

L'inversion du tourbillon potentiel

Algorithme et applications

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Mars 2009

Ertel (1942) Potential Vorticity I

$$DP/Dt = 0$$

where

$$P = \frac{\zeta_3 \nabla_3 \theta}{\rho^*}$$

$$\rho^* P = -\frac{\partial v}{\partial z^*} \frac{\partial \theta}{\partial x} + \frac{\partial u}{\partial z^*} \frac{\partial \theta}{\partial y} + (f + \zeta) \frac{\partial \theta}{\partial z^*} \quad (1)$$

If the wind field $\vec{v}_g = \vec{v} - \vec{v}_a$ where \vec{v}_g is the geostrophic wind and \vec{v}_a the ageostrophic one.

Thermal-wind balance :

$$f \frac{\partial v_g}{\partial z^*} = \frac{g}{\theta_0} \frac{\partial \theta}{\partial x} \quad \text{and} \quad f \frac{\partial u_g}{\partial z^*} = -\frac{g}{\theta_0} \frac{\partial \theta}{\partial y}$$

Ertel (1942) Potential Vorticity II

$$\theta(x, y, z^*, t) = \bar{\theta}(z^*) + \tilde{\theta}(x, y, z^*, t)$$

Therefore :

$$\begin{aligned} \rho^* P = (f + \zeta) \left(\frac{\partial \bar{\theta}}{\partial z^*} + \frac{\partial \tilde{\theta}}{\partial z^*} \right) - \frac{f \theta_0}{g} \left[\left(\frac{\partial u}{\partial z^*} \right)^2 + \left(\frac{\partial v}{\partial z^*} \right)^2 \right] \\ + \frac{f \theta_0}{g} \left[\frac{\partial u}{\partial z^*} \frac{\partial u_a}{\partial z^*} + \frac{\partial v}{\partial z^*} \frac{\partial v_a}{\partial z^*} \right] \quad (2) \end{aligned}$$

Potential Vorticity scaling I

$$\mathcal{M} = [\mathcal{M}]\mathcal{M}'$$

$[\mathcal{M}]$ is a scaling of the corresponding variable and \mathcal{M}' the variable without dimension. The scaling for the subset of independent variables is :

- ▶ $[U] = U$
- ▶ $[x] = [\delta x] = [y] = [\delta y] = L,$
- ▶ $[z^*] = [\delta z^*] = H,$
- ▶ $N^2 = g/\theta_o [\partial\theta/\partial z^*]$ where N^2 is the brunt-vaissala frequency of the reference flow.

Then,

- ▶ $[\zeta] = U/L$
- ▶ $[\partial U/\partial z^*] = U/H$
- ▶ $[\partial\tilde{\theta}/\partial z^*] = f\theta_o LU/gH^2.$

Potential Vorticity scaling II

Using the Rossby $R = U/fL$ and Froude $F = U/NH$ numbers the non-dimensional form of the Potential Vorticity equation is :

$$\begin{aligned} \frac{g}{f\theta_0 N^2} \rho^{*'} P' &= \frac{\partial \bar{\theta}'}{\partial z^{*'}} + F^2 R^{-1} \frac{\partial \tilde{\theta}'}{\partial z^{*'}} + R \zeta' \frac{\partial \bar{\theta}'}{\partial z^{*'}} + F^2 \zeta' \frac{\partial \tilde{\theta}'}{\partial z^{*'}} \\ &- F^2 \left(\frac{\partial u'}{\partial z^{*'}} \right)^2 + F^2 R \frac{\partial u'}{\partial z^{*'}} \frac{\partial u'_a}{\partial z^{*'}} \\ &- F^2 \left(\frac{\partial v'}{\partial z^{*'}} \right)^2 + F^2 R \frac{\partial v'}{\partial z^{*'}} \frac{\partial v'_a}{\partial z^{*'}}. \end{aligned} \quad (3)$$

small Rossby and $F \simeq R^{1/2}$ |

$$\begin{aligned} \frac{g}{f\theta_0 N^2} \rho^{*'} P' &= \frac{\partial \bar{\theta}'}{\partial z^{*'}} + \frac{\partial \tilde{\theta}'}{\partial z^{*'}} + R^1 \zeta' \frac{\partial \bar{\theta}'}{\partial z^{*'}} + R^1 \zeta' \frac{\partial \tilde{\theta}'}{\partial z^{*'}} \\ &- R^1 \left(\frac{\partial u'}{\partial z^{*'}} \right)^2 - R^1 \left(\frac{\partial v'}{\partial z^{*'}} \right)^2. \end{aligned} \quad (4)$$

The R^1 expansion of the left-hand side of this equation is now non-linear.

