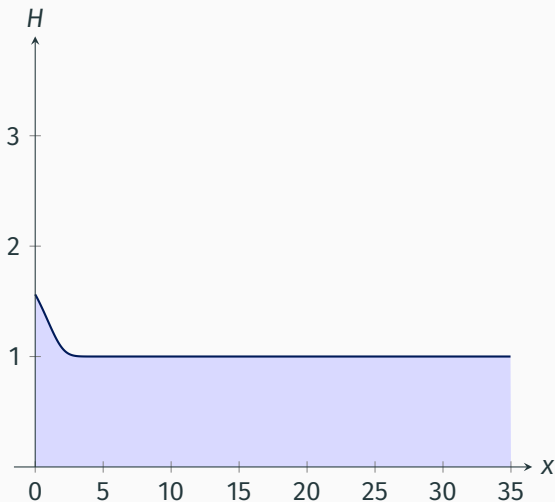
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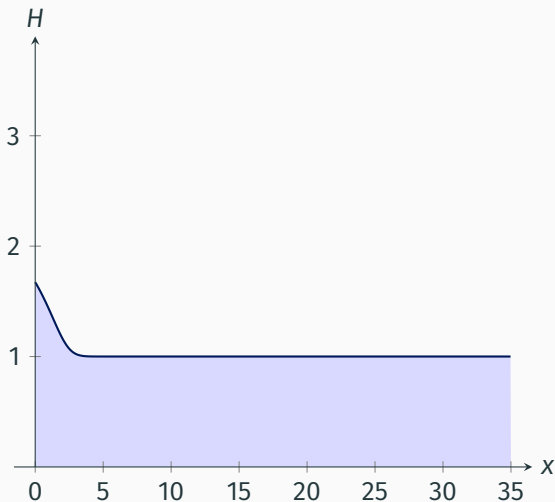
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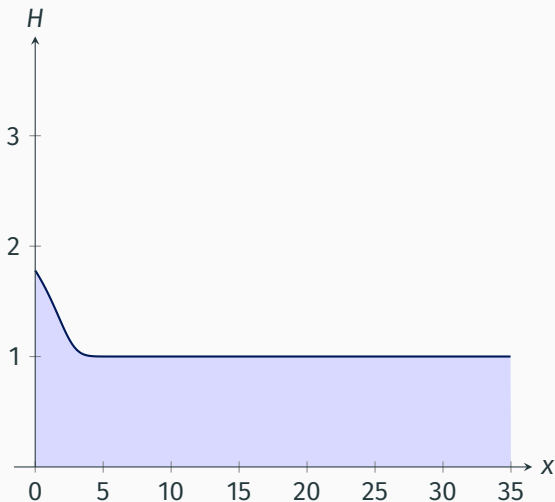
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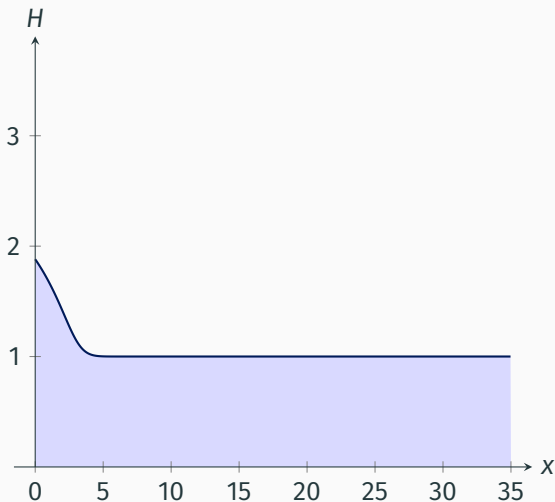
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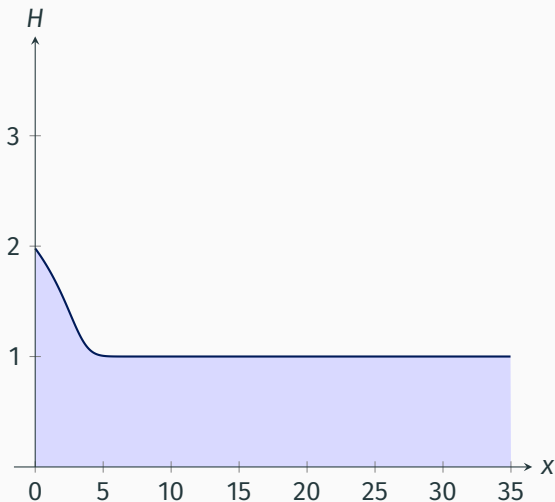
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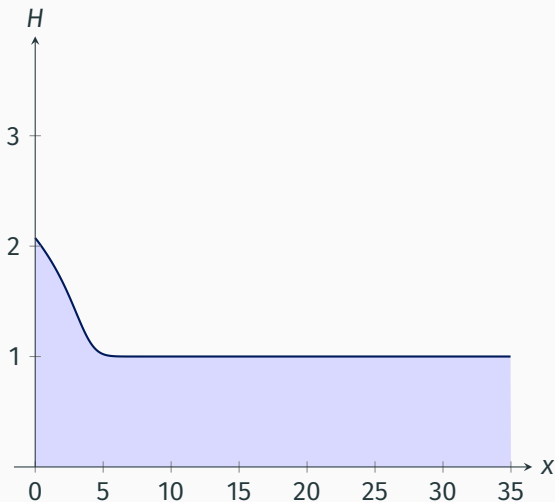
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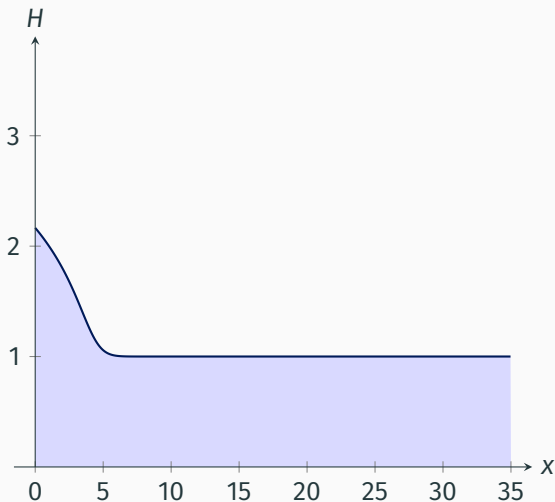
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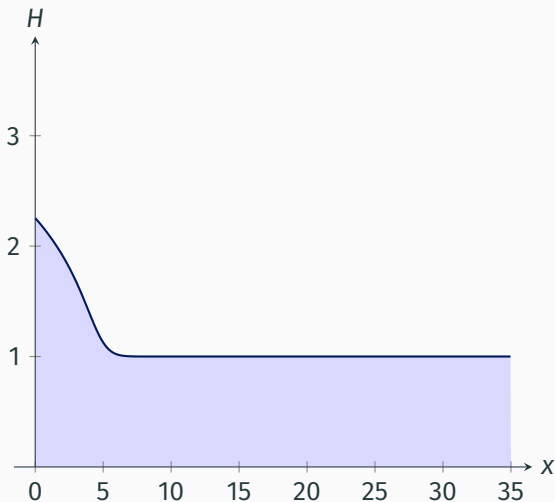
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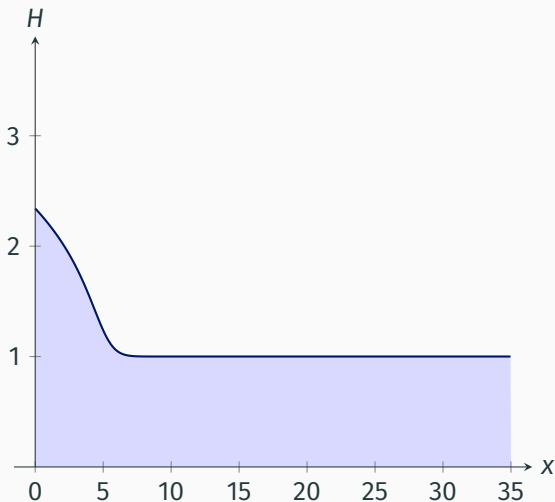
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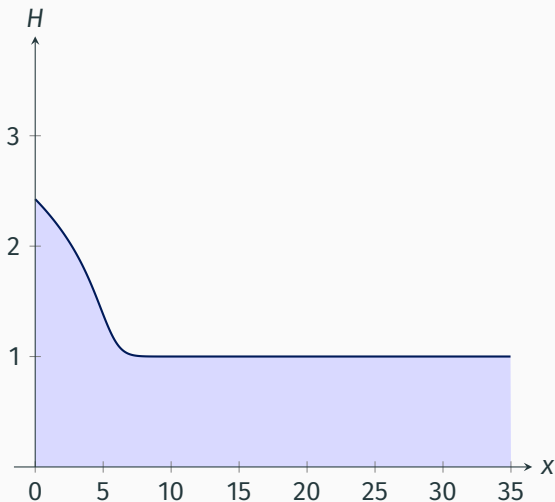
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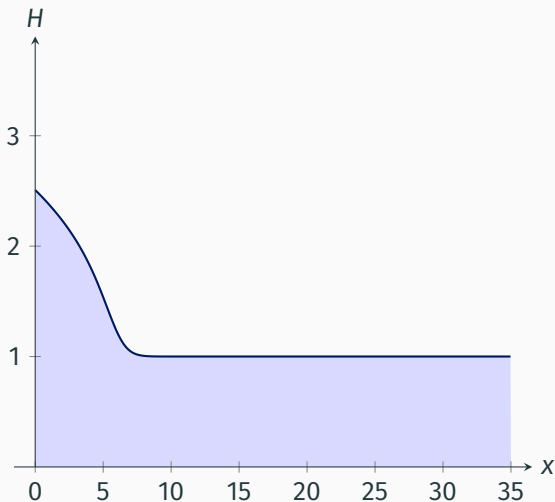
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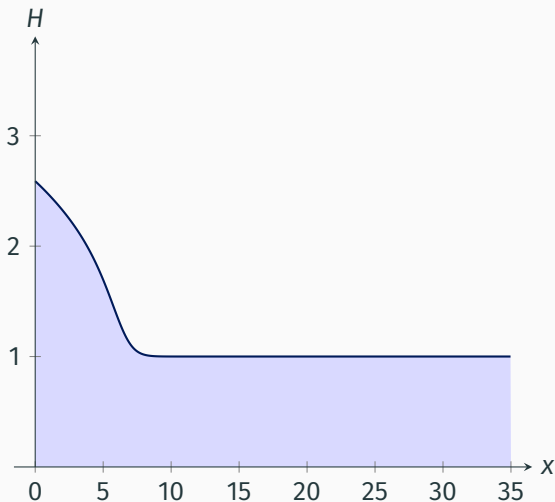
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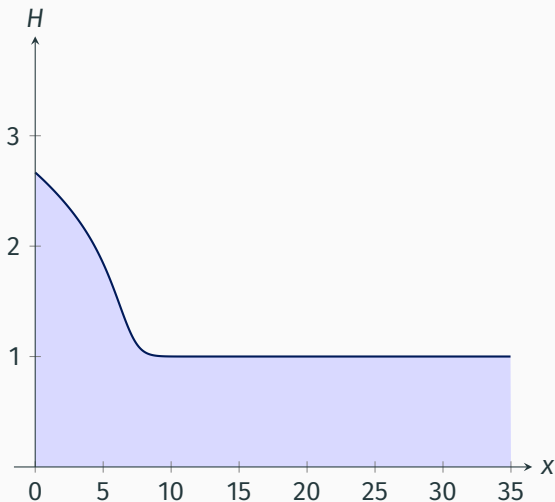
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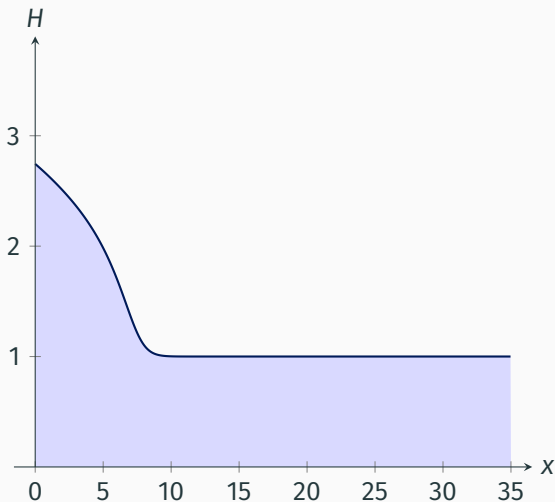
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We consider a 5-year flood for the Garonne river upstream of Toulouse; we take  $F = 0.09$  and  $\varepsilon \simeq 0.175$ .



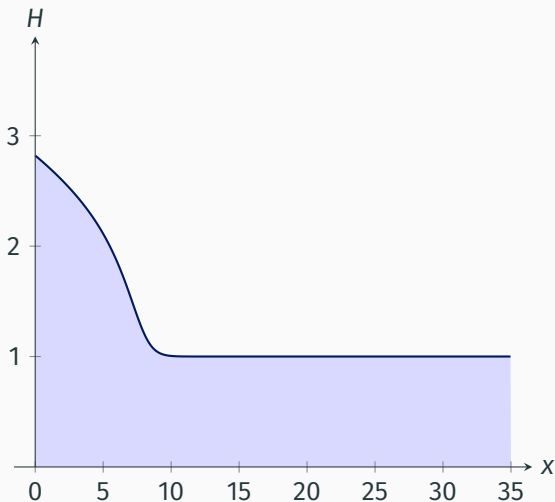
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## Unsteady flood flow

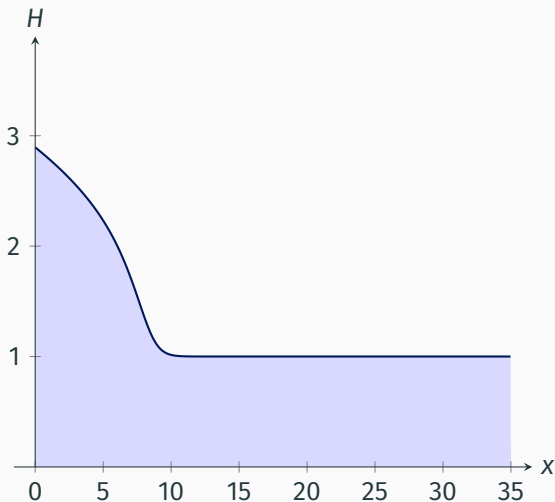
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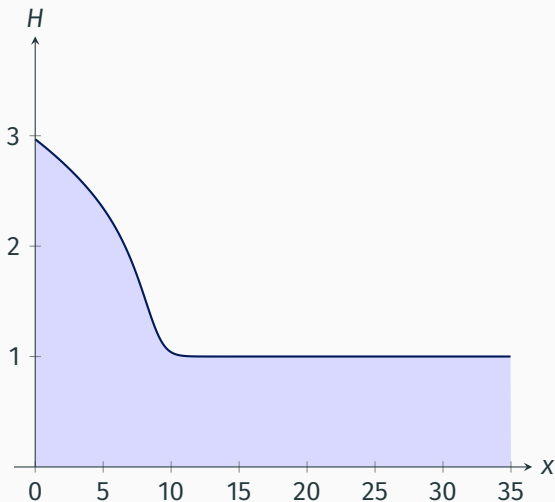
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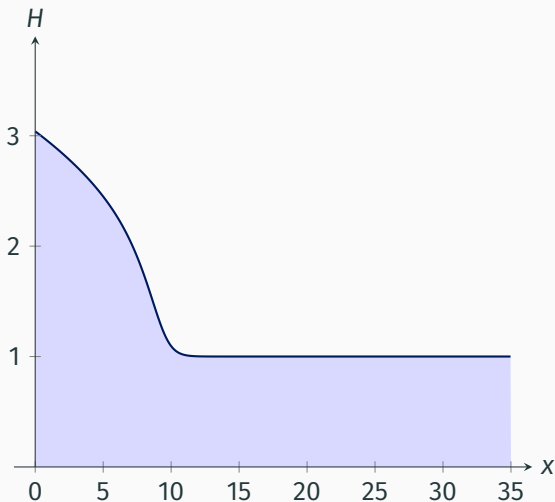
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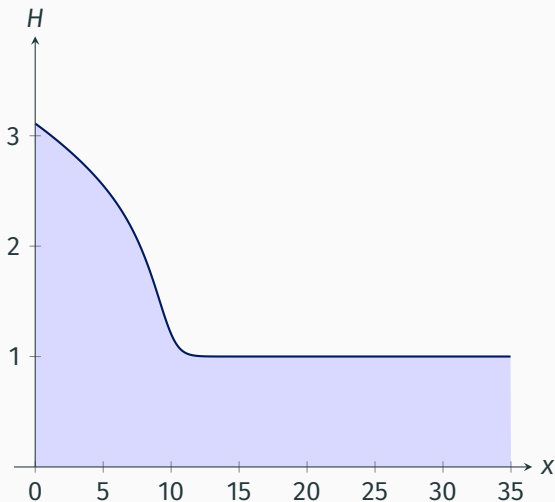
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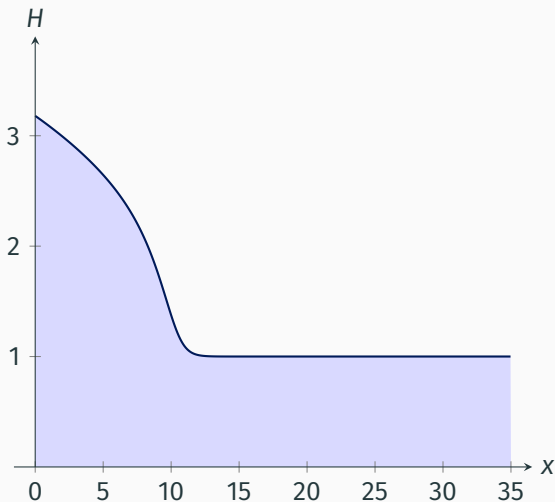
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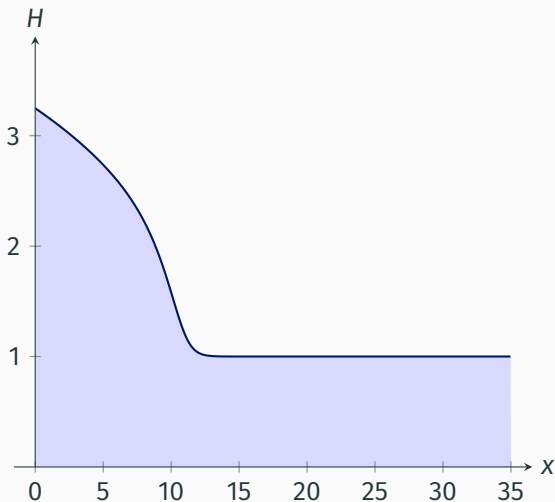
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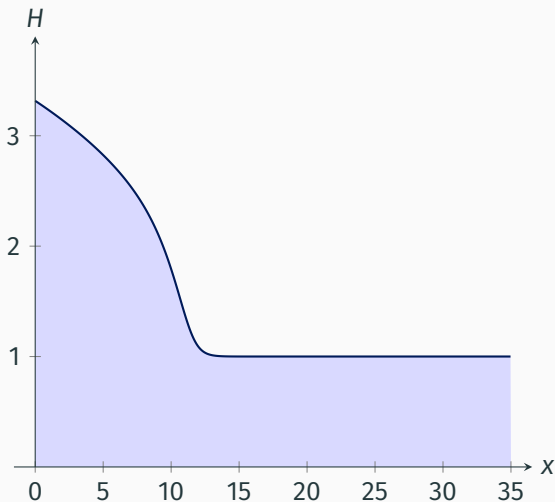
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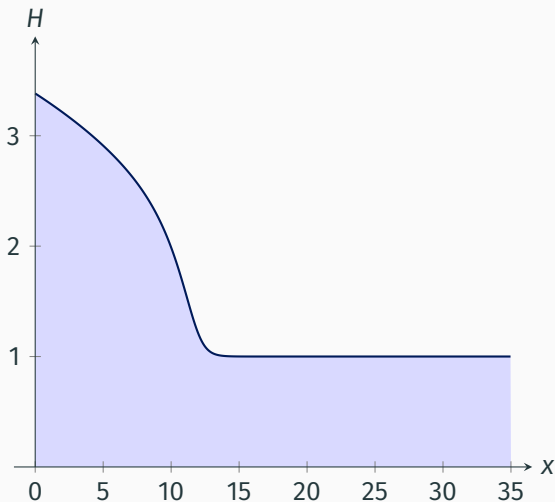
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## Unsteady flood flow

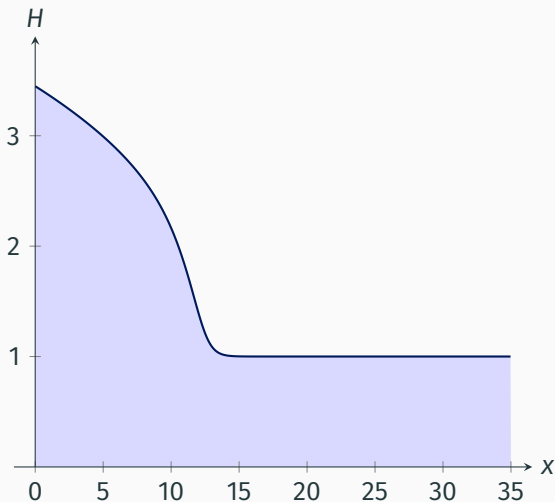
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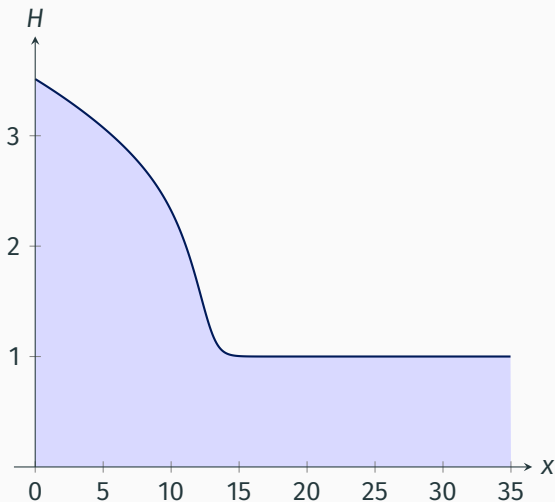
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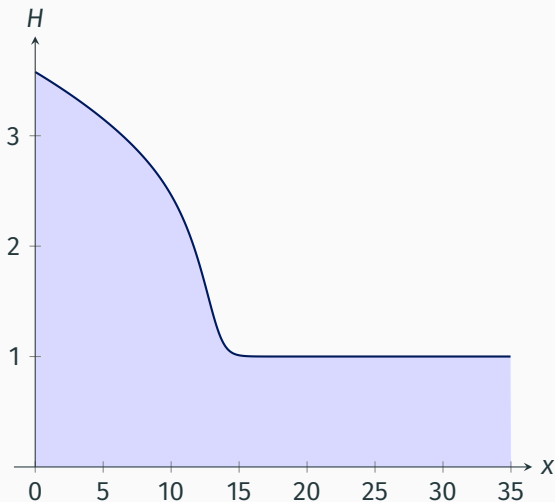
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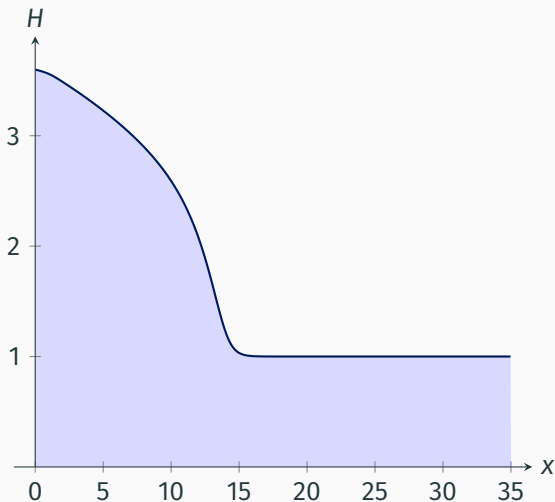
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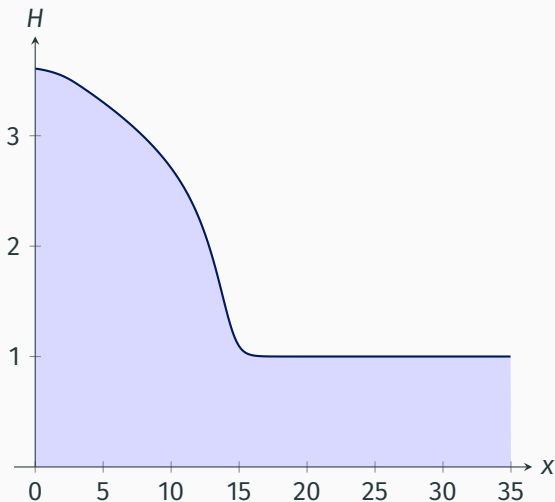
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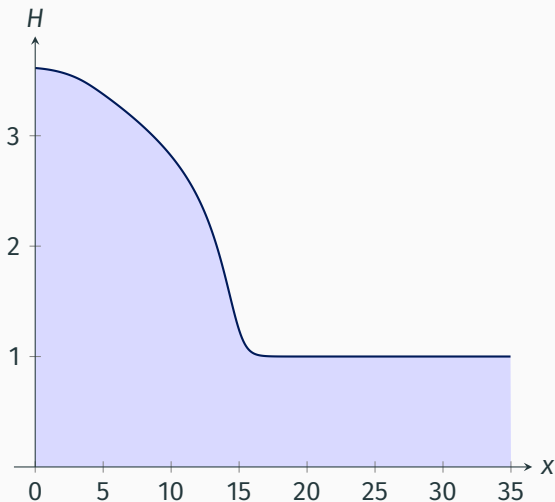
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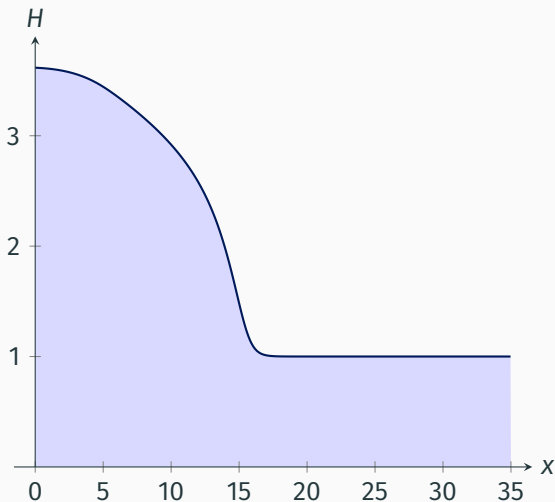
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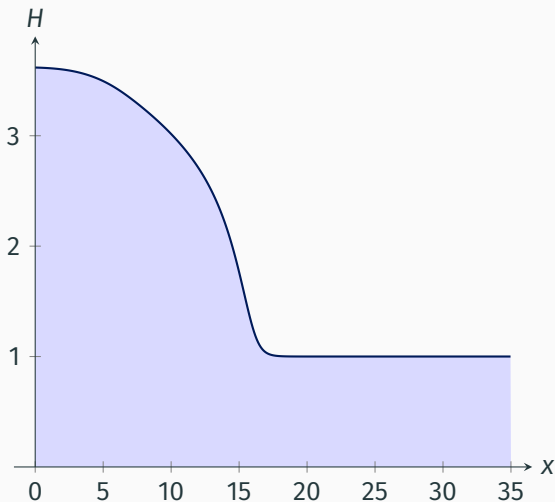
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## Unsteady flood flow

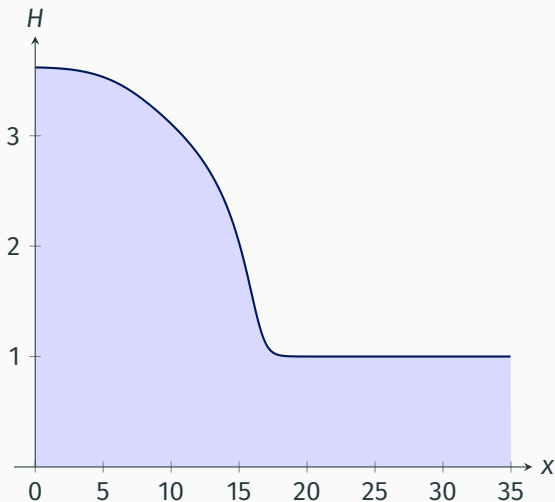
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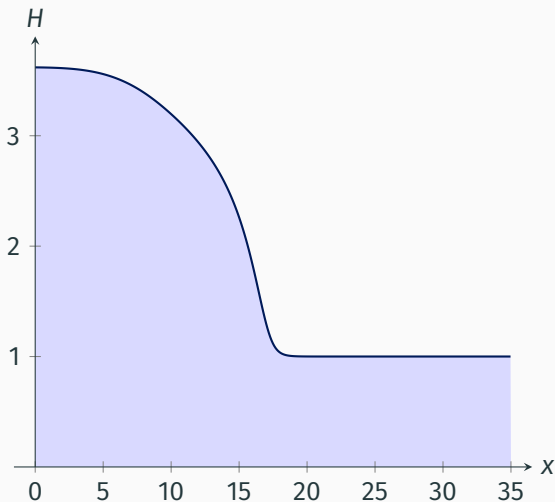
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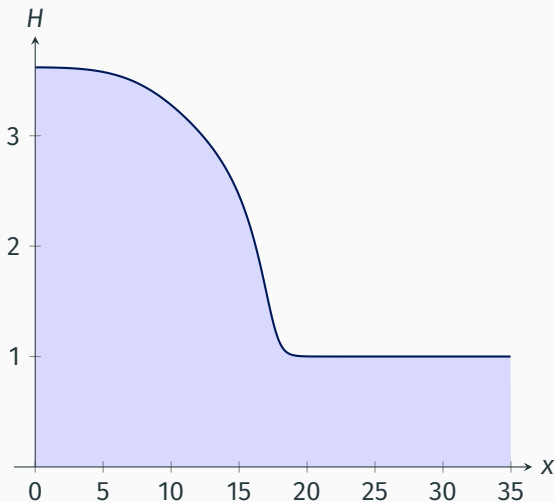
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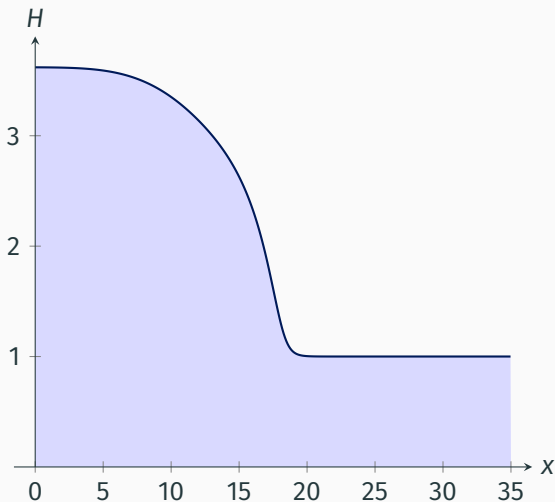
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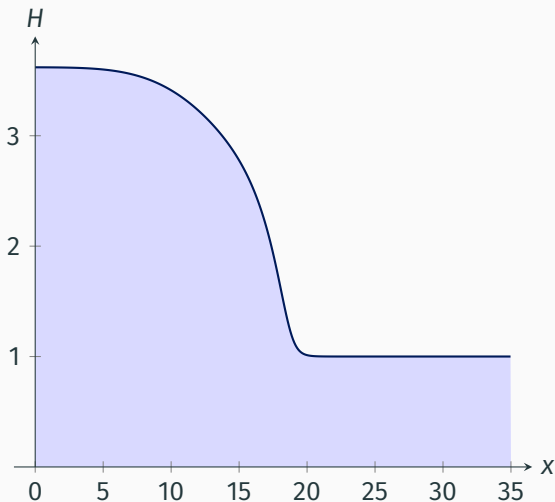
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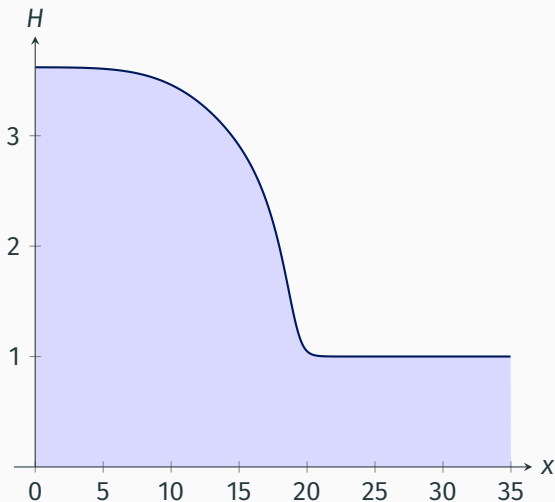
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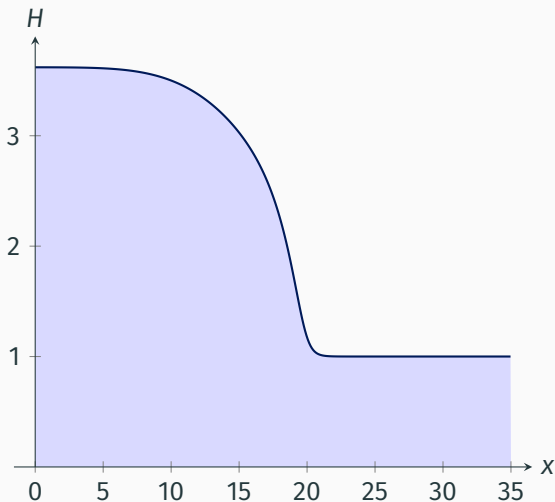
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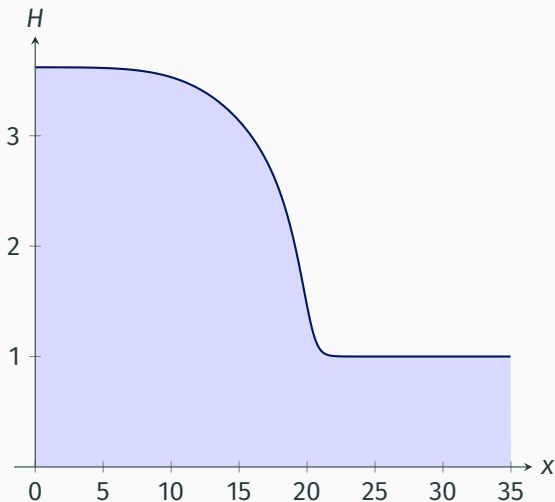
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## Unsteady flood flow