$\phi' = f(\theta')$ hydrostatic relation

$\phi' = \psi'$ geostrophic relation

(5)

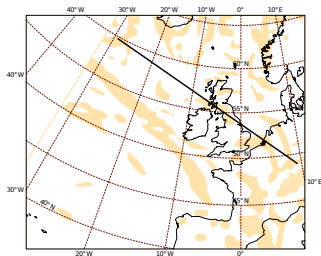
small Rossby and $F \simeq R^{1/2}$ II

Ertel PV is now proportional to the quantity Q where :

$$Q = (f^2 + \nabla^2 \phi) \phi_{zz} - \phi_{xz}^2 - \phi_{yz}^2$$

which belongs to the Monde-Ampere family.

ellipticity criterion I



The shaded area corresponds to atmosphere columns with potential vorticity values less than -0.01 pvu

Linear PV inversion without specifying any balance conditions between velocity and mass I

at small Rossby and Froude/ $O(1)$ expansion :

$$\frac{g}{f\theta_0 N^2} \rho^{*'} P' = \frac{\partial \bar{\theta}'}{\partial z^{*'}} + R^1 \frac{\partial \tilde{\theta}'}{\partial z^{*'}} + R^1 \zeta' \frac{\partial \bar{\theta}'}{\partial z^{*'}}. \quad (6)$$

With physical variables :

$$P = -g \left((f_0 + \zeta) \frac{\partial \bar{\theta}}{\partial p} + f_0 \frac{\partial \theta}{\partial p} \right)$$

For a single layer defined by two isentropic surfaces with uniform density ($\Delta\theta = \theta_U - \theta_b$) whose upper boundary height ϕ_U or pressure U is allowed to vary whereas the bottom ϕ_B/B is fixed.

Linear PV inversion without specifying any balance conditions between velocity and mass II

Δp is defined as the mean depth of the fluid and δp a small amplitude deviation :

$$P = -g \left(\zeta \frac{\Delta\theta}{\Delta p} + f_0 \frac{\Delta\theta}{\Delta p + \delta p} \right)$$

At the first order :

$$P = -g \frac{\Delta\theta}{\Delta p} \left(f_0 + \zeta - f_0 \frac{\delta p}{\Delta p} \right)$$

$$P = -g \frac{\Delta\theta}{\Delta p} \left(f_0 + \frac{\Delta\psi}{f_0} - f_0 \frac{\phi}{\Delta\phi} \right)$$

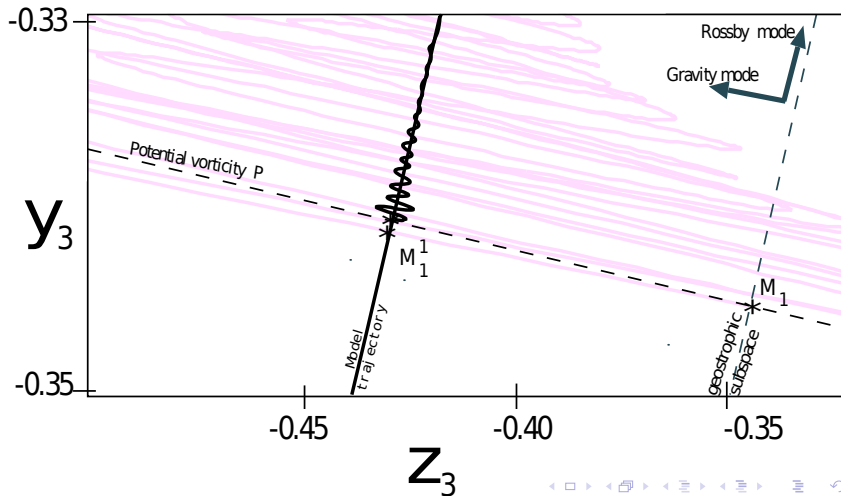
Linear PV inversion without specifying any balance conditions between velocity and mass III