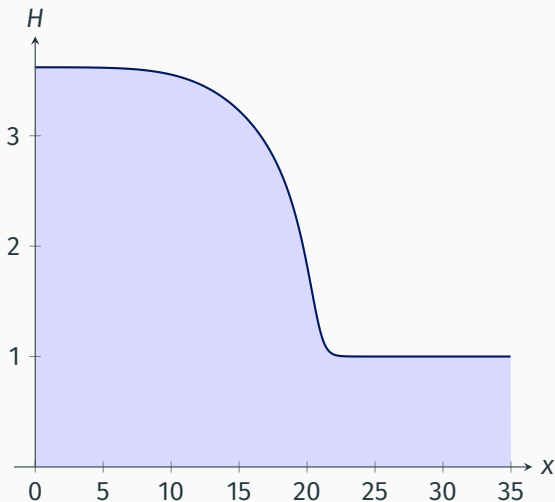
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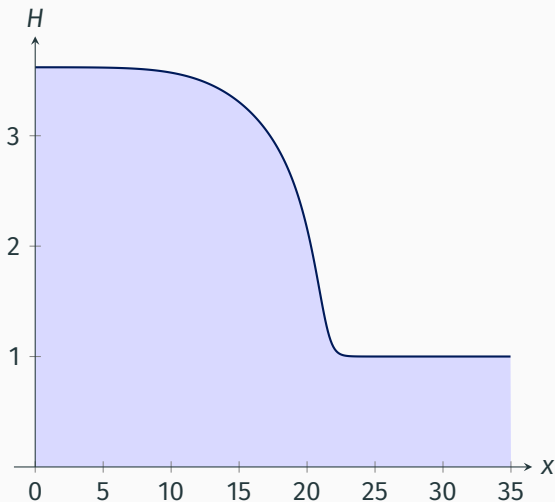
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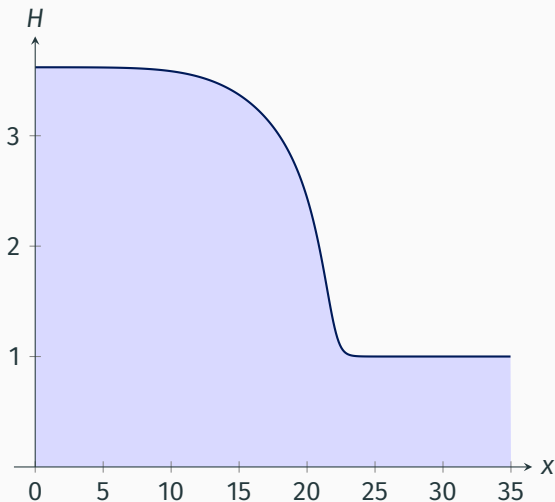
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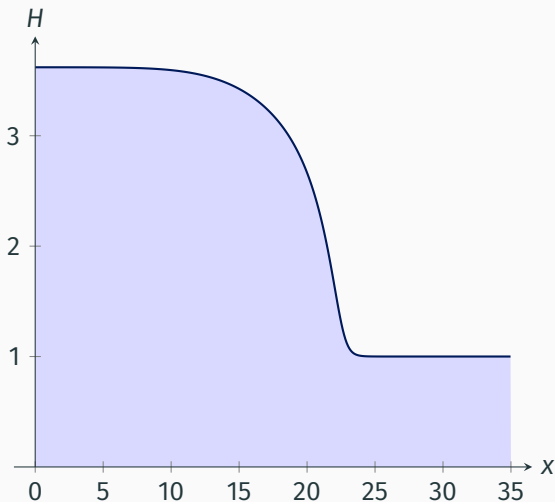
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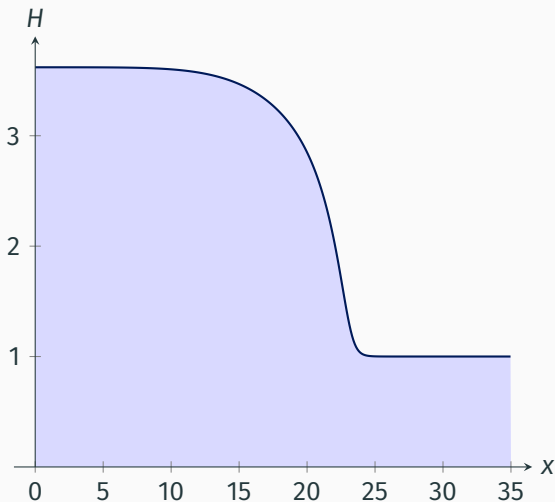
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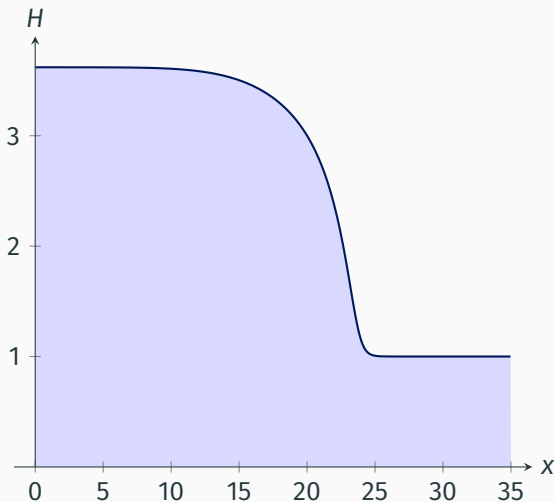
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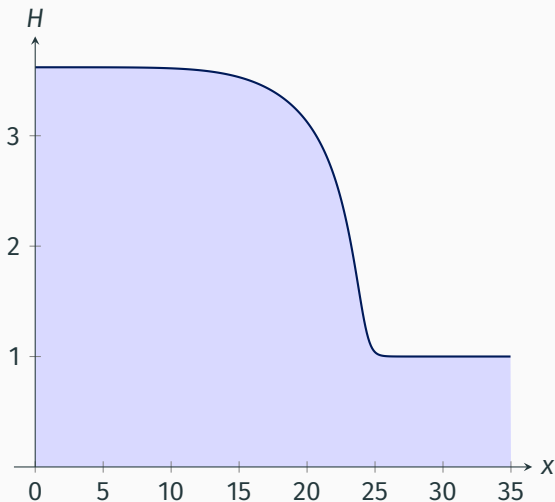
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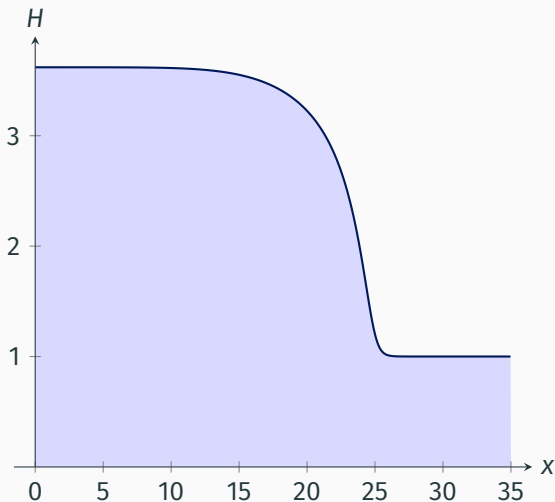
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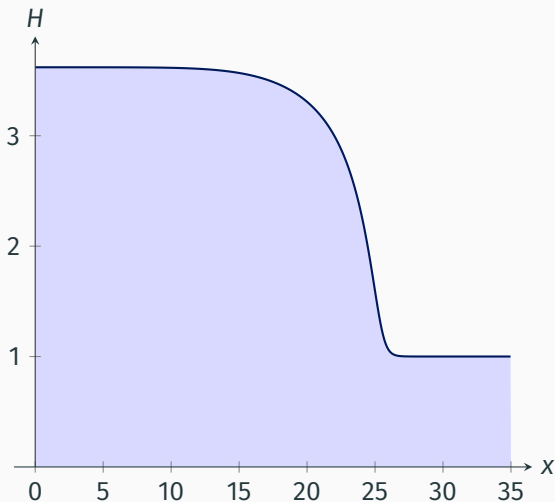
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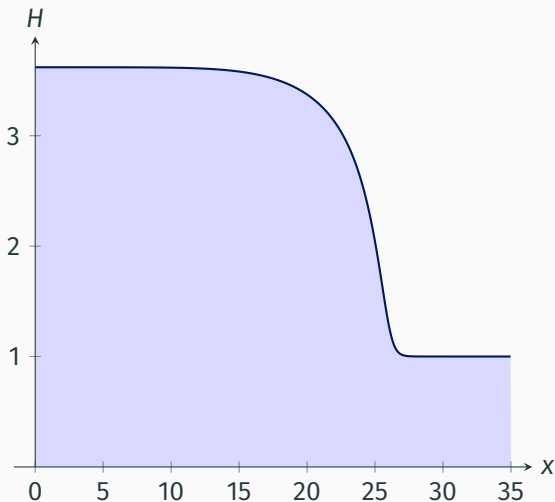
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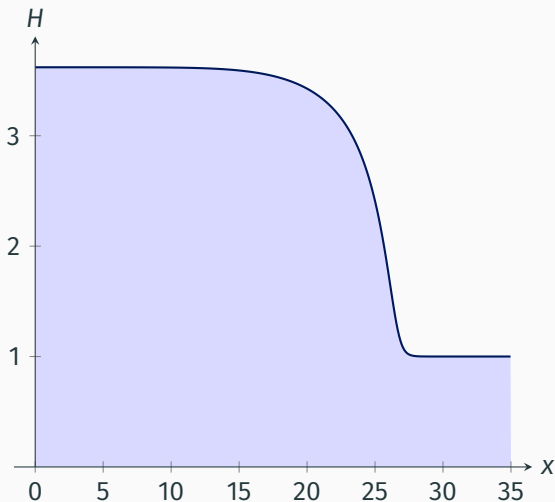
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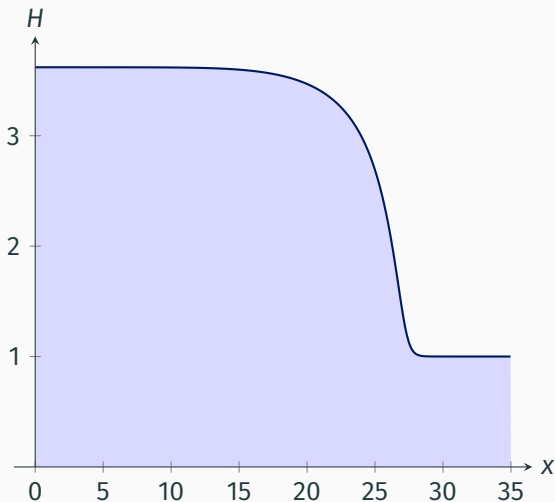
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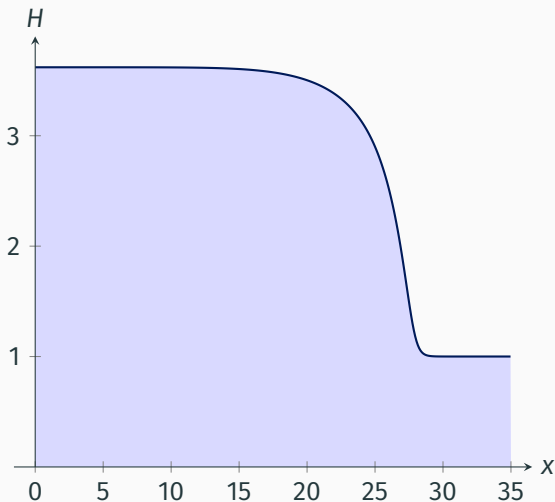
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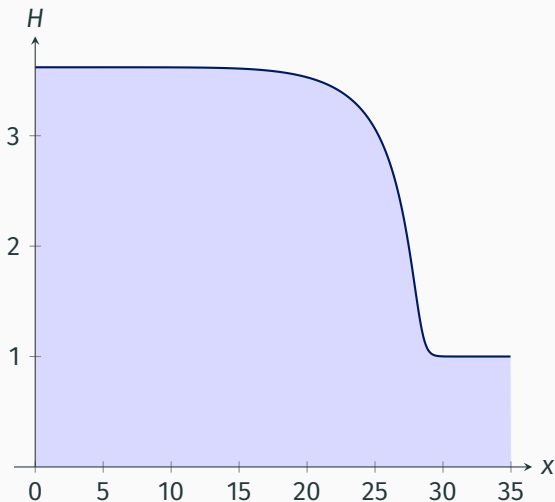
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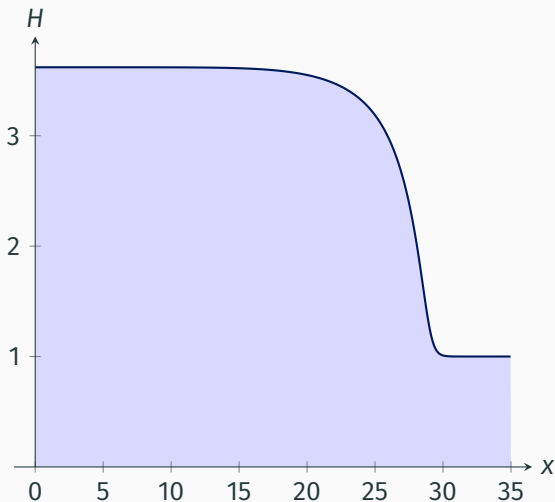
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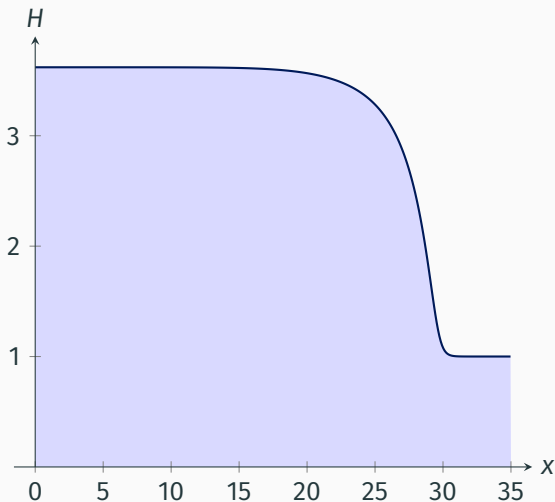
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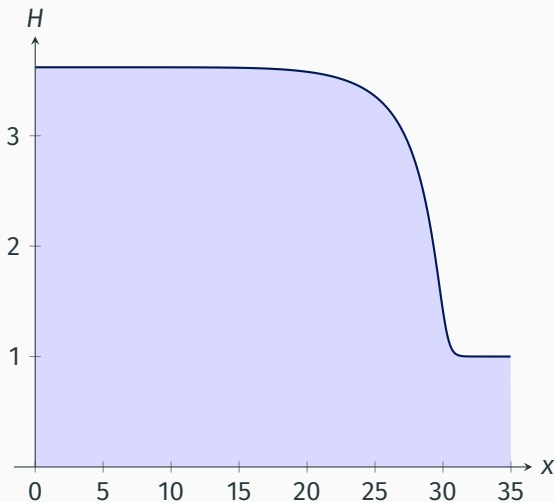
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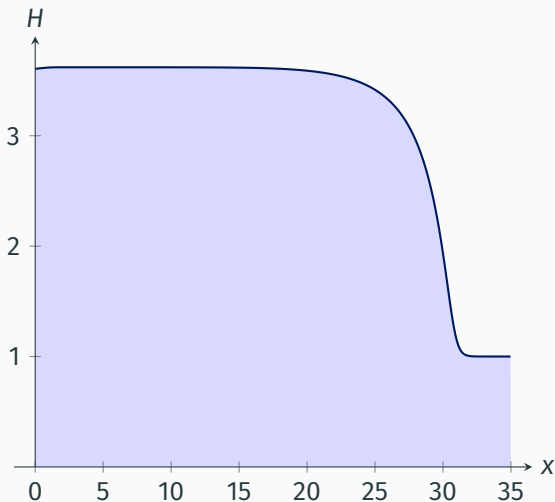
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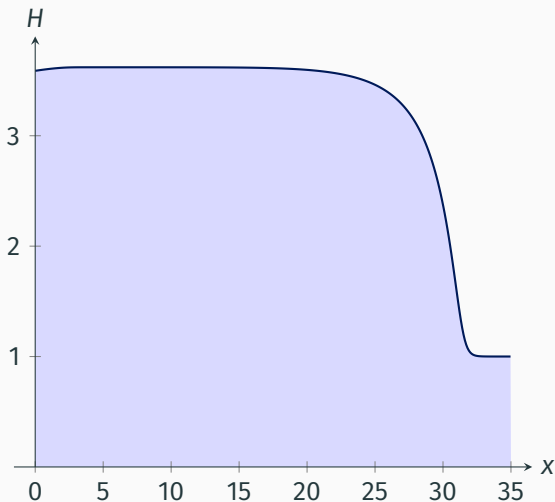
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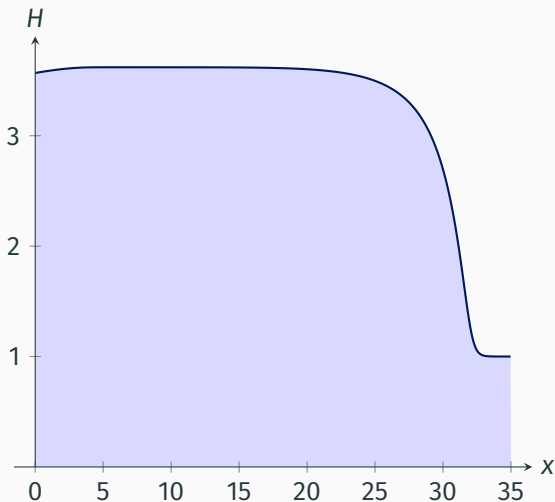
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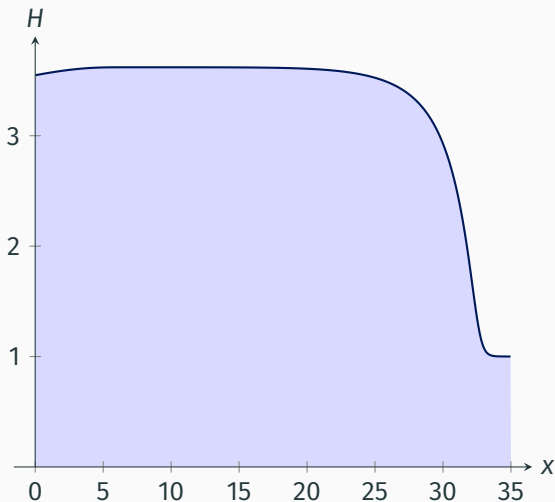
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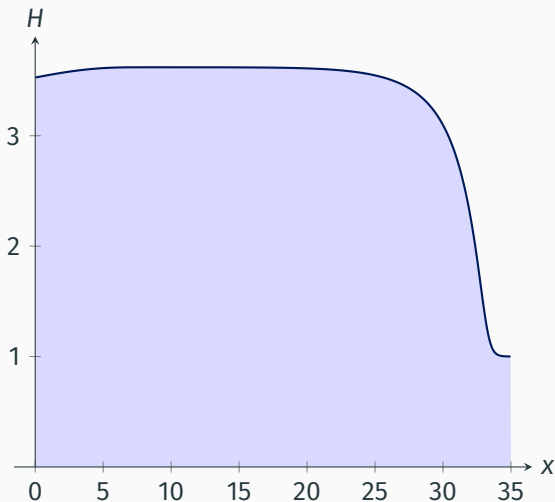
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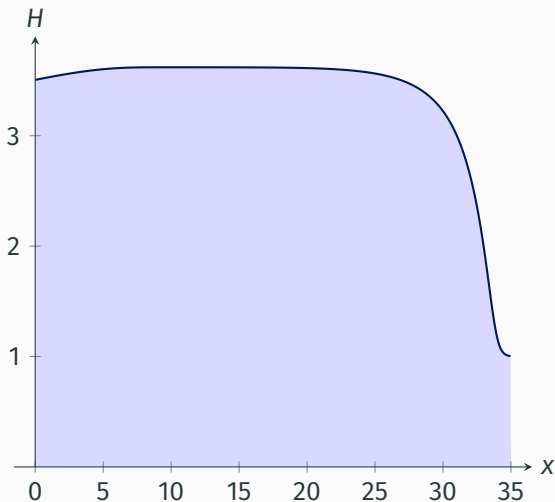
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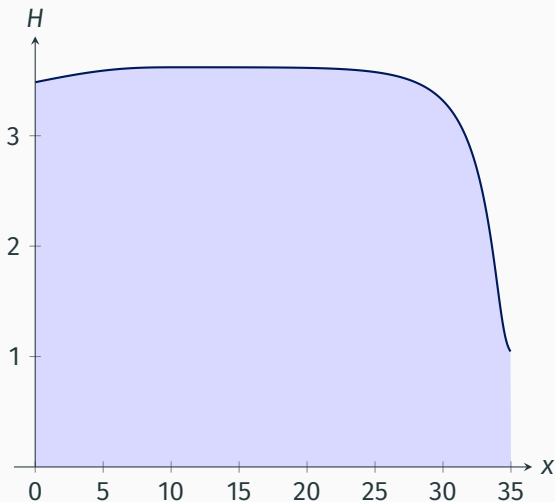
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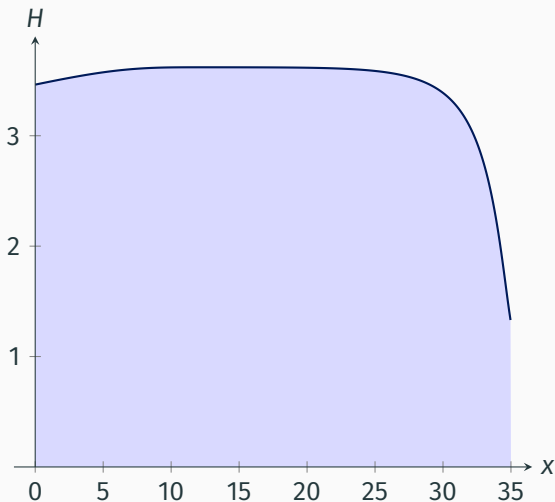
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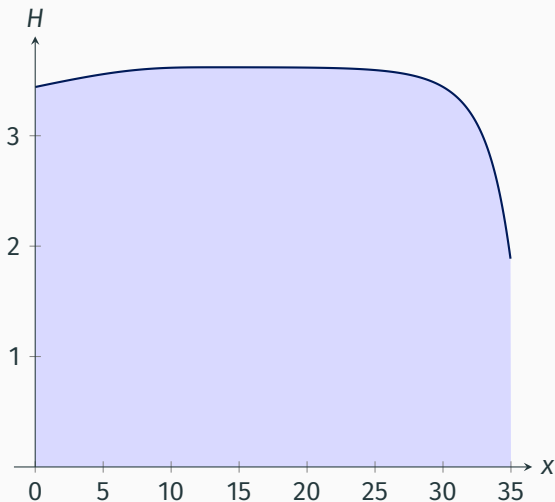
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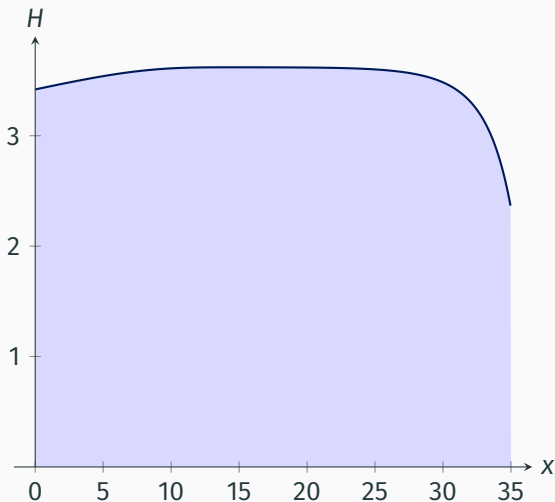
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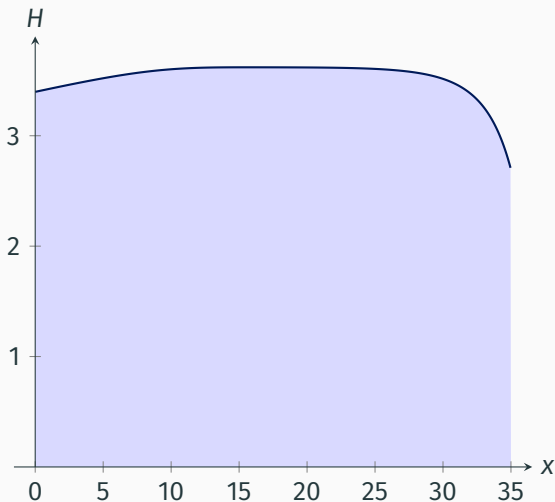
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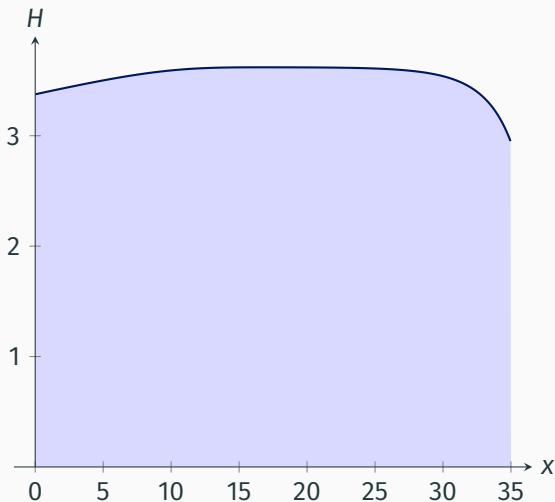
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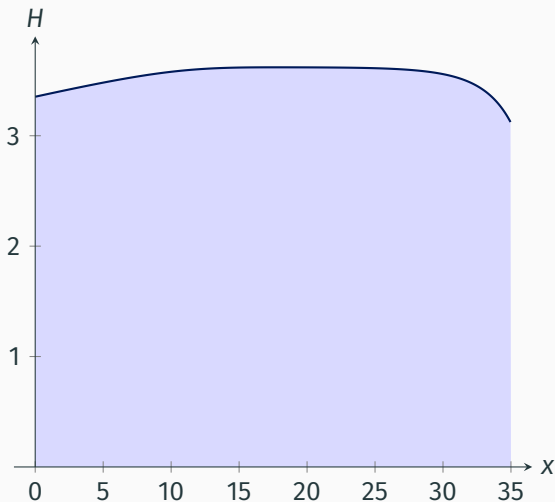
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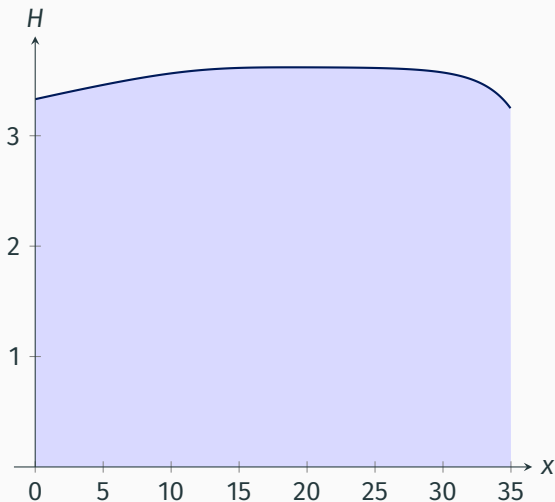
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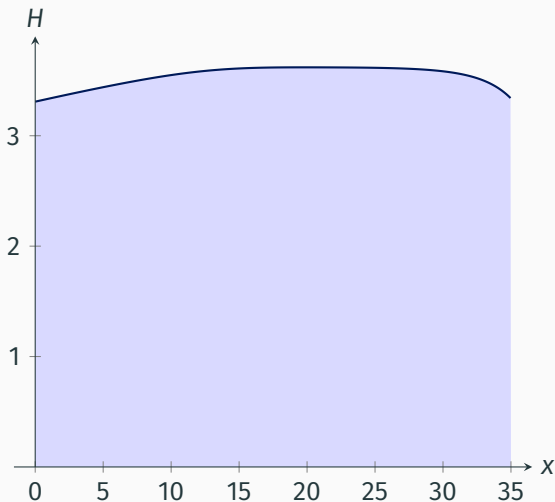
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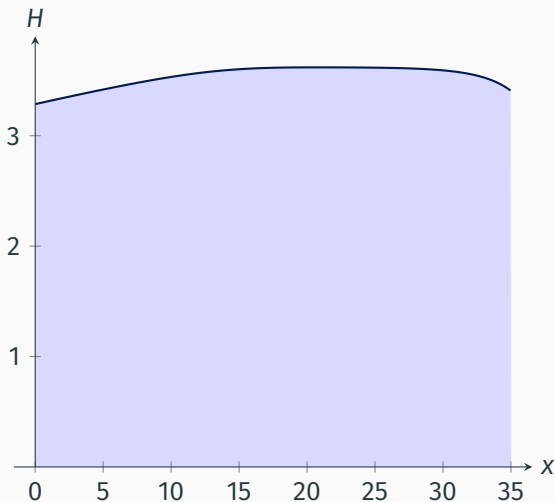
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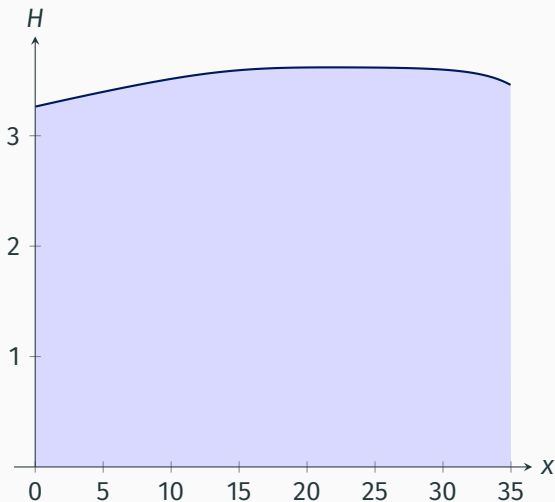
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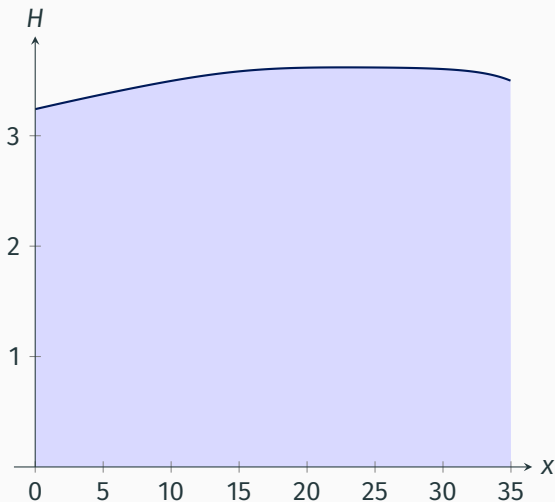
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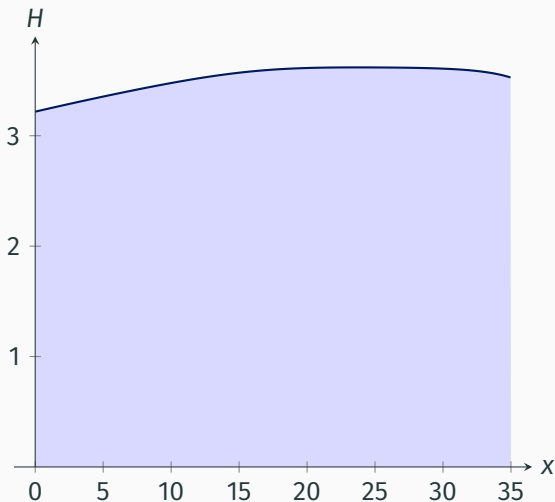
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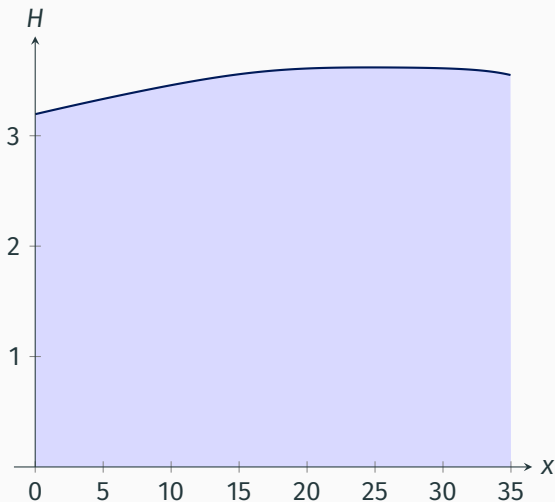
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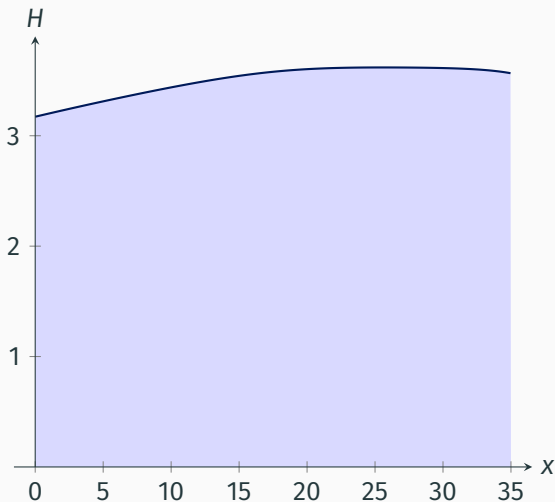
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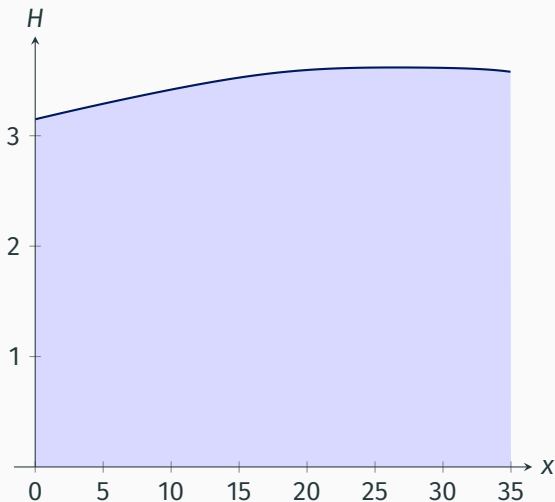
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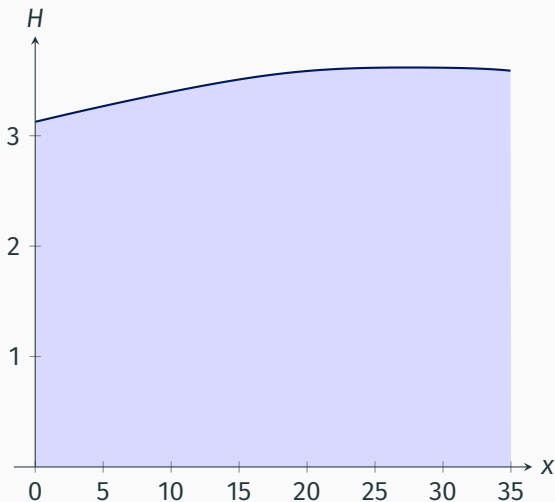
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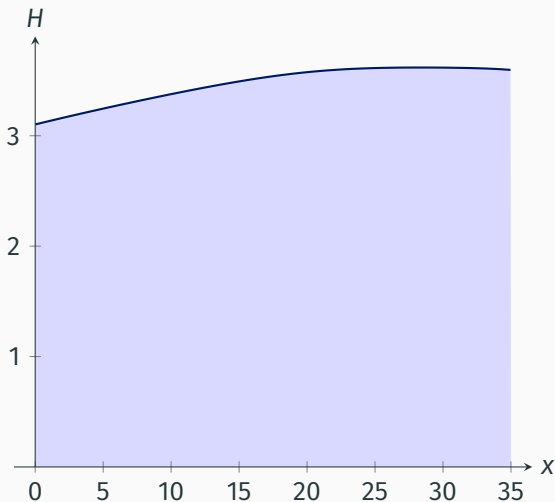
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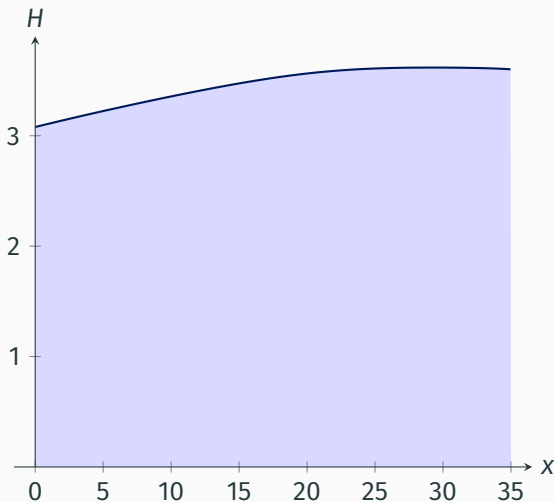
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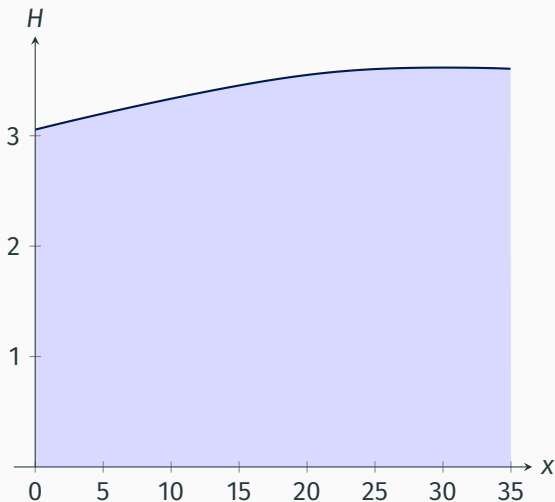
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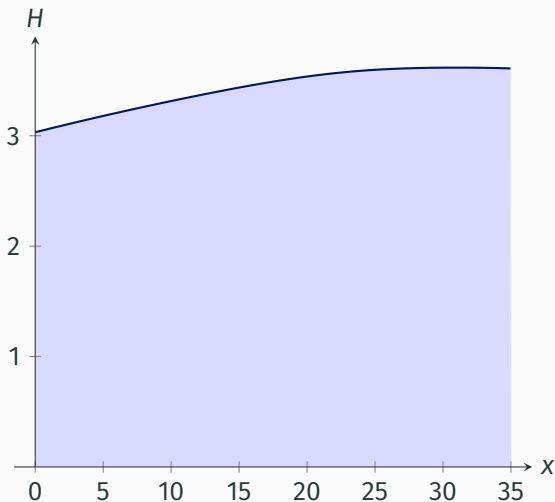
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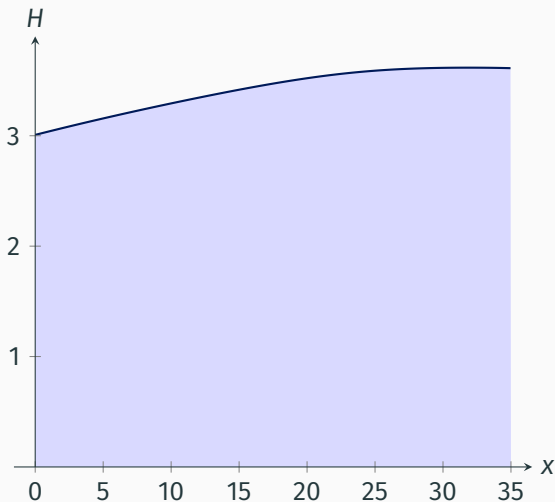
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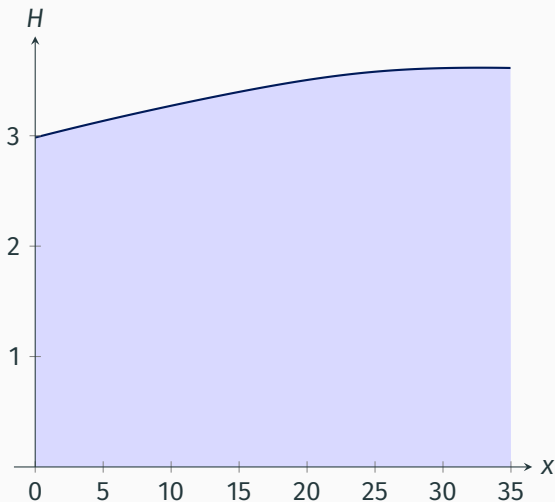
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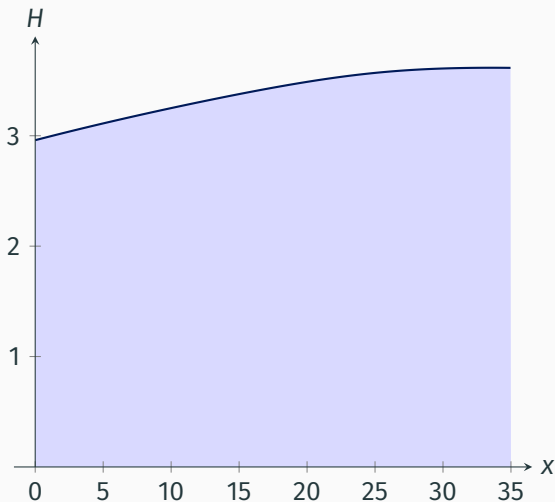
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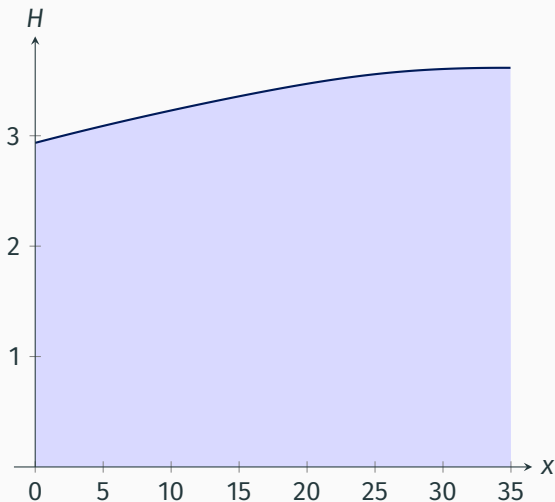
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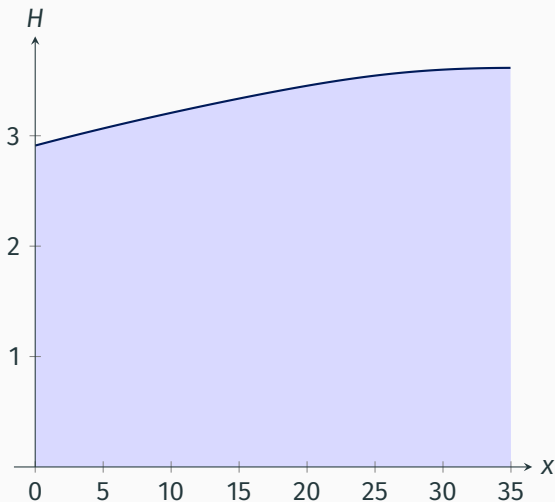
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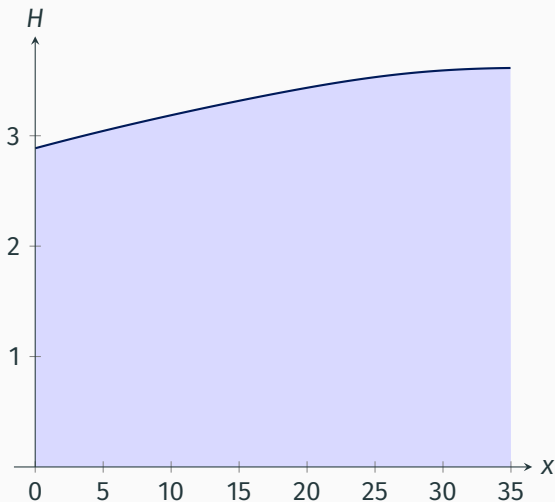
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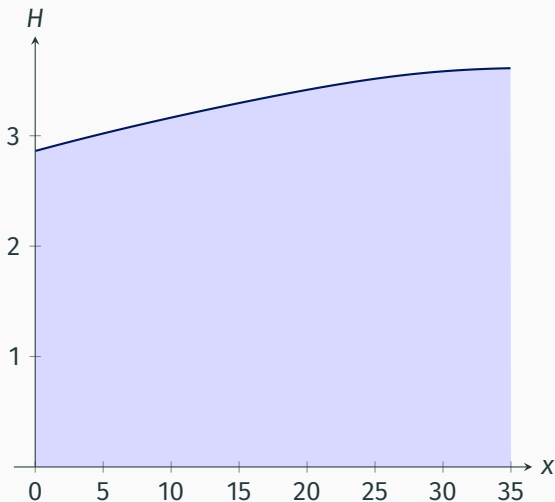
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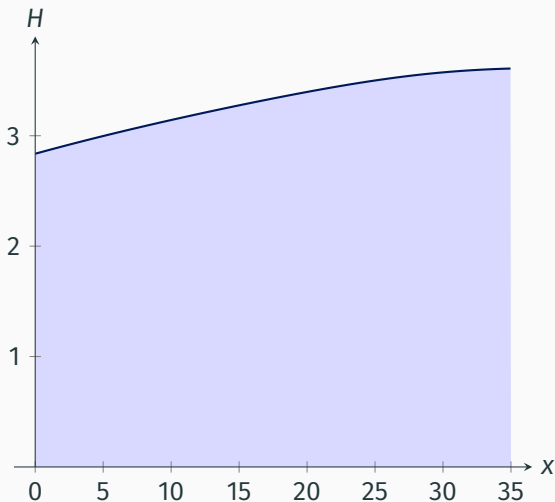
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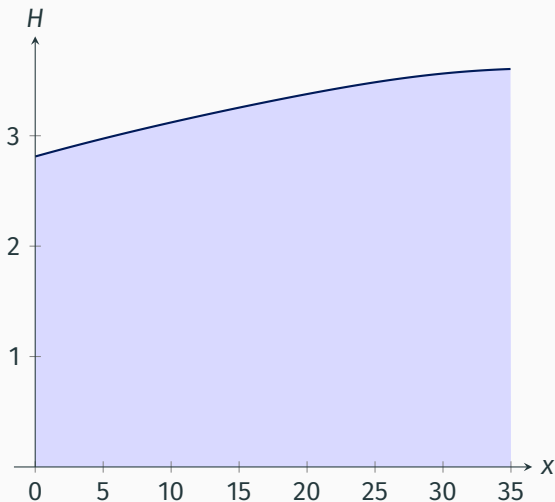
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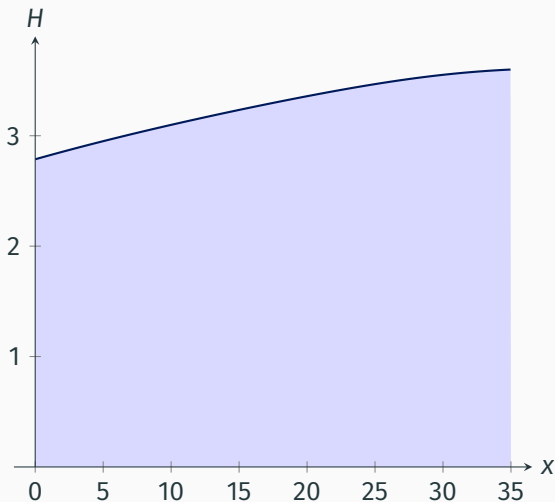
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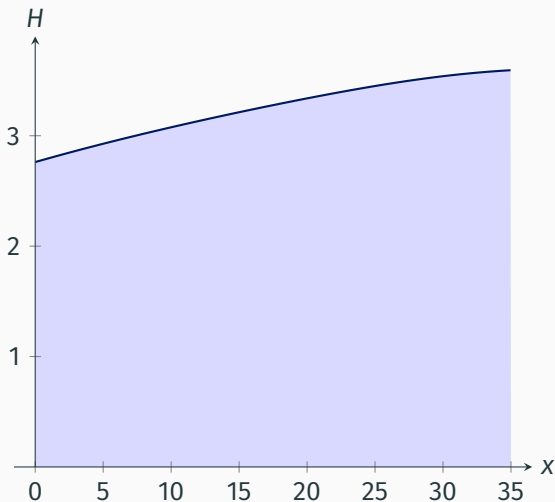
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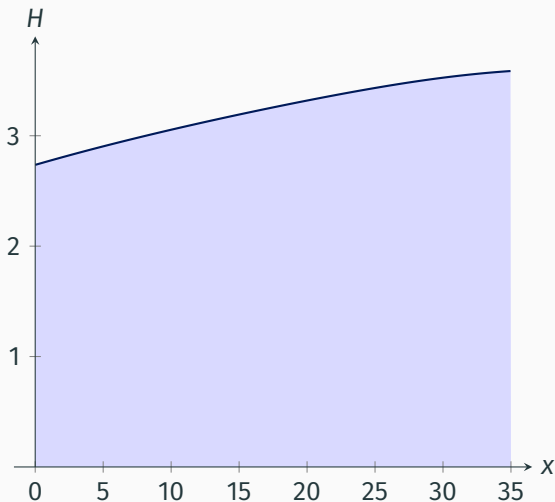
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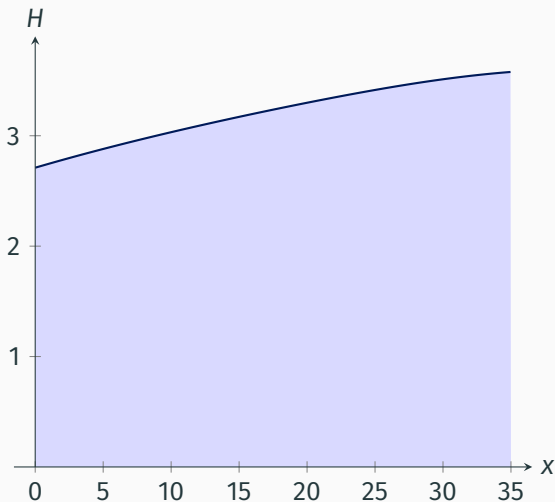
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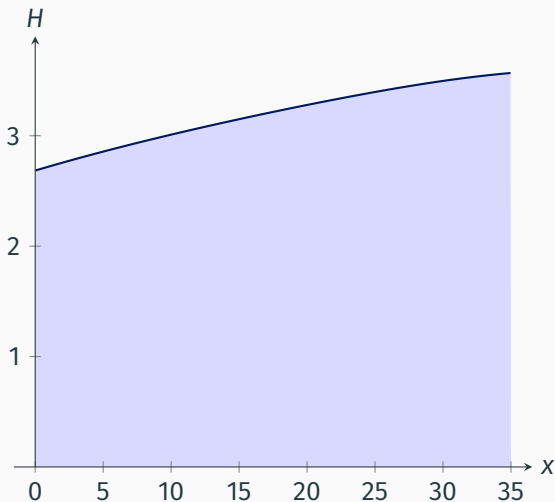
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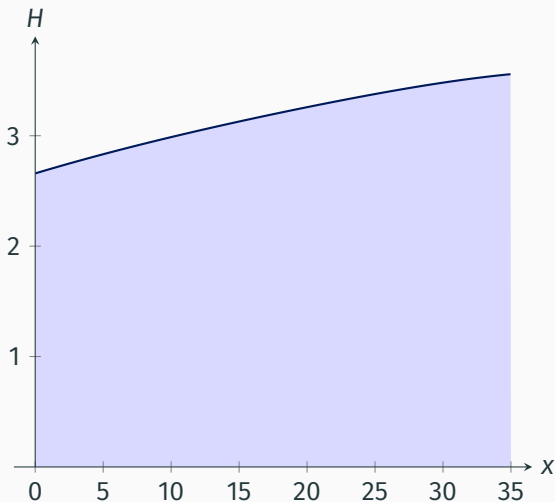
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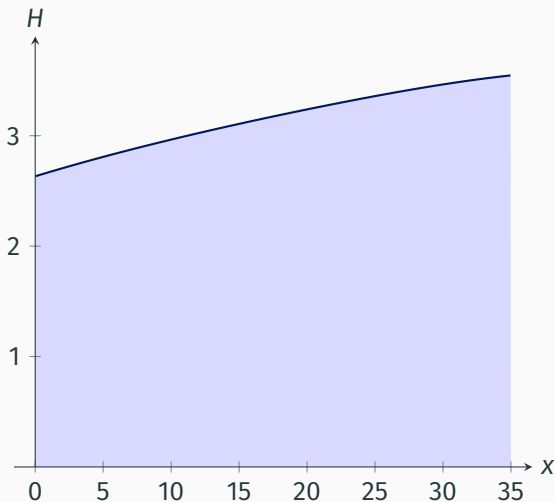
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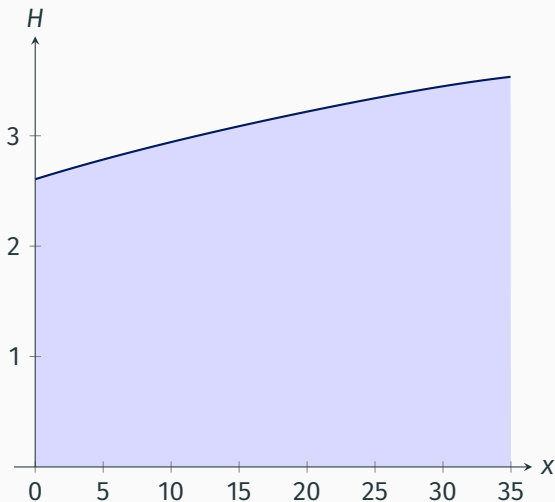
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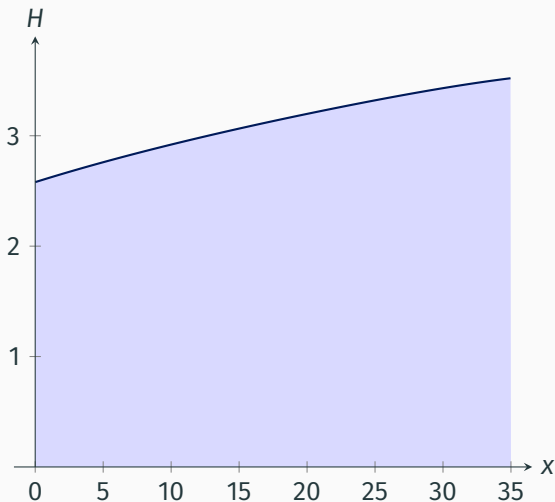
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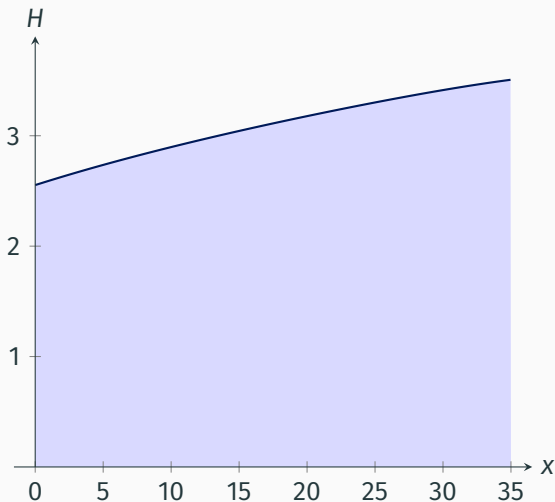
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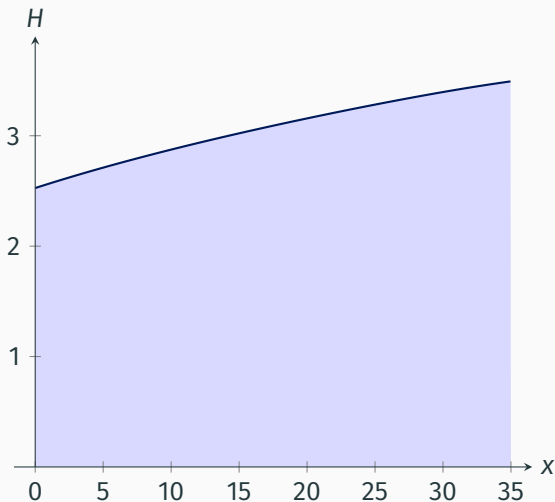
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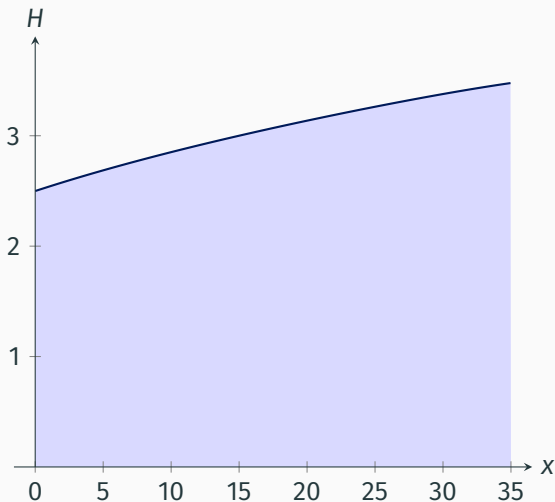