The variable χ is expanded as a sum of $\chi_k e^{ikx}$ functions, the k^{th} component of the expansion of P following the same way writes :

$$P_k = -g \frac{\Delta\theta}{\Delta p} \left(f_0 - \frac{-k^2 \psi_k}{f_0} - f_0 \frac{\phi_k}{\Delta\phi} \right)$$

Lorenz (1980) model which is based on 3 wave numbers.

Linear PV inversion without specifying any balance conditions between velocity and mass IV



Linear PV inversion using a minimum energy constraint ; A variational approach of PV inversion I

$$P = -f_0 g \frac{\partial \theta}{\partial p} - g \frac{\partial \bar{\Theta}}{\partial p} \zeta \quad (7)$$

The heart of the inversion is based on the minimization of :

$$J = \int_{\Omega} \left\{ \frac{1}{2} \zeta \Delta^{-1} \zeta - \sigma \frac{1}{2} \frac{\partial \bar{\Theta}^{-1}}{\partial p} \theta^2 \right\} d\Omega - \int_{\Omega} \Lambda \left(-f_0 g \frac{\partial \theta}{\partial p} - g \frac{\partial \bar{\Theta}}{\partial p} \zeta - P \right) d\Omega$$

$$\sigma = -\frac{R}{p} \left(\frac{p}{p_0} \right)^{R/C_p}$$

Linear PV inversion using a minimum energy constraint ; A variational approach of PV inversion II

The saddle point of J obeys to :

$$\frac{\partial J}{\partial \Lambda} = 0, \quad \frac{\partial J}{\partial \theta} = 0, \quad \frac{\partial J}{\partial \zeta} = 0 \quad \text{and} \quad \frac{\partial J}{\partial D} = 0$$

leading to the following Euler-Lagrange equations :

$$P = -f_0 g \frac{\partial \theta}{\partial p} - g \frac{\partial \bar{\Theta}}{\partial p} \zeta \quad (9)$$

$$\zeta = g \frac{\partial \bar{\Theta}}{\partial p} \Delta \Lambda \quad (10)$$

$$\theta = g f_0 \sigma^{-1} \frac{\partial \bar{\Theta}}{\partial p} \frac{\partial \Lambda}{\partial p} \quad (11)$$

Linear PV inversion using a minimum energy constraint ; A variational approach of PV inversion III

Eqs.(9)-(11) can be rewritten as follows :

$$\frac{\Delta\Lambda}{f_0} + f_0 \frac{\partial \bar{\Theta}^{-2}}{\partial p} \frac{\partial}{\partial p} \left(\sigma^{-1} \frac{\partial \bar{\Theta}}{\partial p} \frac{\partial \Lambda}{\partial p} \right) = - \frac{1}{f_0 g^2} \frac{\partial \bar{\Theta}^{-2}}{\partial p} P \quad (12)$$

A variational approach of non-linear PV inversion I

$$P = -f_0 g \frac{\partial \theta}{\partial p} - g \frac{\partial \bar{\Theta}}{\partial p} \zeta + \text{non-linear terms} \quad (13)$$

$$J = \int_{\Omega} \left\{ \frac{1}{2} (\zeta - \zeta_i) \Delta^{-1} (\zeta - \zeta_i) - \sigma \frac{1}{2} \frac{\partial \bar{\Theta}^{-1}}{\partial p} (\theta - \theta_i)^2 \right\} d\Omega \\ - \int_{\Omega} \Lambda \left(-f_0 g \frac{\partial \theta}{\partial p} - g \frac{\partial \bar{\Theta}}{\partial p} \zeta - P \right) d\Omega. \quad (14)$$

Applications à l'expertise humaine

Assimilation directe



Observations
canaux « vapeur d'eau »
METEOSAT 8
Ozone...