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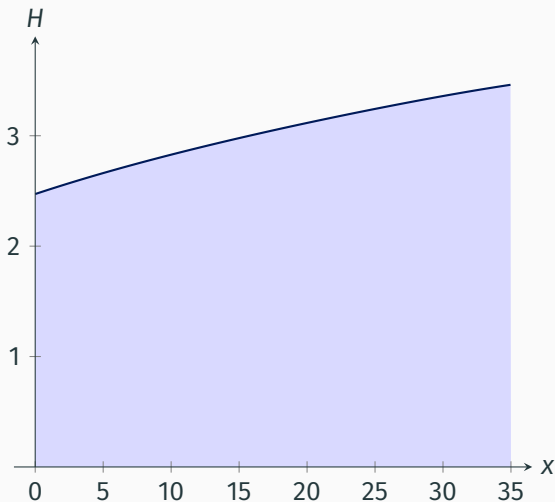
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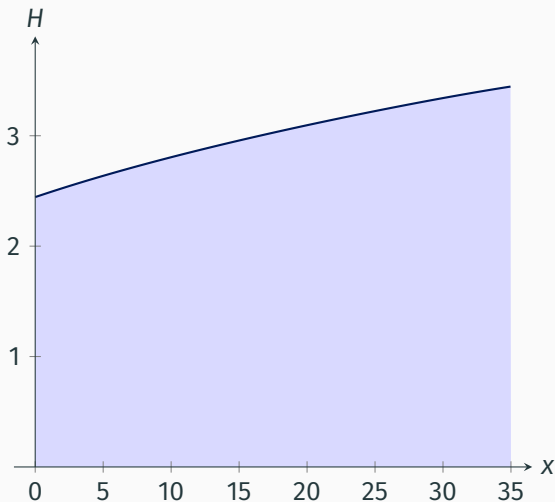
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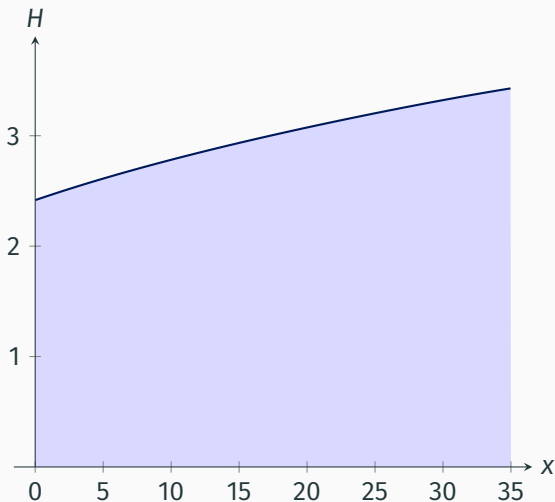
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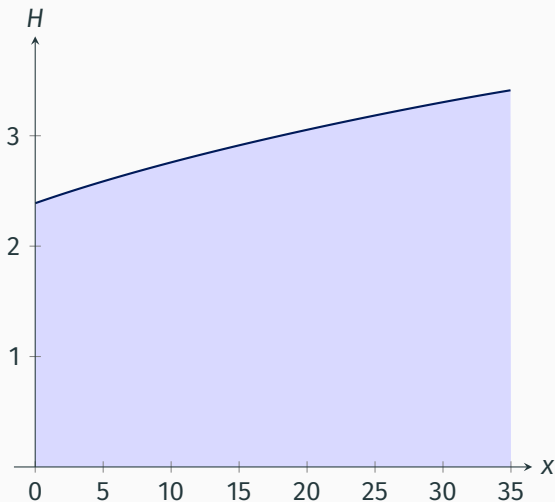
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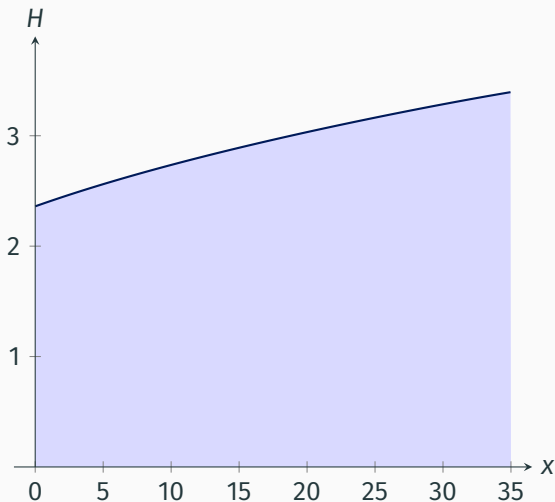
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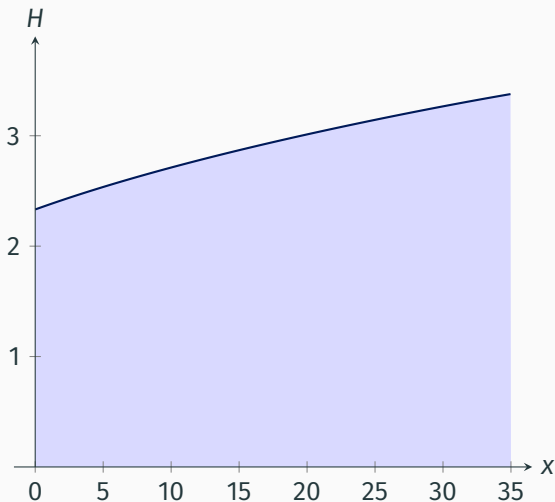
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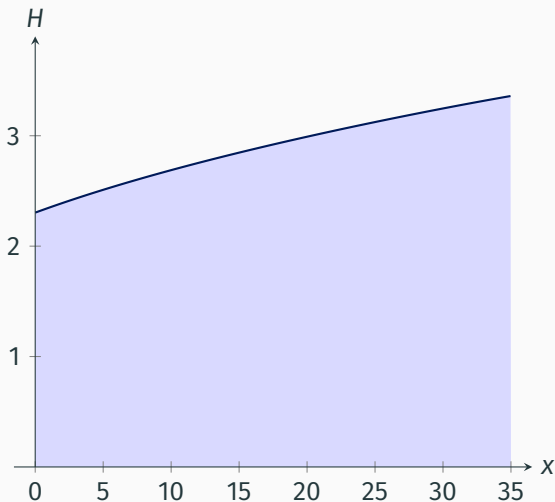
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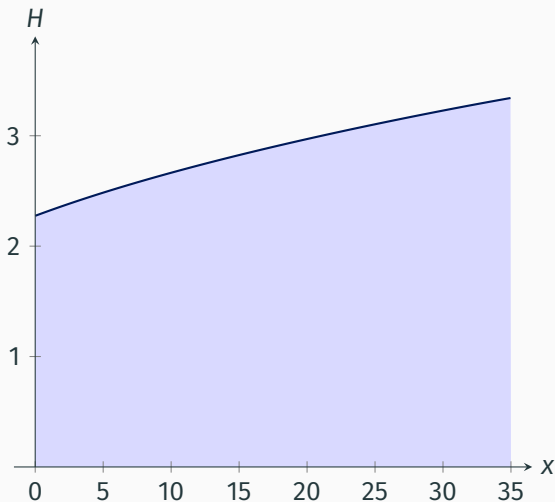
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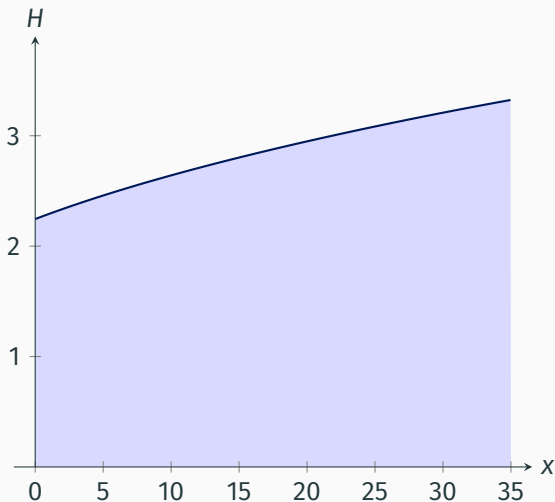
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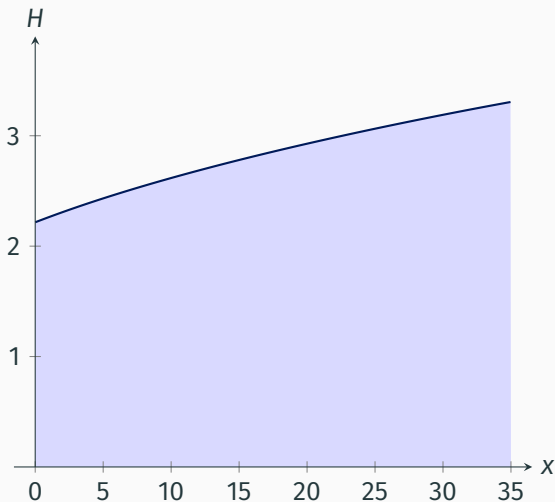
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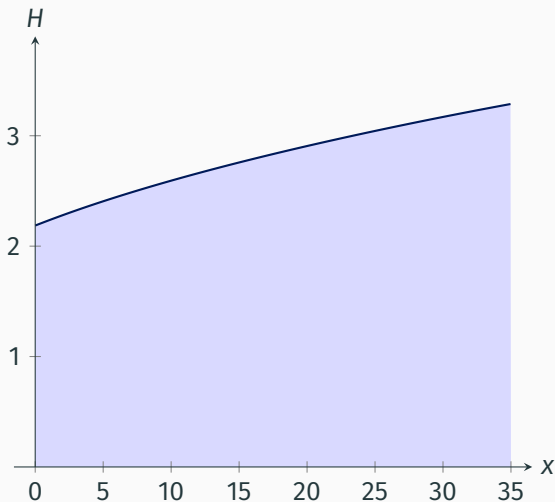
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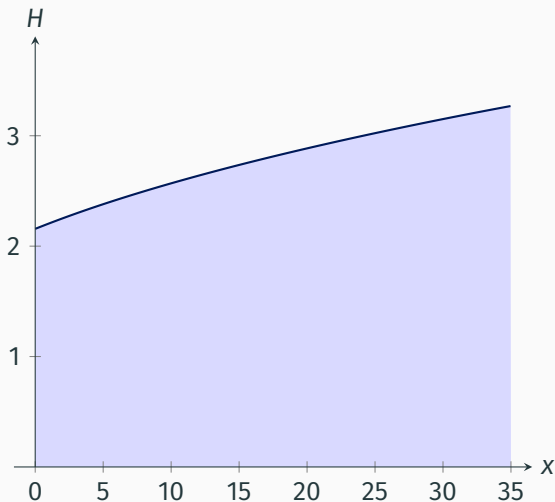
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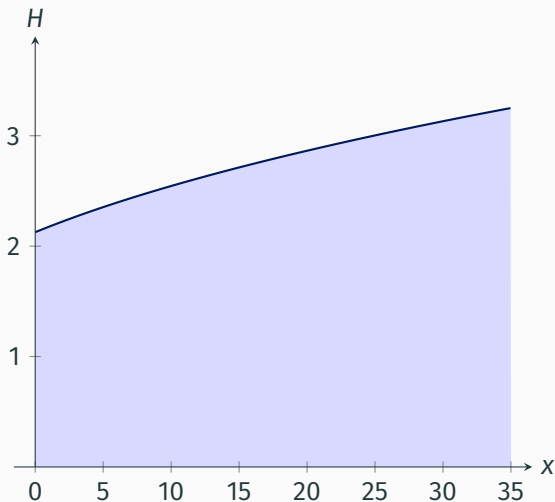
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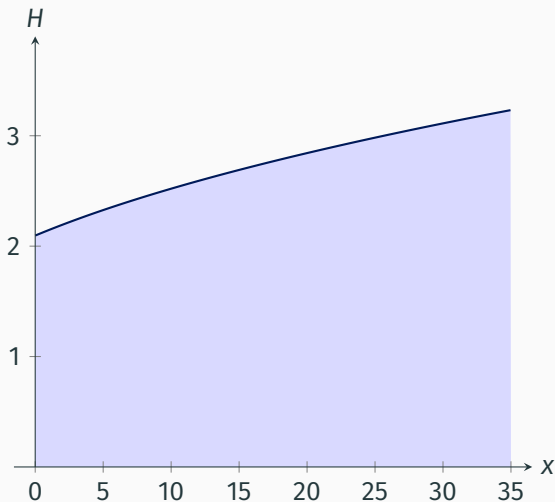
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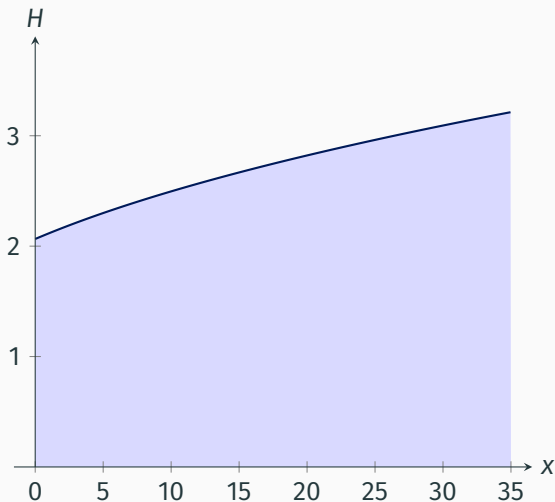
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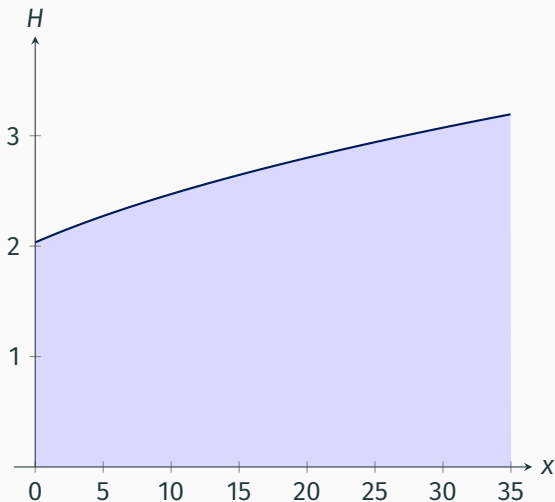
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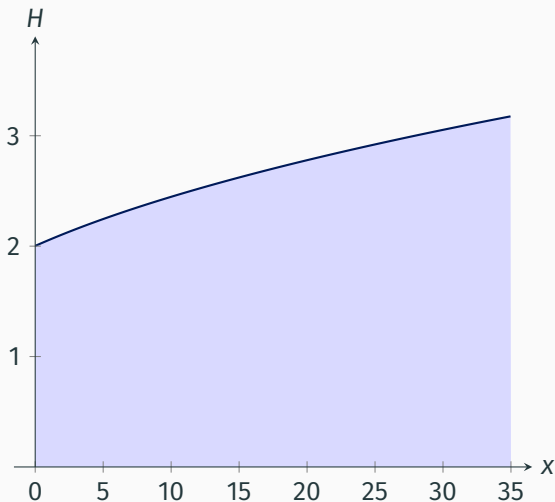
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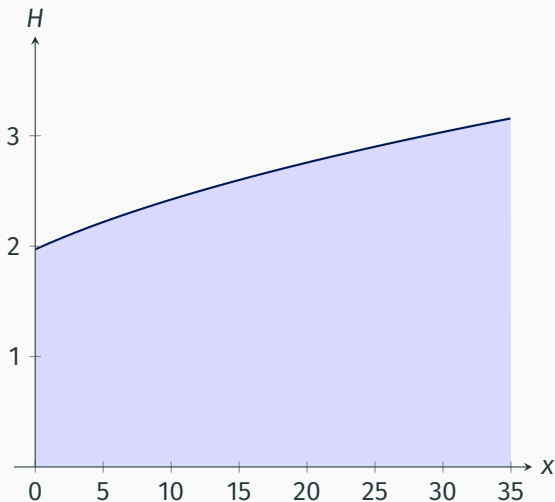
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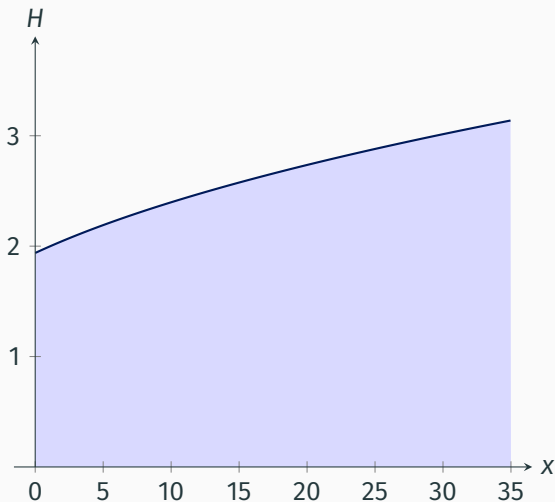
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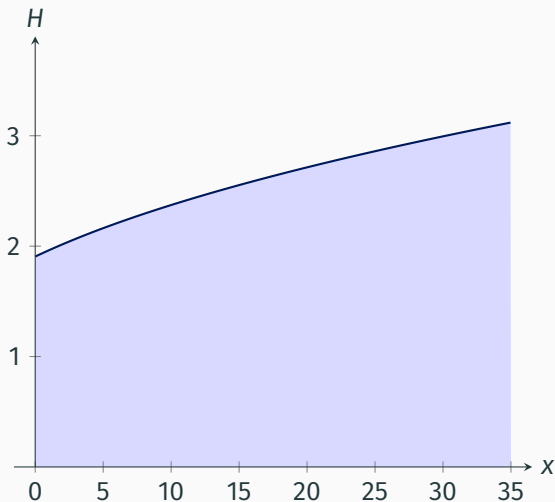
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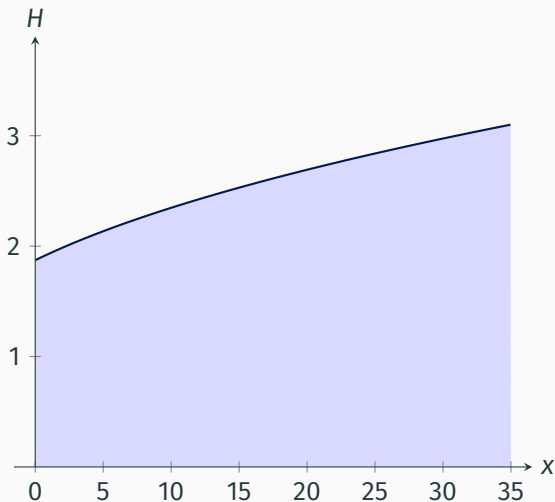
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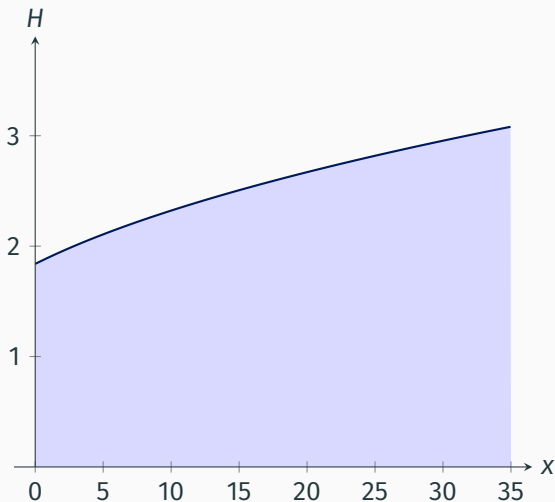
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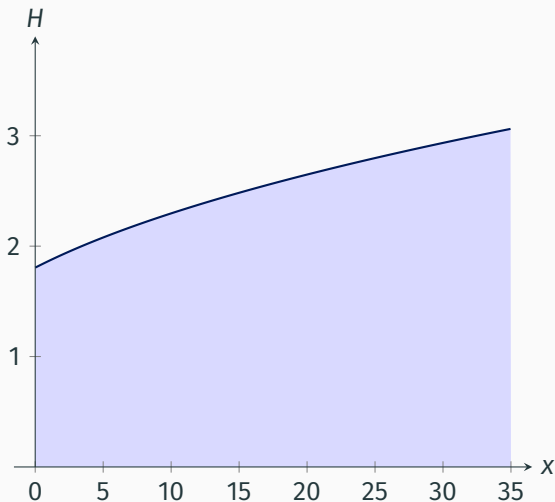
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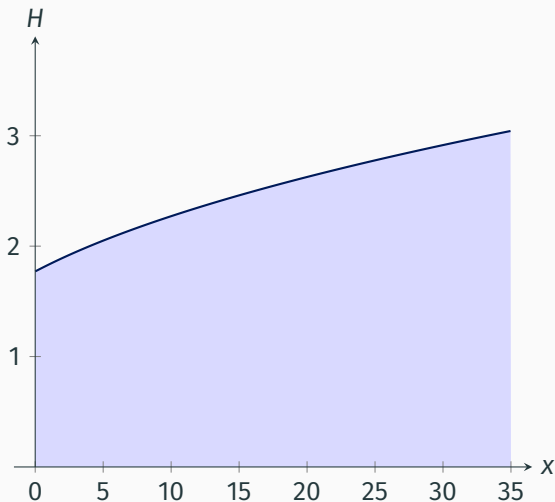
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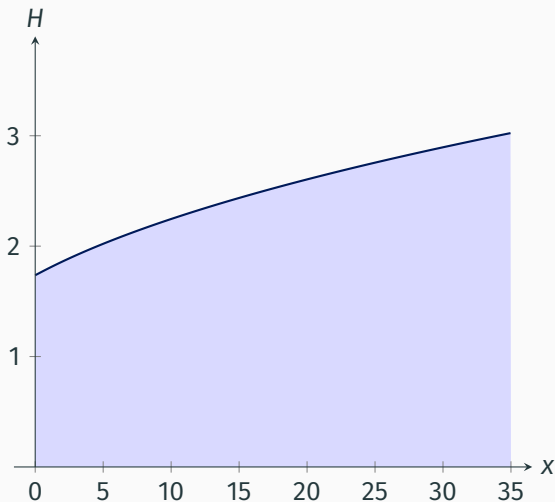
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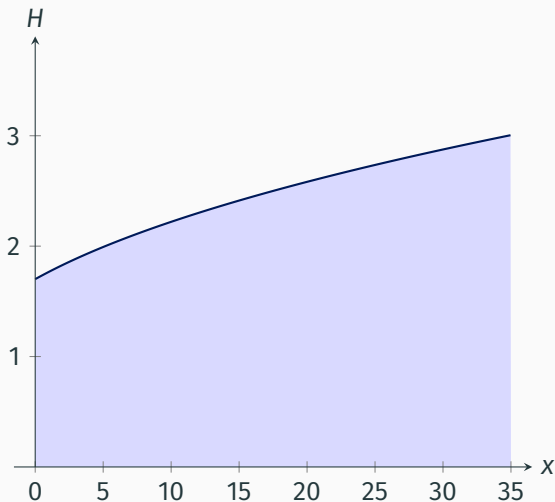
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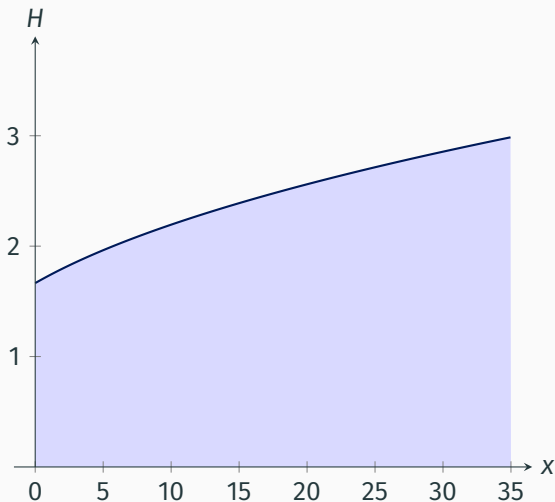
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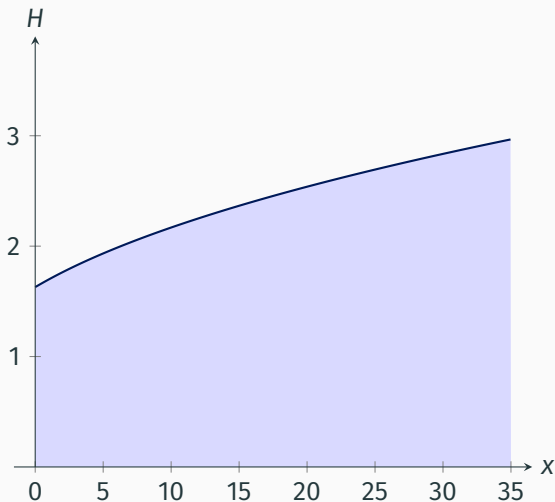
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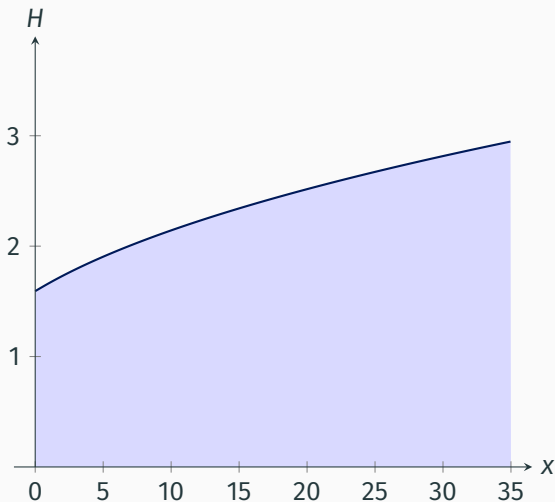
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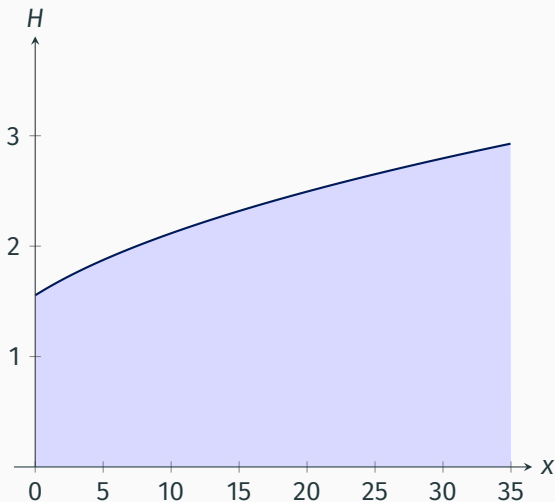
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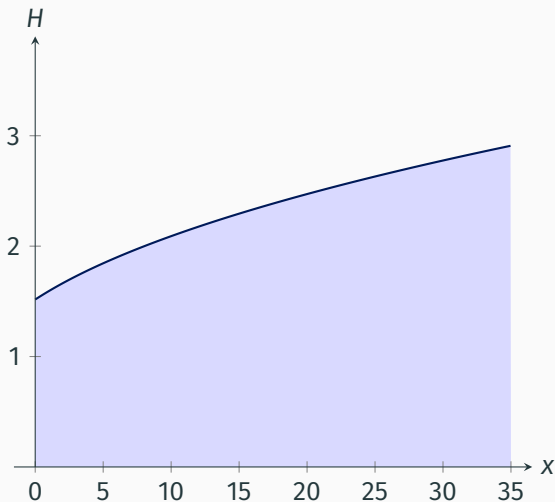
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## Unsteady flood flow

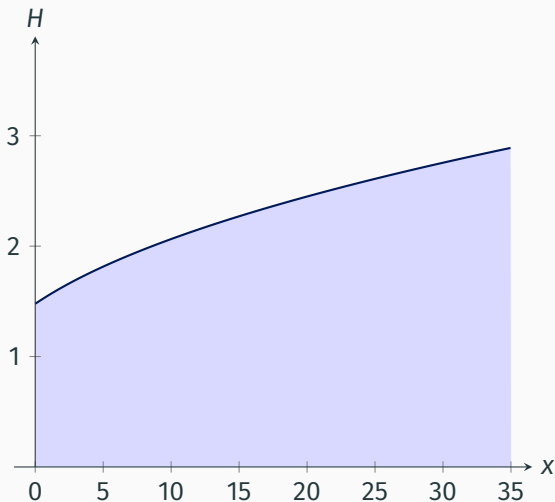
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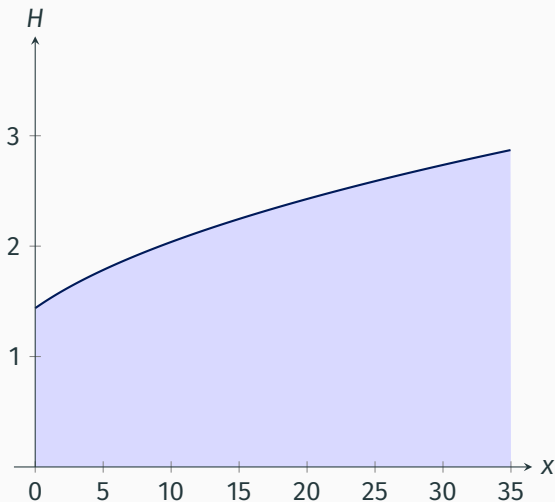
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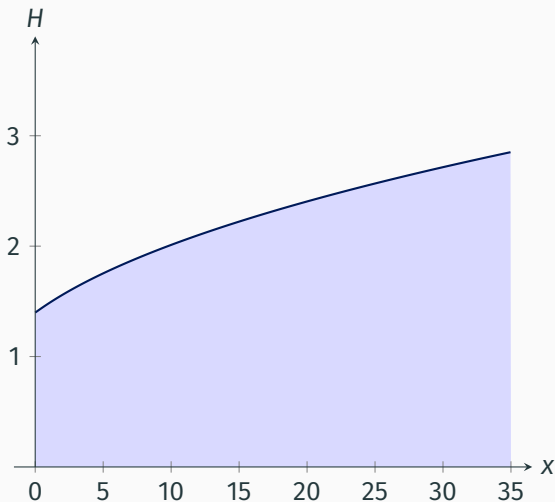
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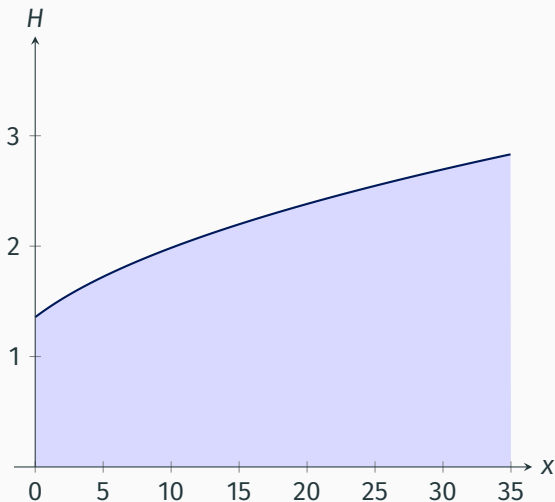
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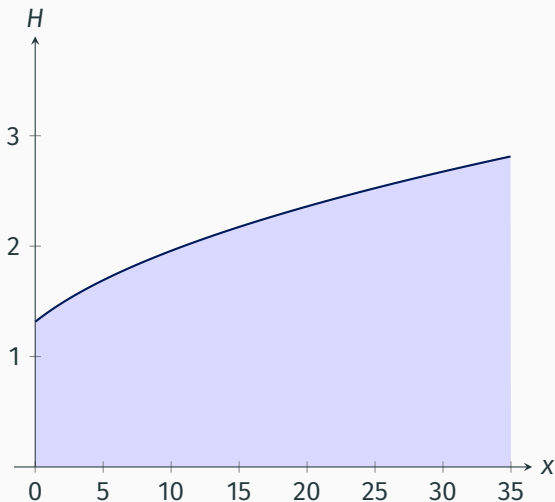
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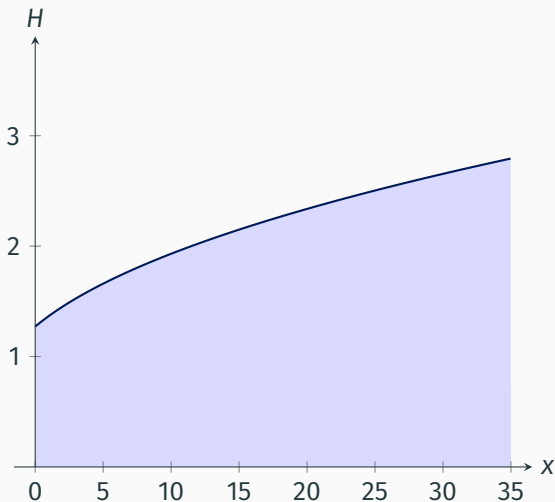
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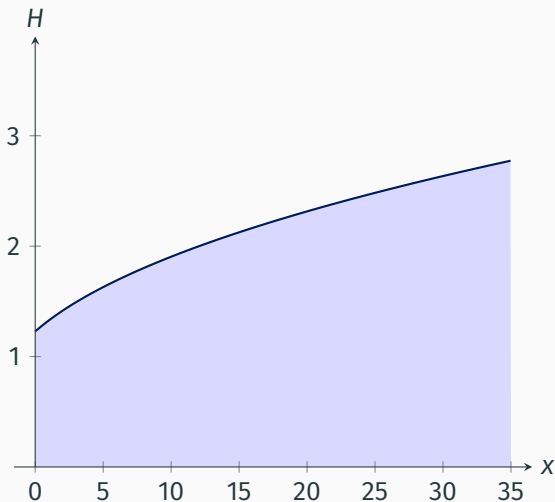
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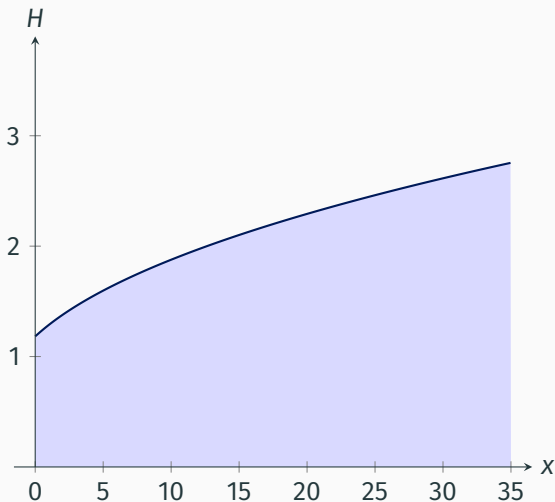
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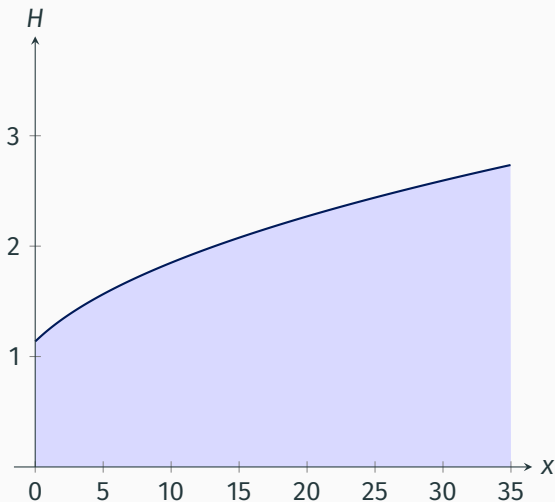
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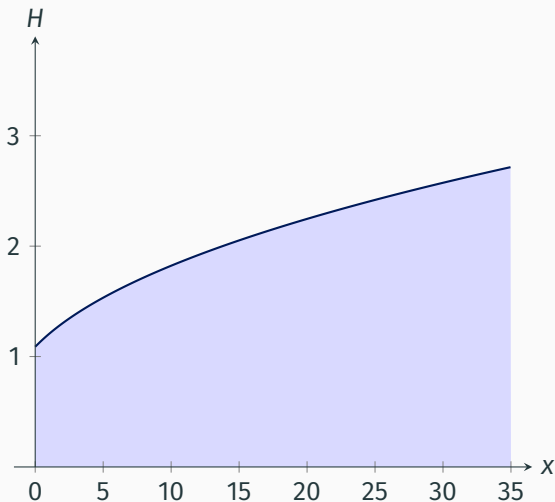
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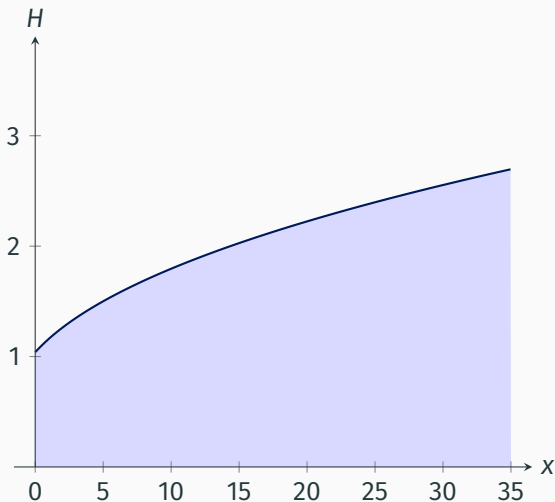
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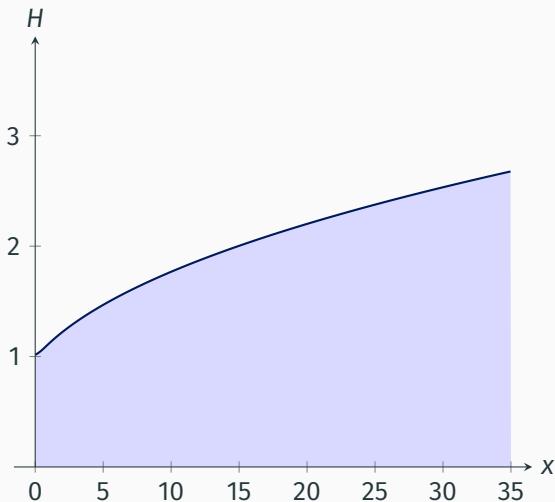
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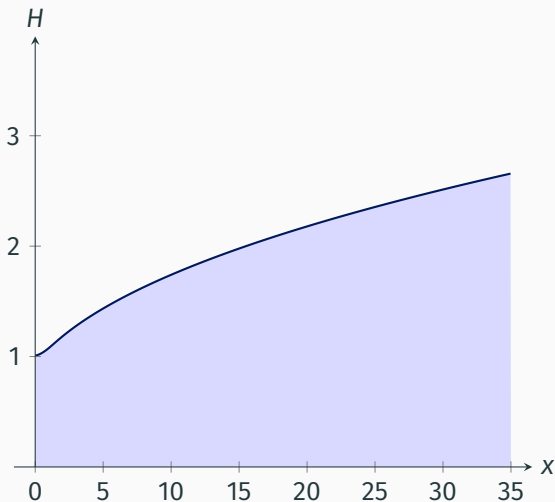
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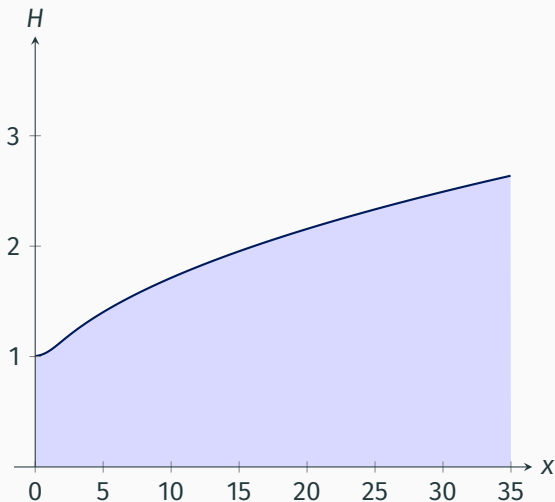
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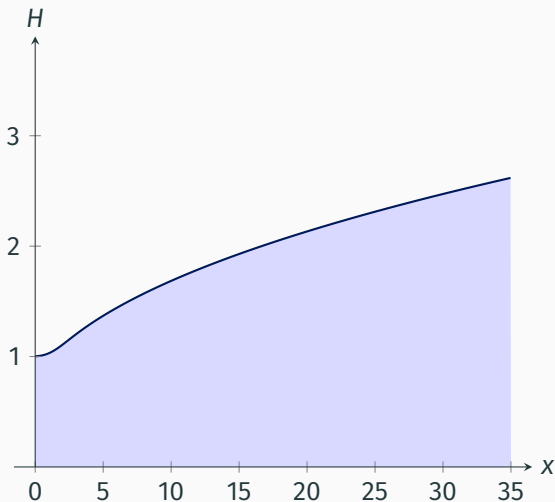
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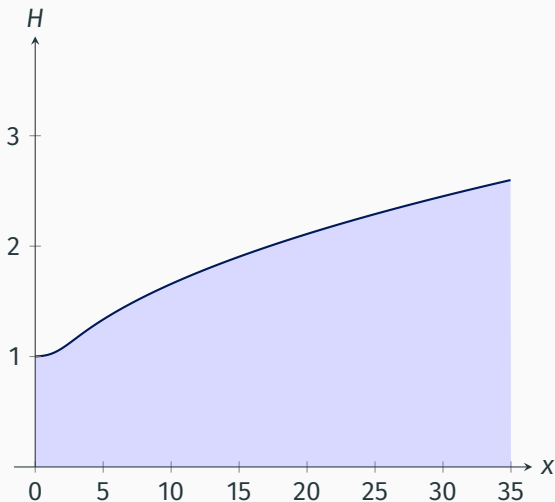
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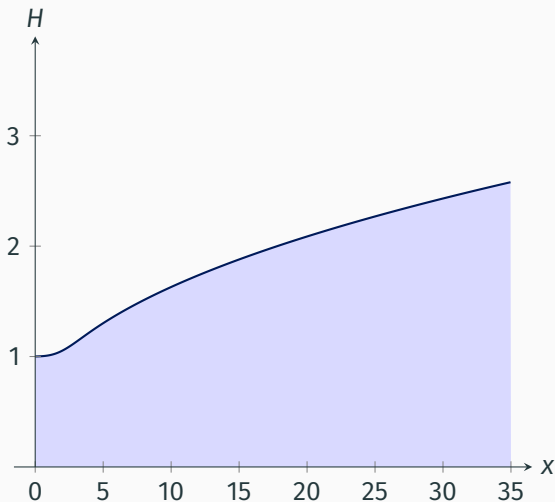
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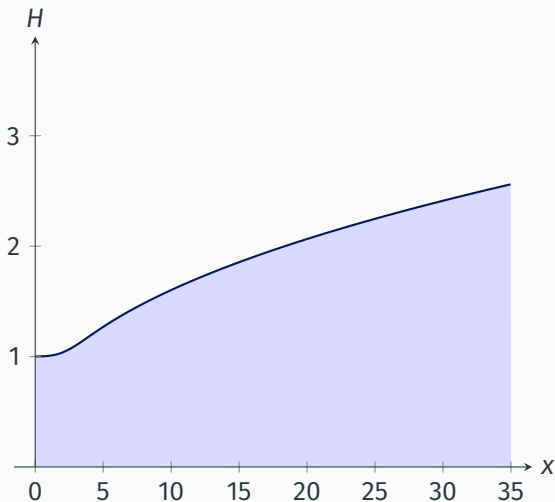
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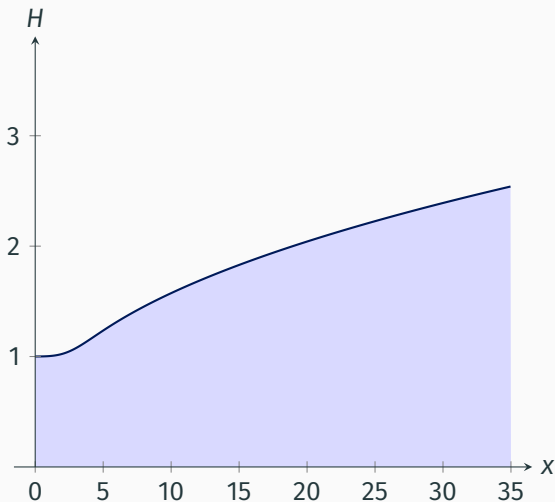
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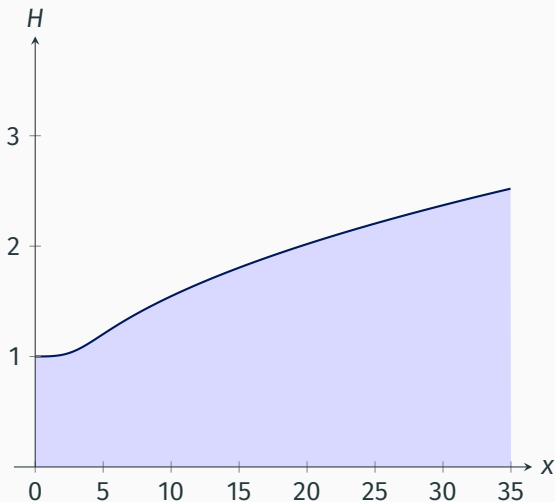
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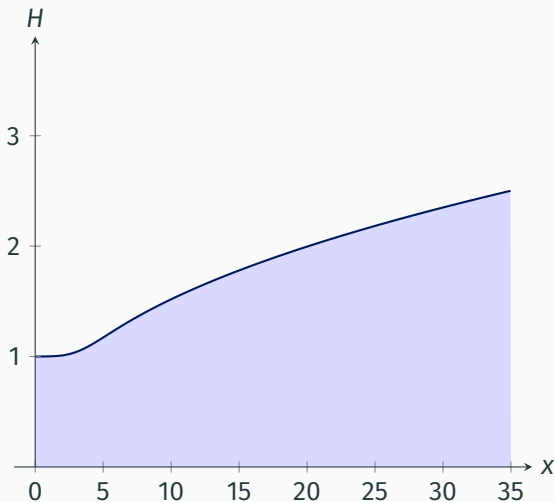
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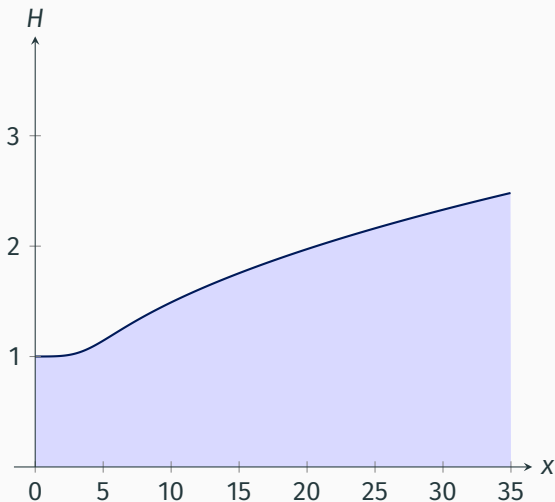
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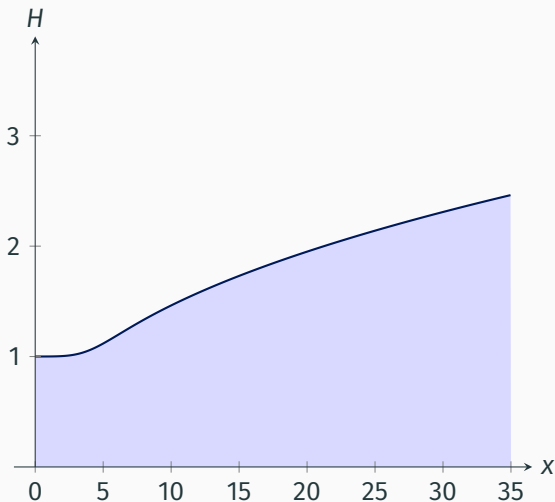
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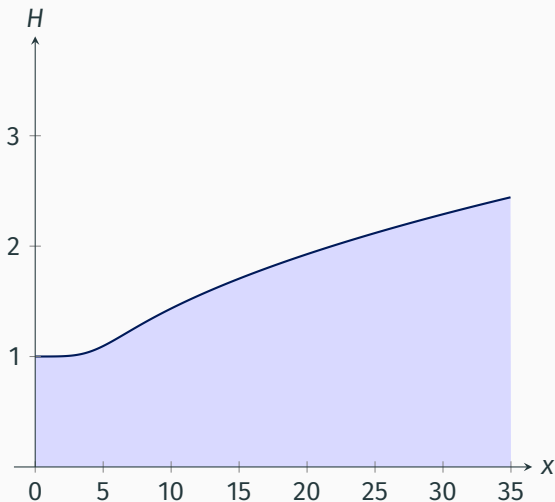
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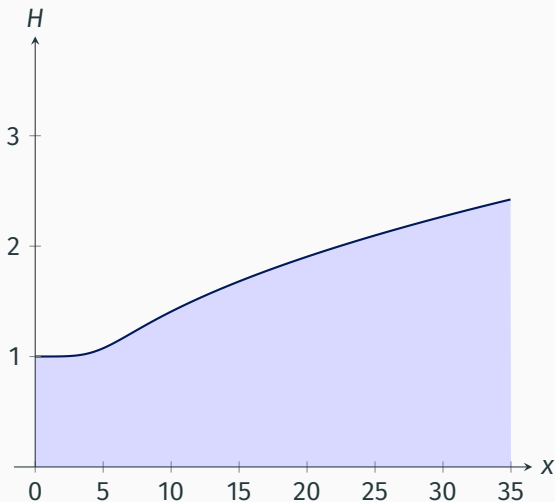
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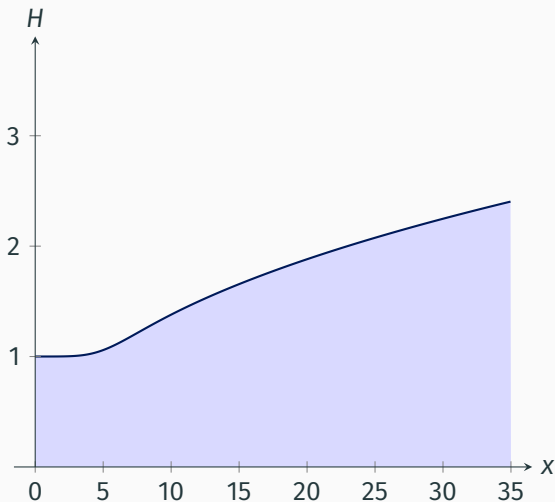
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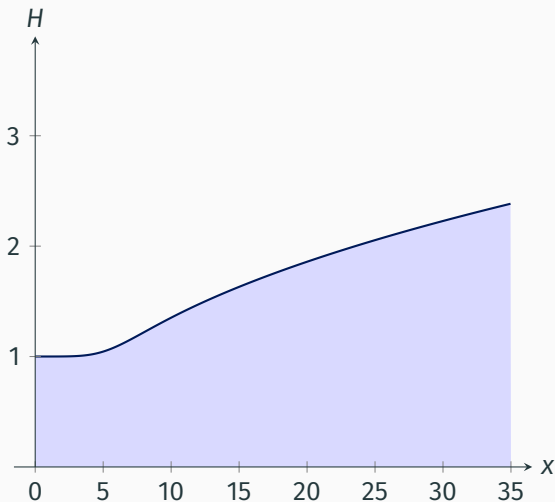
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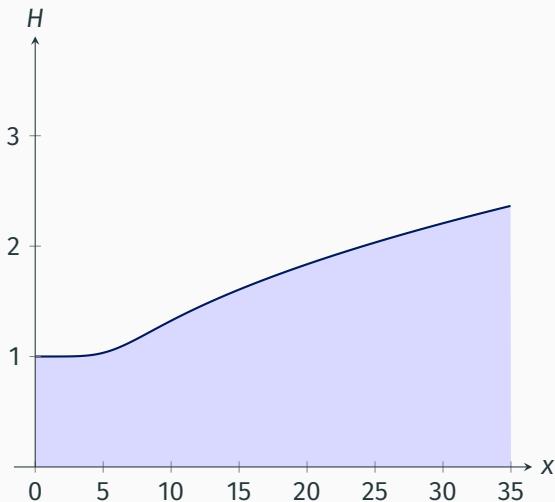
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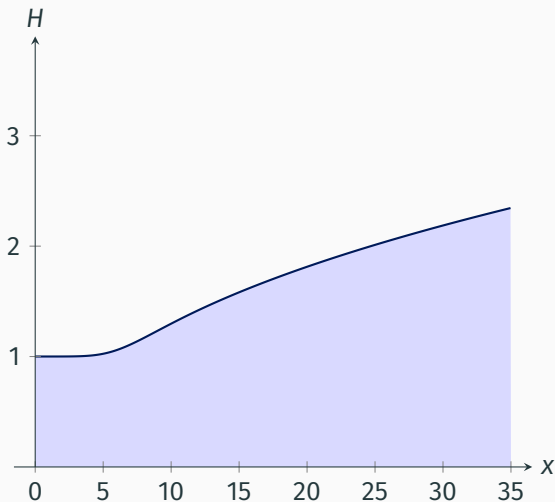
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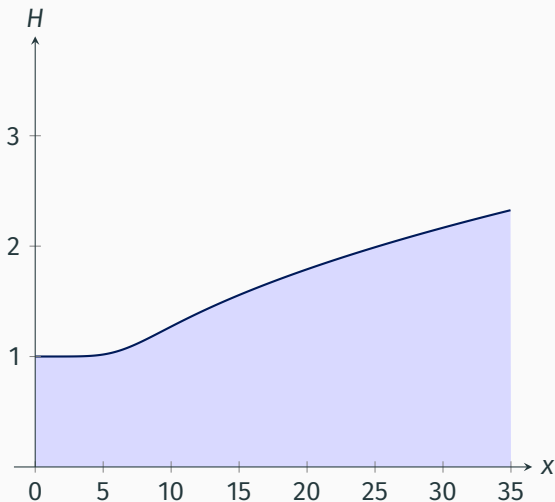
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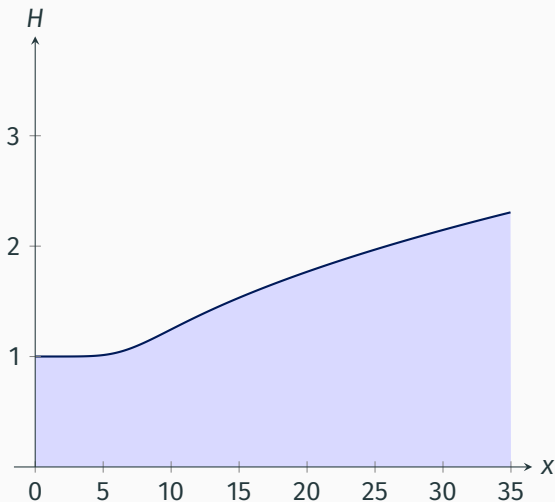
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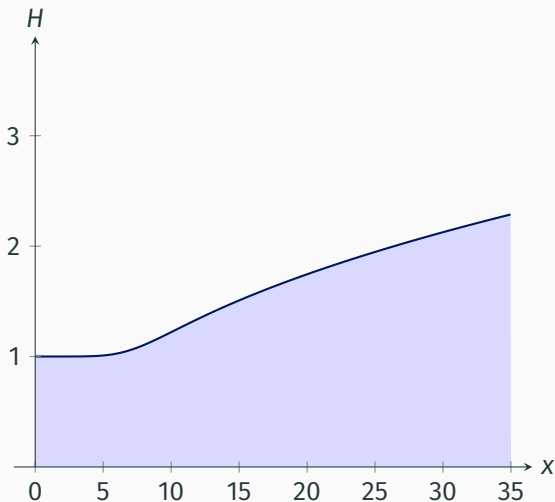