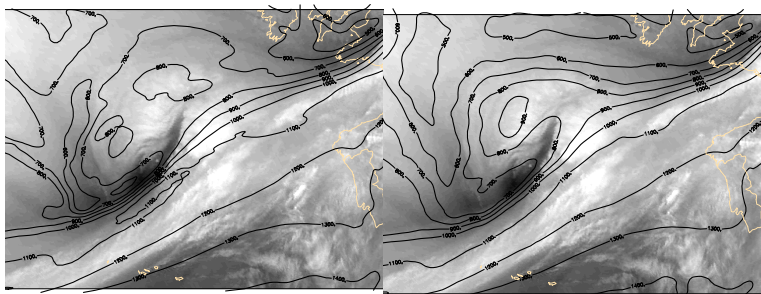
Variables du modèle
Vent, température
Humidité



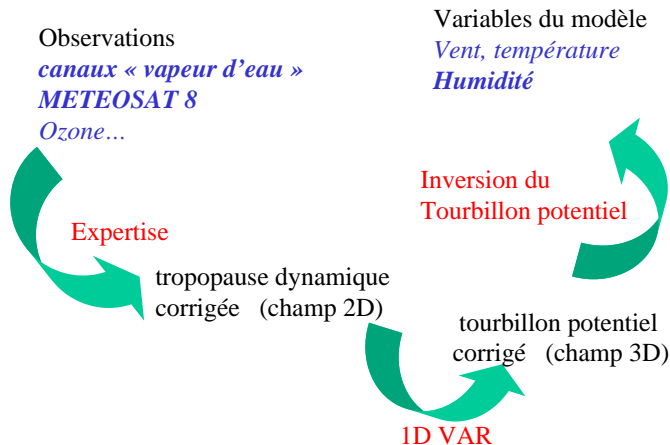
Expertise puis correction

Applications à l'expertise humaine

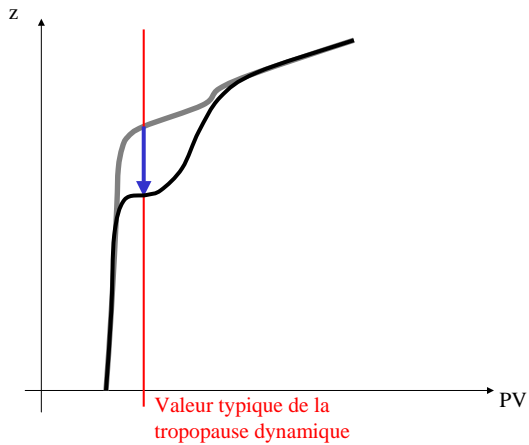
A gauche : avant modifications. A droite : après modifications