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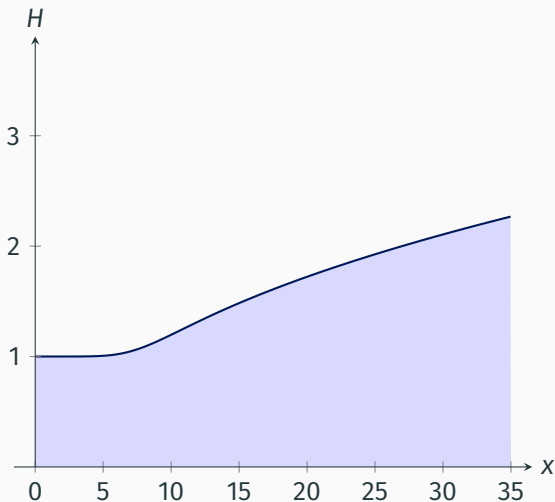
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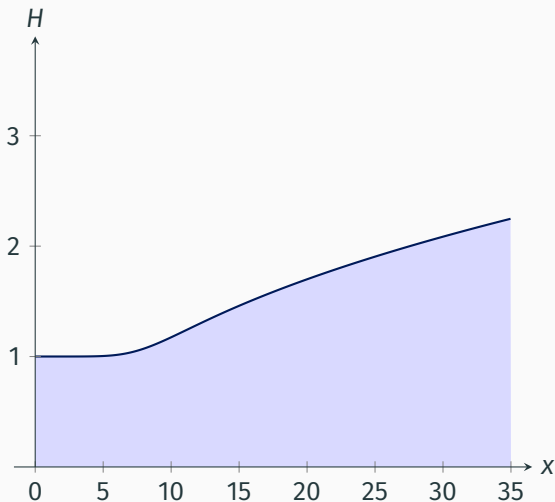
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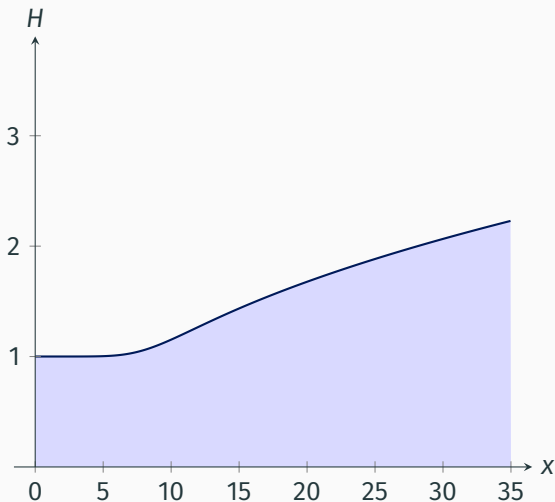
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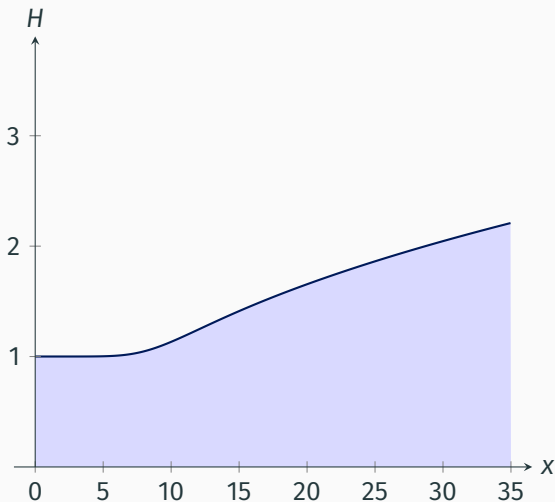
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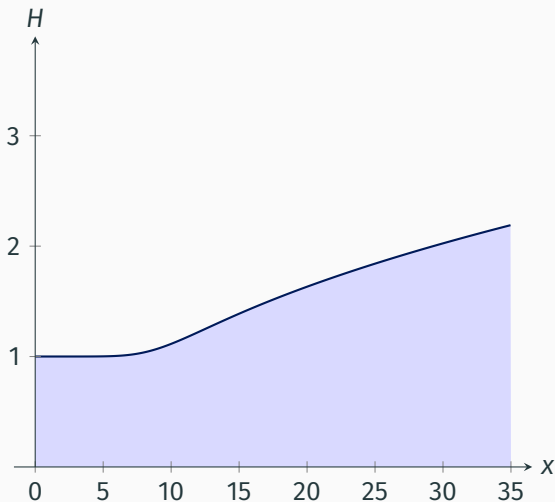
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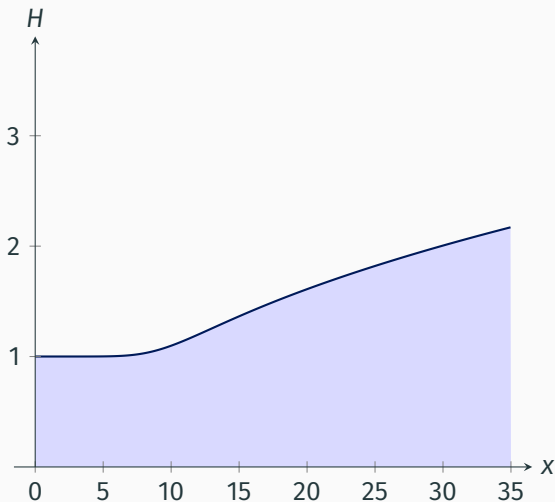
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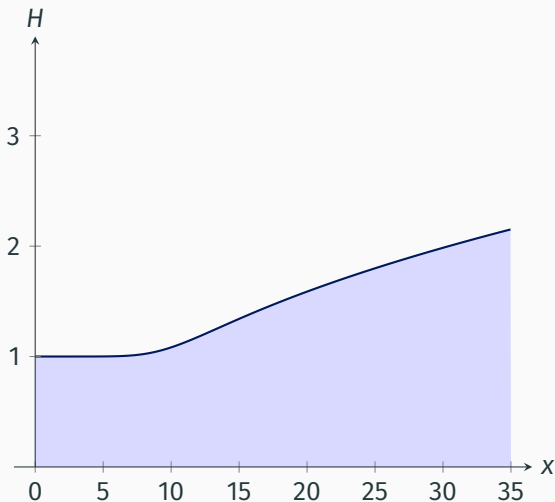
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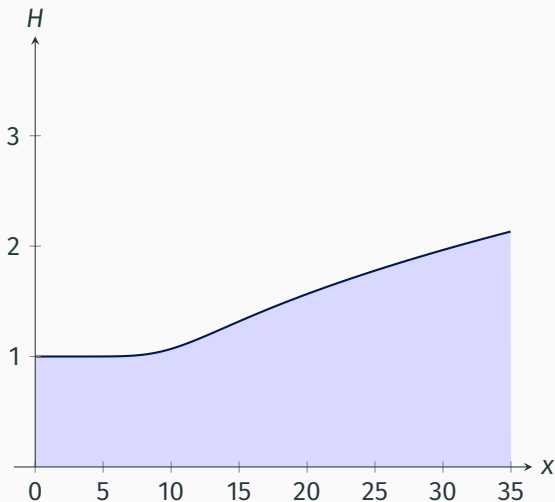
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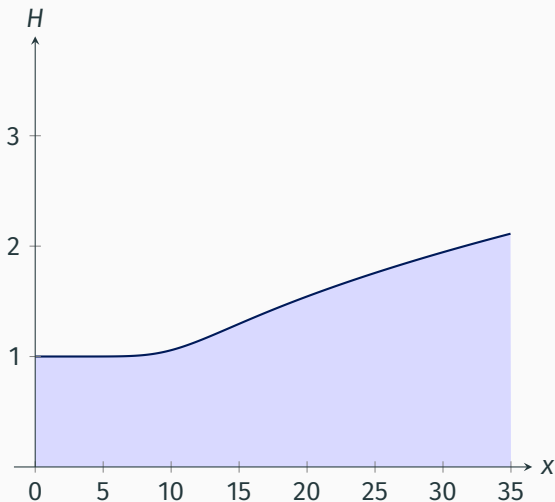
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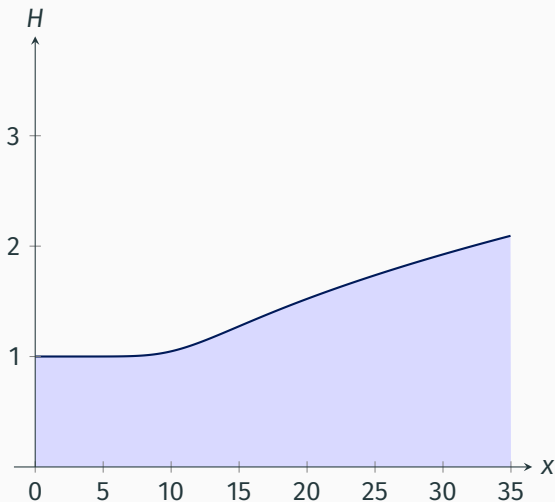
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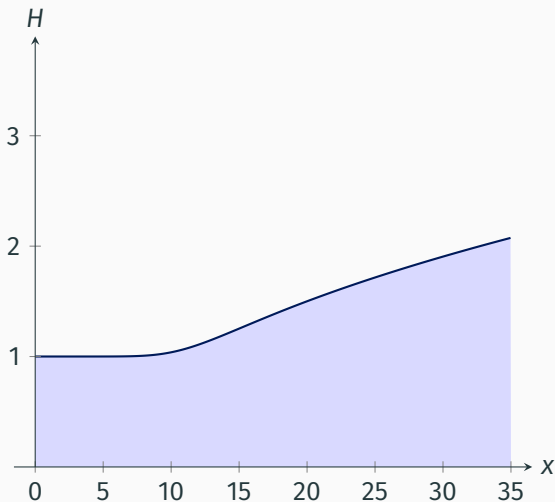
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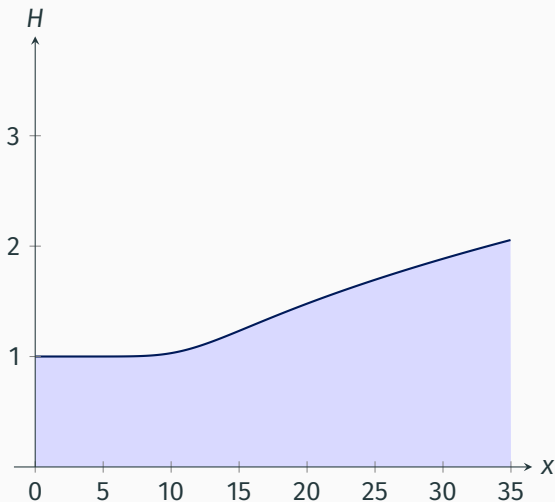
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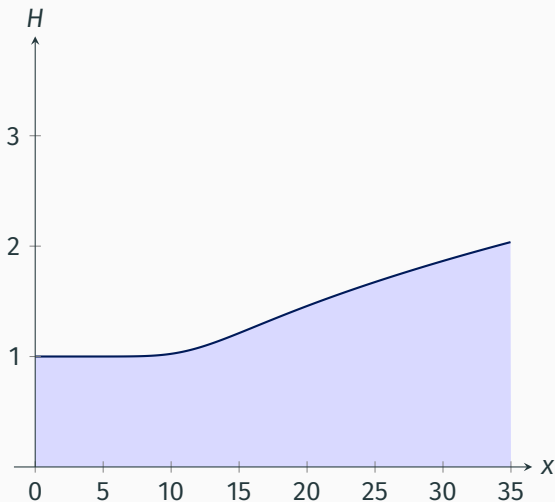
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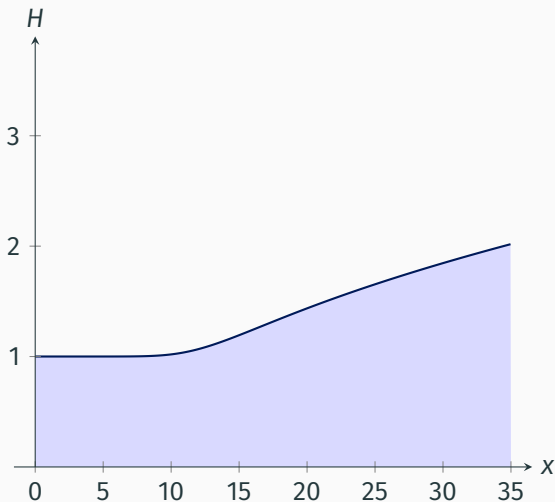
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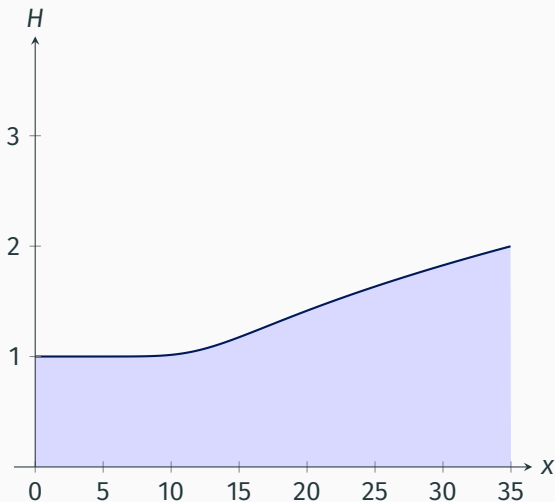
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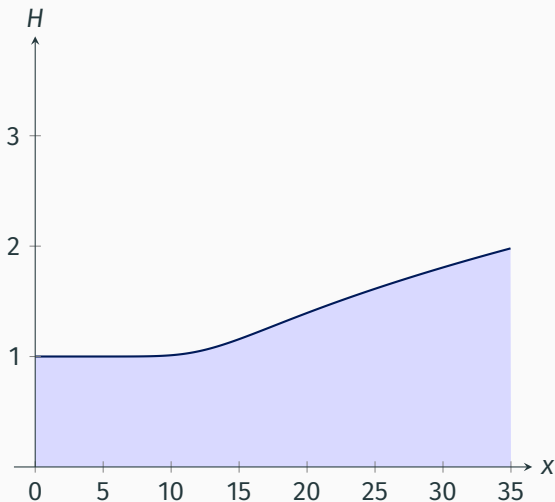
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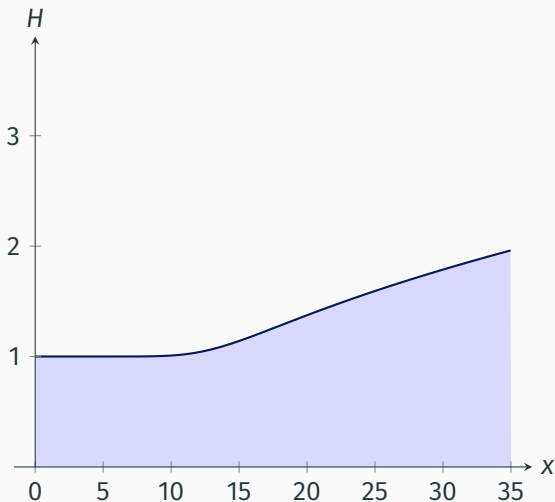
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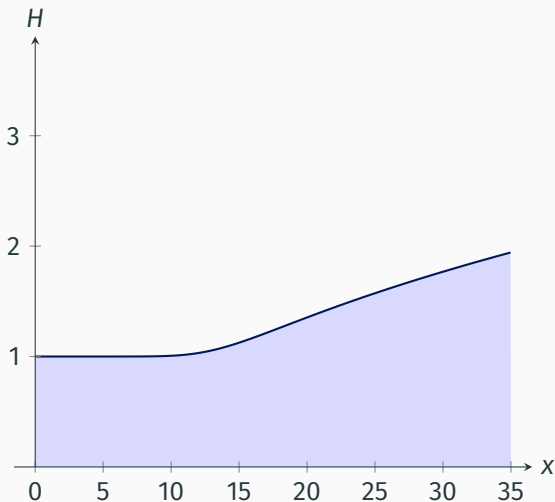
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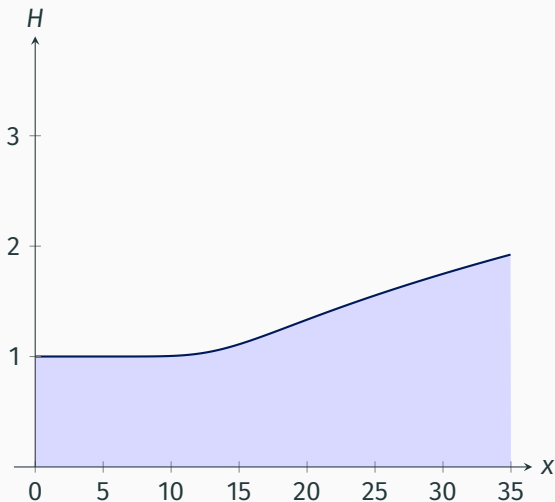
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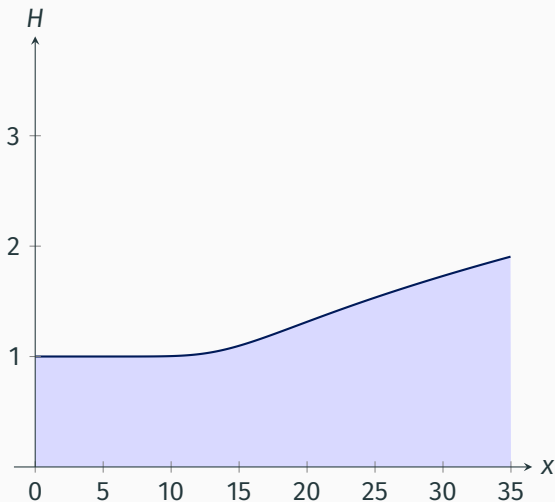
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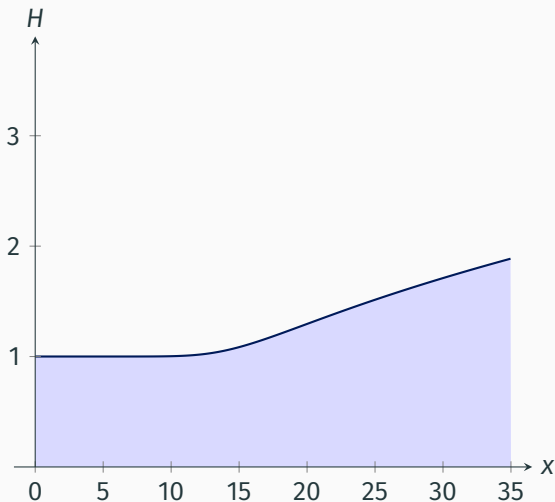
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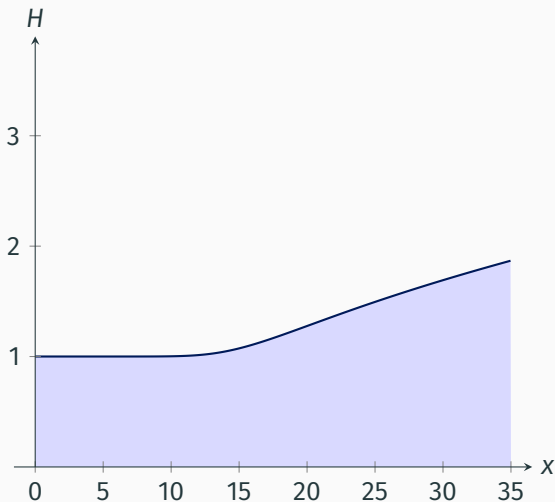
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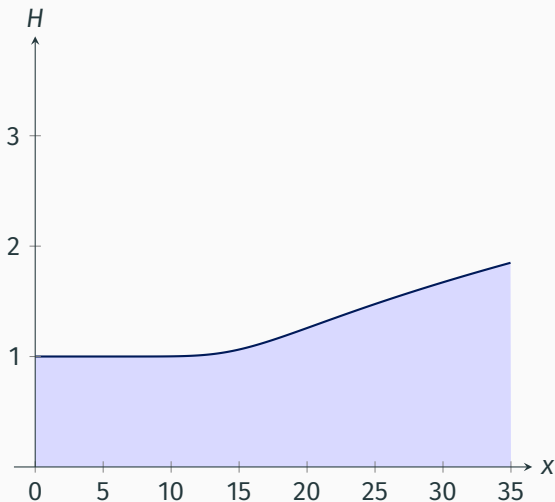
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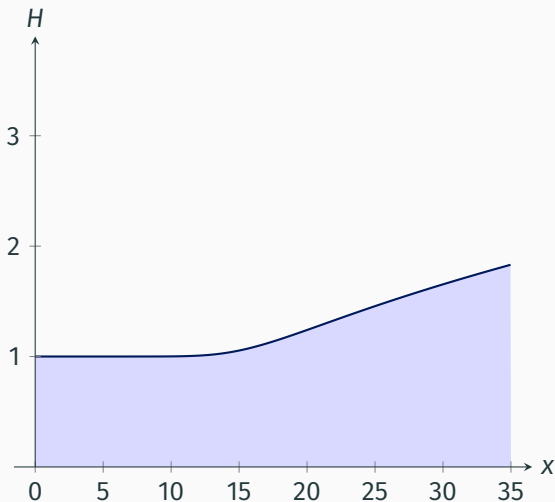
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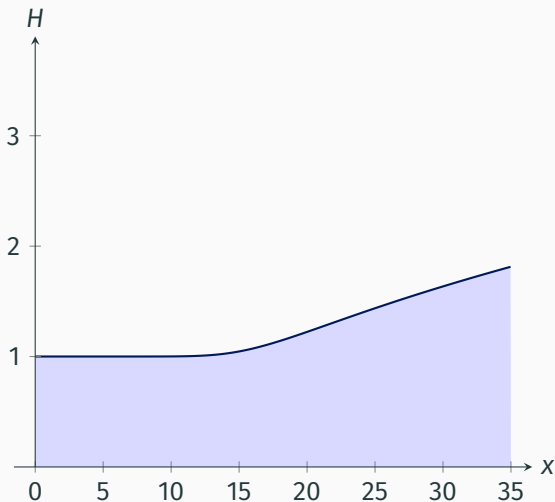
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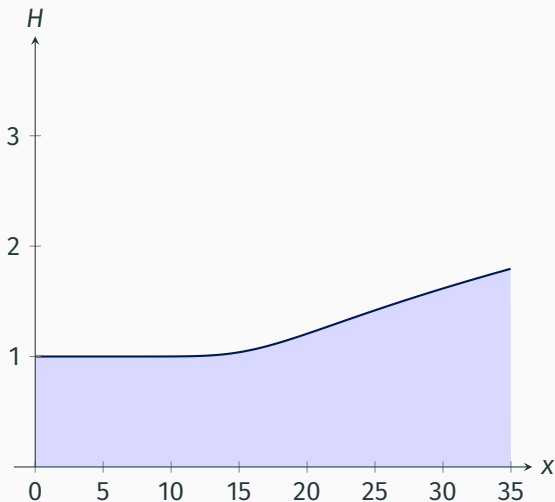
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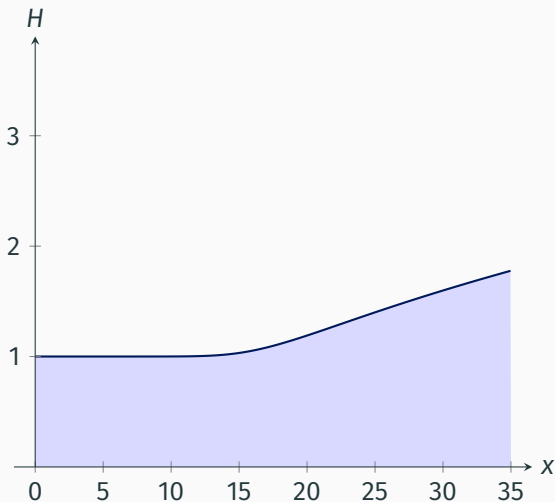
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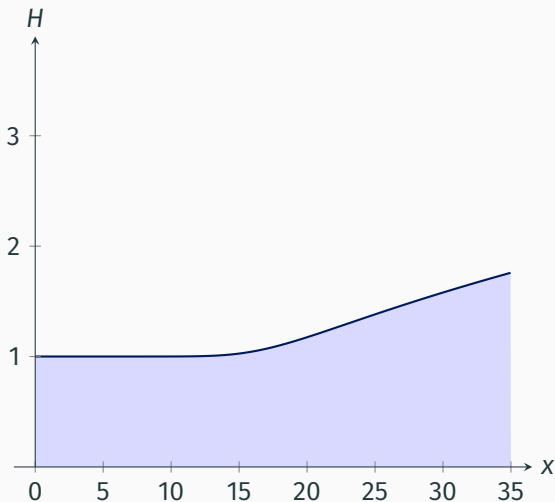
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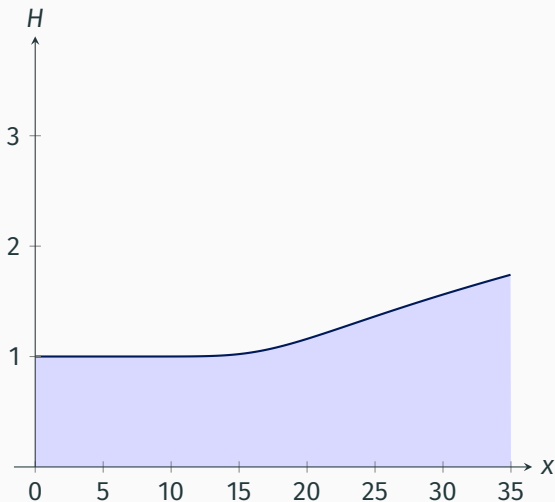
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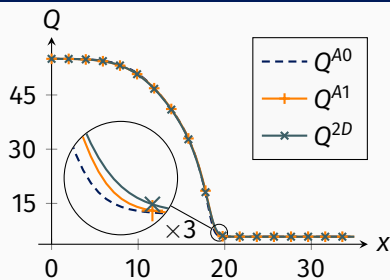
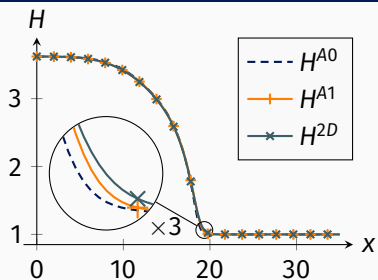