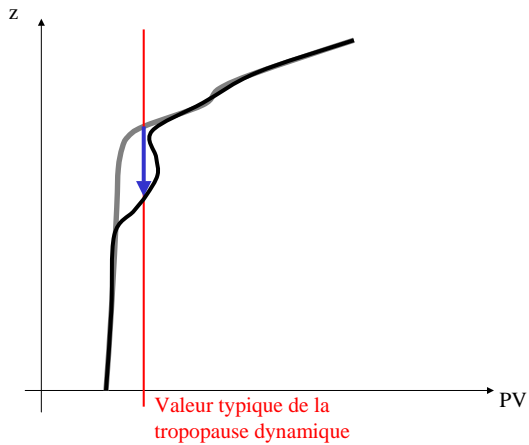
Applications à l'expertise humaine



Reconstruction de profil : ce que l'on veut



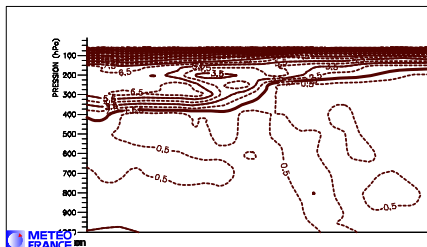
Après analyse 1DVAR d'une observation



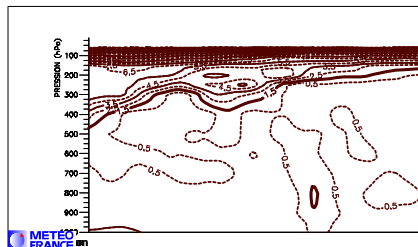
Applications à l'expertise humaine

A gauche : avant modifications. A droite : après modifications

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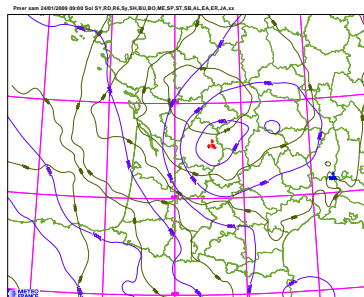
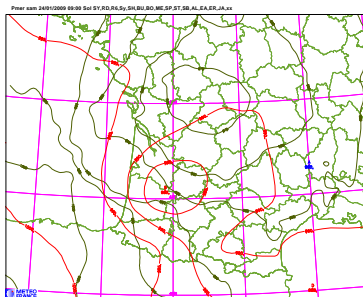


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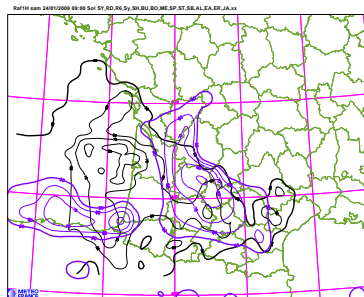
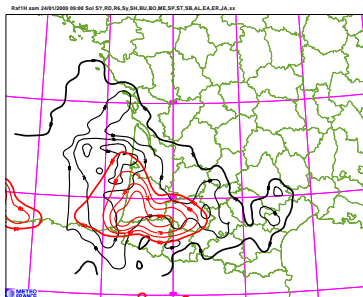
Applications à l'expertise humaine

PMER (à gauche : avant modifications. A droite : après modifications)



Applications à l'expertise humaine

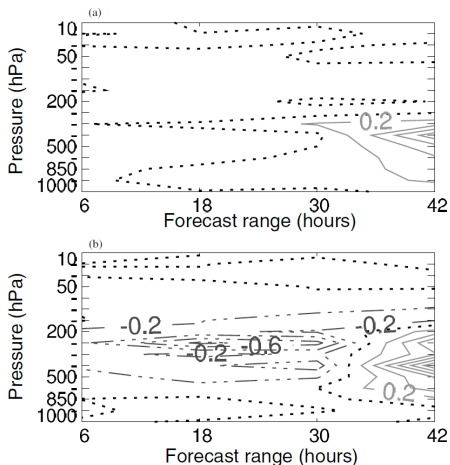
Rafales (à gauche : avant modifications. A droite : après modifications)



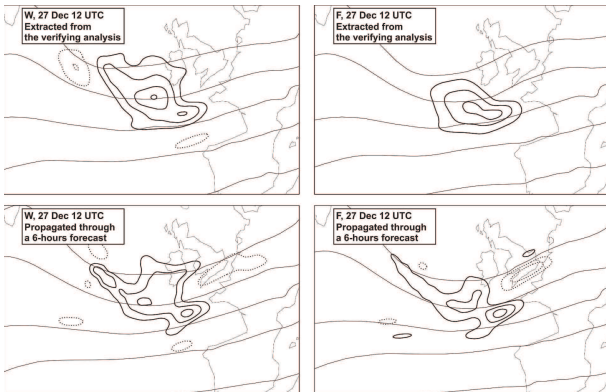
Assimilation of corrections as pseudo-observations within 4DVAR I

PV Correction considered as observations and assimilated (Guérin et al. 2005) Main point : estimation of observation error

Assimilation of corrections as pseudo-observations within 4DVAR II



Coherent structure depiction using wavelets (Plu et al. 2008) I



Perspectives I

- ▶ Non-linéarités
- ▶ Eventuels problèmes de conditionnement
- ▶ Quelles approches pour traiter les forts gradients ?