## Unsteady flood flow

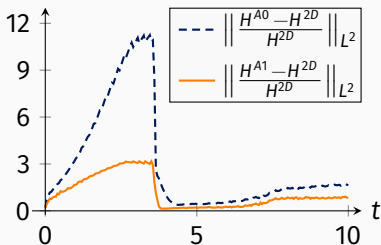
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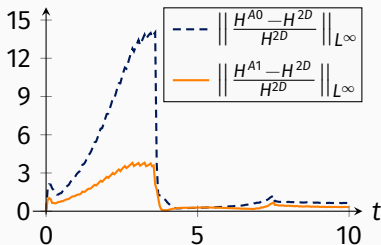
# Unsteady flood flow (2D: ref. sol., A0: 0<sup>th</sup>-order, A1: 1<sup>st</sup>-order)



error (%)



error (%)



1. Governing equations
2. Asymptotic expansions
3. Transverse averaging
4. A zeroth-order model
5. Numerical validation of the model
6. Conclusion and perspectives

# Conclusion

We have developed a **new 1D model**, based on the 2D shallow water equations, that is:

- **consistent**, up to first-order, with the 2D model in the asymptotic regime corresponding to a **river flow**:
  - ▶ the **zeroth-order** is obtained with a **new explicit friction term**,
  - ▶ the **first-order** relies on **new equations** describing the evolution of the energy;
- **hyperbolic**;
- **easily implementable** and **numerically validated**.

The preprint related to these results is available on HAL:

V. Michel-Dansac, P. Noble et J.-P. Vila, **Consistent section-averaged shallow water equations with bottom friction**, 2018.

<https://hal.archives-ouvertes.fr/hal-01962186>



## **Work related to the model:**

- improve the treatment of the river meanders by going to the first-order instead of the zeroth-order
- adapt this methodology to treat confluences
- consider a time-dependent topography to model the effects of sedimentation

## **Work related to the implementation and scientific computation:**

- compare the 1D results to the ones given by a fully 2D code, in real test cases (Garonne, Lèze, Gironde, Amazon, ...)
- couple the 1D and 2D equations in the context of the Gironde estuary (collaboration in progress with the SHOM)

Thank you for your attention!

# First-order model

The first-order model is:

$$\left\{ \begin{array}{l} S_t + Q_x = 0, \\ Q_t + \left( \frac{Q^2}{S} + \Psi \right)_x + \left( 1 - \frac{S\Psi_{2D}^{(0)}}{(Q_{2D}^{(0)})^2} \right) \frac{SH_x}{F^2} = \frac{1}{\varepsilon} S \left( \mathcal{J} - \mathcal{J} - \frac{S\Psi_{2D}^{(0)}}{(Q_{2D}^{(0)})^2} (\mathcal{J} - \mathcal{J}_\Psi) \right), \\ \left( \frac{1}{2} \frac{Q^2}{S} + \frac{1}{2} \Psi \right)_t + \left( \frac{Q}{S} \left( \frac{1}{2} \frac{Q^2}{S} + \frac{1}{2} \Pi \right) \right)_x + \frac{QH_x}{F^2} = \frac{1}{\varepsilon} Q (\mathcal{J} - \mathcal{J}), \\ \left( \frac{1}{2} (\Pi - 3\Psi) \right)_t = \frac{1}{\varepsilon} Q \frac{S\Pi_{2D}^{(0)}}{(Q_{2D}^{(0)})^2} (\mathcal{J}_\Psi - \mathcal{J}_\Pi). \end{array} \right.$$

It ensures the correct asymptotic regime, that is to say

$$Q = Q_{2D}^{(0)} + \varepsilon Q_{2D}^{(1)} + \mathcal{O}(\varepsilon^2).$$

In addition, it is hyperbolic and linearly stable.

# Non-dimensional form of the 2D shallow water system

To emphasize the different scales of the flow, we perform a non-dimensionalization of the 2D system.

We introduce the following dimensionalization scales and related non-dimensional quantities (which are denoted with a bar, like  $\bar{x}$ ):

$$h := \mathcal{H}\bar{h}, \quad u := \mathcal{U}\bar{u}, \quad v := \mathcal{V}\bar{v}, \quad x := \mathcal{X}\bar{x}, \quad y := \mathcal{Y}\bar{y}, \quad t := \mathcal{T}\bar{t}, \quad \mathcal{T} := \frac{\mathcal{X}}{\mathcal{U}}.$$

The mass conservation equation

$$\frac{\partial h}{\partial t} + \frac{\partial hu}{\partial x} + \frac{\partial hv}{\partial y} = 0$$

then becomes

$$\frac{\mathcal{H}}{\mathcal{T}} \frac{\partial \bar{h}}{\partial \bar{t}} + \frac{\mathcal{H}\mathcal{U}}{\mathcal{X}} \frac{\partial \bar{h}\bar{u}}{\partial \bar{x}} + \frac{\mathcal{H}\mathcal{V}}{\mathcal{Y}} \frac{\partial \bar{h}\bar{v}}{\partial \bar{y}} = 0.$$

# Non-dimensional form of the 2D shallow water system

The non-dimensional conservation equation is

$$\frac{\mathcal{H}}{\mathcal{T}} \frac{\partial \bar{h}}{\partial \bar{t}} + \frac{\mathcal{H}\mathcal{U}}{\mathcal{X}} \frac{\partial \bar{h}\bar{u}}{\partial \bar{x}} + \frac{\mathcal{H}\mathcal{V}}{\mathcal{Y}} \frac{\partial \bar{h}\bar{v}}{\partial \bar{y}} = 0, \text{ i.e. } \frac{\partial \bar{h}}{\partial \bar{t}} + \frac{\partial \bar{h}\bar{u}}{\partial \bar{x}} + \frac{\mathcal{V}}{\mathcal{U}} \frac{\mathcal{X}}{\mathcal{Y}} \frac{\partial \bar{h}\bar{v}}{\partial \bar{y}} = 0.$$

We set  $R_u := \mathcal{V}/\mathcal{U}$  and  $R_x := \mathcal{Y}/\mathcal{X}$ , to get

$$\frac{\partial \bar{h}}{\partial \bar{t}} + \frac{\partial \bar{h}\bar{u}}{\partial \bar{x}} + \frac{R_u}{R_x} \frac{\partial \bar{h}\bar{v}}{\partial \bar{y}} = 0.$$

We have

- $\mathcal{V} \ll \mathcal{U}$  (quasi-unidimensional flow)  $\implies R_u \ll 1$ ,
- $\mathcal{Y} \ll \mathcal{X}$  (quasi-unidimensional geometry)  $\implies R_x \ll 1$ .

We assume  $R_u = R_x$  to keep the mass conservation equation unchanged from the dimensional case.

# Non-dimensional form of the 2D shallow water system

Regarding the geometry, we assume that  $Z(x, y) = b(x) + \phi(x, y)$ , where:

- $b(x)$  represents the main **longitudinal topography**, driving the flow from upstream to downstream;
- $\phi(x, y)$  represents **small** longitudinal and transverse **variations**.

The related non-dimensional quantities are

$$b = \mathcal{B}\bar{b}\left(\frac{x}{\mathcal{X}}\right) \quad \text{and} \quad \phi = \mathcal{H}\bar{\phi}\left(\frac{x}{\mathcal{X}}, \frac{y}{\mathcal{Y}}\right).$$

The non-dimensional topography gradient then reads:

$$\nabla Z = \begin{pmatrix} \frac{\mathcal{B}}{\mathcal{X}} \frac{\partial \bar{b}}{\partial \bar{x}}(\bar{x}) + \frac{\mathcal{H}}{\mathcal{X}} \frac{\partial \bar{\phi}}{\partial \bar{x}}(\bar{x}, \bar{y}) \\ \frac{\mathcal{H}}{\mathcal{Y}} \frac{\partial \bar{\phi}}{\partial \bar{y}}(\bar{x}, \bar{y}) \end{pmatrix}.$$

# Non-dimensional form of the 2D shallow water system

Regarding the friction, we take  $C_h = \mathcal{C} \bar{C}(\bar{x}, \bar{y})$ .

The non-dimensional friction source term then reads:

$$\frac{\mathbf{u} \|\mathbf{u}\|}{C_h^2 h^p} = \begin{pmatrix} \frac{u}{\mathcal{C}\mathcal{H}^p} \cdot \frac{\bar{u} \sqrt{\mathcal{U}^2 \bar{u}^2 + \mathcal{V}^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p} \\ \frac{v}{\mathcal{C}\mathcal{H}^p} \cdot \frac{\bar{v} \sqrt{\mathcal{U}^2 \bar{u}^2 + \mathcal{V}^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p} \end{pmatrix} = \begin{pmatrix} \frac{\mathcal{U} |\mathcal{U}|}{\mathcal{C}\mathcal{H}^p} \cdot \frac{\bar{u} \sqrt{\bar{u}^2 + R_U^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p} \\ \frac{\mathcal{V} |\mathcal{U}|}{\mathcal{C}\mathcal{H}^p} \cdot \frac{\bar{v} \sqrt{\bar{u}^2 + R_U^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p} \end{pmatrix}.$$

# Non-dimensional form of the 2D shallow water system

We are finally able to write the non-dimensional form of the 2D shallow water system: from the dimensional system

$$\begin{cases} h_t + \nabla \cdot (h\mathbf{u}) = 0, \\ \mathbf{u}_t + \mathbf{u} \cdot \nabla \mathbf{u} + g\nabla h = g\left(-\nabla Z - \frac{\mathbf{u}\|\mathbf{u}\|}{C_h^2 h^p}\right), \end{cases}$$

we get the following non-dimensional form:

$$\begin{cases} \bar{h}_{\bar{t}} + (\bar{h}\bar{u})_{\bar{x}} + (\bar{h}\bar{v})_{\bar{y}} = 0, \\ \frac{\mathcal{U}^2}{\mathcal{X}} \bar{u}_{\bar{t}} + \frac{\mathcal{U}^2}{\mathcal{X}} \bar{u}\bar{u}_{\bar{x}} + \frac{\mathcal{U}\mathcal{V}}{\mathcal{Y}} \bar{v}\bar{u}_{\bar{y}} + \frac{g\mathcal{H}}{\mathcal{X}} (\bar{h} + \bar{\Phi})_{\bar{x}} = g\left(-\frac{\mathcal{U}|\mathcal{U}|}{C_h^p} \frac{\bar{u}\sqrt{\bar{u}^2 + R_u^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p} - \frac{\mathcal{B}}{\mathcal{X}} \bar{b}_{\bar{x}}\right), \\ \frac{\mathcal{V}\mathcal{U}}{\mathcal{X}} \bar{v}_{\bar{t}} + \frac{\mathcal{V}\mathcal{U}}{\mathcal{X}} \bar{u}\bar{v}_{\bar{x}} + \frac{\mathcal{V}^2}{\mathcal{Y}} \bar{v}\bar{v}_{\bar{y}} + \frac{g\mathcal{H}}{\mathcal{Y}} (\bar{h} + \bar{\Phi})_{\bar{y}} = g\left(-\frac{\mathcal{V}|\mathcal{U}|}{C_h^p} \frac{\bar{v}\sqrt{\bar{u}^2 + R_u^2 \bar{v}^2}}{\bar{C}^2 \bar{h}^p}\right). \end{cases}$$



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# Non-dimensional form of the 2D shallow water system

We introduce:

- $F^2 = \frac{u^2}{g\mathcal{H}}$  the reference Froude number,
- $\delta = \frac{\mathcal{H}}{\mathcal{X}}$  the shallow water parameter,
- $l_0 = \frac{\mathcal{B}}{\mathcal{X}}$  and  $J_0 = \frac{u|u|}{c\mathcal{H}^p}$  the topography and friction slopes.

With  $\frac{g\mathcal{X}}{u^2} = \frac{g\mathcal{H}}{u^2} \frac{\mathcal{X}}{\mathcal{H}} = \frac{1}{\delta F^2}$  and  $\frac{g\mathcal{H}\mathcal{X}}{vuy} = \frac{g\mathcal{H}}{u^2} \frac{u}{v} \frac{\mathcal{X}}{y} = \frac{1}{R_u^2 F^2}$ , we finally get:

$$\begin{cases} \bar{h}_{\bar{t}} + (\bar{h}\bar{u})_{\bar{x}} + (\bar{h}\bar{v})_{\bar{y}} = 0, \\ \bar{u}_{\bar{t}} + \bar{u}\bar{u}_{\bar{x}} + \bar{v}\bar{u}_{\bar{y}} + \frac{1}{F^2} (\bar{h} + \bar{\phi})_{\bar{x}} = \frac{1}{\delta F^2} \left( -J_0 \frac{\bar{u}\sqrt{\bar{u}^2 + R_u^2\bar{v}^2}}{\bar{c}^2\bar{h}^p} - l_0\bar{b}_{\bar{x}} \right), \\ \bar{v}_{\bar{t}} + \bar{u}\bar{v}_{\bar{x}} + \bar{v}\bar{v}_{\bar{y}} + \frac{1}{R_u^2 F^2} (\bar{h} + \bar{\phi})_{\bar{y}} = -\frac{J_0}{\delta F^2} \frac{\bar{v}\sqrt{\bar{u}^2 + R_u^2\bar{v}^2}}{\bar{c}^2\bar{h}^p}. \end{cases}$$